

PROGRESS REPORT

2010 ESKAY (SIB) & MITCHELL PROPERTY EXPLORATION

Eskay Creek Camp, Northwestern British Columbia

**Eskay: Latitude 56° 35' N
Longitude 130° 29' W
NTS 104B 9E and 10E**

FOR

ESKAY MINING CORP.

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EXECUTIVE SUMMARY

This report provides a summary of the work carried out by Cambria personnel on the Eskay Property and the Mitchell Property from July 18 to October 1 2010, plus recommendations for future work on the properties.

Location: The mineral properties of Eskay Mining Corp. are located in northwestern British Columbia, 70 kilometres northwest of Stewart, B.C. The reference map sheets are NTS sheets 104B9 and 104B10. The Eskay Creek Area properties, including **the SIB claim block on the Eskay (SIB) Property** and the contiguous **Corey Property** surround and abut the Eskay Creek mine property of Barrick Gold Corporation and the past producing Eskay Creek mine.

Mineral Tenures: Eskay Mining Corp. holds an interest in 177 mineral claims over 44,750 hectares in the region. In detail, Eskay Mining Corp. holds an up to 80% interest over 33,000 hectares (the Eskay (SIB) Property and the Mitchell claims) via an option agreement with St. Andrew Goldfields Ltd. **Eskay Mining Corp. has a 100% interest in the remaining 37 claims** (the Corey Property), subject to a 2% net smelter return royalty. All claims are currently in good standing.

The Eskay property: The property adjoins Barrick's prolific past-producing Eskay Creek mine property. The principal target on Kenrich's Eskay Property is the Lulu Zone, a gold, silver and base metal-enriched zone of stringer and semi-massive sulphides having the same geochemical and geological characteristics as the Eskay deposit. The property also encompasses the Hexagon Zone, which is a large hydrothermal system most likely related to Volcanogenic Massive Sulphide (VMS) type mineralization.

The 2010 Eskay Property exploration program comprised:

- approximately 6 km x 1 km of geological mapping and lithogeochemical sampling; mapping was conducted at 1:5000 scale.
- 5 drillholes for 3856.7 metres of drilling (an additional 44.2 metres were drilled in hole "138A" which had to be abandoned and restarted as Hole 138).

The 2010 Mitchell Property exploration program comprised:

- Approximately 1.5 km x 0.75 km of reconnaissance geological mapping and geochemical sampling

Diamond drilling in 2010 concentrated on the extension of the Lulu Zone host rocks on the footwall side of the Coulter Creek Thrust Fault (CCTF). All drillholes have successfully intercepted favourable Salmon River Formation stratigraphy, including Eskay-type rhyolites and basalts. Two of the holes intercepted mineralization with Eskay-like geochemical characteristics (highly anomalous in Au, Ag, Zn, Pb, As and Sb).

Hole EK10-137 intersected the same stockwork mineralization hosted in rhyolite. **A 10.0 metre intercept from 463.0 to 473.0 metres depth returned assays of**

0.65 g/t Au, 0.26% Pb, 0.38% Zn and 6 g/t Ag. This includes a 0.5 metre intercept from 472.5 to 473.0 metres depth that returned 2.35 g/t Au, 0.97% Pb, 1.89% Zn and 25 g/t Ag. The mineralization is very similar in style to that intersected in Hole EK08-134 located 75 metres to the south. This extends the zone of stockwork-mineralized rhyolite and its location above the CCTF in this hole suggests that the mineralization in Hole 137 and in Hole 134 both occur in the fault's hangingwall block.

Hole EK10-138 intersected a sequence of interlayered Salmon River Formation rhyolite and basalt below the CCTF. Here, a 2-metre interval contained 5% pyrite-sphalerite-galena stringers within a sheared and veined argillite unit. Within this, a **0.9 metre interval from 556.9 to 557.8 metres depth returned assay grades of 0.06g/t Au, 0.2% Pb, 0.5% Zn and 5 g/t Ag.**

It is apparent that the Eskay Rift succession in the footwall to the CCTF is not an exact duplicate of the succession at the Eskay Creek mine as rhyolites occur both above and below the basalts. The basalt, suspected to lie stratigraphically above the Lulu Zone position, thus provides a potentially important marker unit to one's position in the rock sequence. Further to this, holes 139 & 140 (drilled from opposite directions on the same section south of the Lulu Zone) reveal that the Salmon River Formation in the footwall to the CCTF is at least 500 metres thick, a far greater thickness than is apparent from outcrop in the fault's hangingwall.

A continuing coordinated, systematic program of diamond drilling, geological mapping, lithogeochemical sampling and geophysical data modeling is recommended, including an IP survey, utilizing the in-drillhole electrodes already in place, to identify targets for future drill testing. Diamond drilling should primarily focus on testing the Eskay Rift rocks beneath the CCTF. Given that the basalts encountered in the footwall of the CCTF by the 2010 drilling likely lie stratigraphically above the Lulu Zone position, those parts of the footwall panel that are downsection from the basalts are priority targets and should be tested for Lulu Zone-equivalent mineralization. The laterally extensive zone of Eskay-type rhyolite-hosted stockwork mineralization in holes EK08-134 and EK10-137 should also be followed up and down dip.

INTRODUCTION

The mineral properties of Eskay Mining Corp. (the Company) are located in northwestern British Columbia, 70 kilometres northwest of Stewart, B.C (see Figure 1). The reference map sheets are NTS sheets 104B9 and 104B10. The Eskay (SIB) property and the contiguous Corey Property surround and abut the Eskay Creek mine property of Barrick Gold Corporation and the past producing polymetallic, precious metal rich Eskay Creek mine.

The Company holds an interest in mineral tenures comprising 177 claims over 44,750 hectares in the region. In detail, the Company holds a 100% interest in 37 mineral tenures comprising 11,500 ha (the Corey Property), subject to a 2% Net Smelter Return royalty. On May 8, 2008 the Company announced an option agreement with St Andrew Goldfields Ltd. (SAS) to acquire up to an 80% undivided interest in 33,000 hectares of mineral tenures that include the historic SIB/Lulu areas (the Eskay (SIB) Property) and claims east of Seabridge's Iron Cap Zone (the Mitchell Property). All mineral tenures are currently in good standing.

Acquisition of the St Andrew properties allows Eskay Mining Corp. to expand its ongoing exploration in the region and apply its accumulated geological expertise onto this highly prospective new ground closer to the Eskay Creek mine. Significantly, portions of the St Andrew property lie within 1 kilometer of the Eskay deposit. Of greatest immediate significance to Kenrich is the SIB claim block. The SIB block adjoins Barrick's Eskay Creek mine property and contains a continuous succession of Eskay Rift rhyolite and mudstone, the host units to the Eskay deposits, plus the Lulu Zone, a gold, silver and base metal-enriched zone of stringer and semi-massive sulfides having the same geochemical and geological characteristics as the Eskay deposit (the best drill intercept at Lulu returned a value of 14.43 g/t Au over 14.3 metres; McGuigan, 2002) and the Hexagon Zone, a large Volcanogenic Massive Sulfide (VMS) style hydrothermal system on the east side of the SIB claims (McGuigan, 2002).

The Company's mining properties are accessed by helicopter from the Eskay Mine access road that extends from Highway 37 to the Eskay Mine. Staging areas for helicopter operations are located at a fuel cache located along the Eskay Creek Mine road, about five kilometers west from the mine. Additionally, well serviced helicopter pads and a fueling station are located at the nearby Bell II Lodge located on Highway 37 east of the Company's properties.

Valley bottoms are densely forested with mature stands of fir, Sitka spruce, cedar, hemlock, aspen, alder, and maple. A thick undergrowth of ferns, salmonberry, huckleberry and devil's club is usually present.

The Eskay Property's SIB claim block and parts of the Corey Property are located within the Unuk River watershed. Major tributaries include the Unuk and South Unuk Rivers, Sulphurets Creek, Coulter Creek and Storie Creek. All rivers and creeks originate from glacial melt waters, and reach peak flow conditions in the summer months. The region is mountainous with elevations ranging from 250 meters on the Unuk River to approximately 2,150 meters at John Peaks. Mountain slopes are moderate to very steep. The tree line occurs at about 1,200 meters and at higher



elevations valleys are generally filled with glaciers. Semi-permanent ice and snow may be encountered on north facing slopes. Snow conditions are extreme in alpine areas while river bottom areas receive snow seasonally. However, precipitation in the form of rain occurs all year round.

The following report summarizes the work carried out on the Eskay and Mitchell Properties from July 18 to October 1 2010, and provides recommendations for continuing work on the property. Sean McKinley, M.Sc., P.Geo. supervised the project (the project's Qualified Person), Stephen Tennant Ph.D. directed day-to-day field operations, Alanna Ramsay B.Sc., GIT logged the majority of the drillcore and Christopher Sebert P.Eng. carried out geological mapping and outcrop sampling. Falcon Drilling Ltd. of Prince George, B.C. undertook the diamond drilling.

PROPERTY DESCRIPTION AND LOCATION

The subject properties are located in northwestern British Columbia, 70 km northwest of Stewart and 900 kilometres northwest of Vancouver (Figure 1). Reference maps are NTS Sheets 104B 9W and 10E. The properties are centered at approximately 56 degrees 35 minutes north and 130 degrees 29 minutes west.

The Eskay/SIB properties abut and surround the past-producing Eskay Creek gold mine, owned by Barrick Gold Corporation who ceased production in 2008. Portions of the Eskay property are less than 1 kilometre from the Eskay Creek mine.

The Mitchell Property is non-contiguous with the main Eskay Property and is located 20 kilometres southeast of the Eskay Creek Mine. It lies east of the Mitchell and Iron Cap properties of Seabridge Gold and northeast of the Snowfield Property of Silver Standard.

A complete listing of the mineral tenures that comprise the Eskay & Mitchell properties is included in Appendix C.

The mining properties of Eskay Mining Corp. are accessed by helicopter from the Eskay Mine access road that extends from Highway 37 to the Eskay Mine. Staging areas for helicopter operations are located at a fuel cache located along the Eskay Creek Mine road, about five kilometers west from the mine. Additionally, well serviced helicopter pads and a fueling station are located at the nearby Bell II Lodge located on Highway 37 east of the Eskay properties.

Valley bottoms are densely forested with mature stands of fir, Sitka spruce, cedar, hemlock, aspen, alder, and maple. A thick undergrowth of ferns, salmonberry, huckleberry and devil's club is usually present.

The Eskay property area is located within the Unuk River watershed. Major tributaries include the South Unuk River and Sulphurets Creek. The Mitchell Property lies on and adjacent to the Mitchell Glacier at the headwaters of Mitchell Creek which flows westward to Sulphurets Creek. All rivers and creeks originate from glacial melt waters, and reach peak flow conditions in the summer months. The region is mountainous with elevations ranging from 250 metres on the Unuk River to

approximately 2,150 metres at John Peaks. Mountain slopes are moderate to very steep. The tree line occurs at about 1,200 metres and at higher elevations valleys are generally filled with glaciers. Semi-permanent ice and snow may be encountered on north facing slopes. Snow conditions are extreme in alpine areas while river bottom areas receive snow seasonally. However, precipitation in the form of rain occurs all year round.

BACKGROUND – ESKAY CREEK CAMP

Eskay Creek Mine – Barrick Gold Corp.

The **Eskay Creek Mine**, operated by Barrick Gold, is located in northwestern British Columbia, 75 km northwest of Stewart, B.C. The property is accessed from Highway 37 and the nearby Eskay Creek Mine road. The Eskay property is 10 km from the northern border of the Corey property of Kenrich.

The mine property contains several deposits of gold- and silver-rich polymetallic sulfide and sulfosalt mineralization as volcanogenic and replacement massive sulfide, debris flow breccias, and discordant veins and stockworks.

The Eskay Creek deposits are examples of shallow subaqueous hot spring deposits, an important new class of submarine mineral deposits that has only recently been recognized in modern geological environments. They are relatively under explored and poorly recognized within the geological record. The deposit type is transitional between subaerial hot spring Au-Ag deposits and deeper water, volcanogenic massive sulfide exhalites (Kuroko or Besshi types) and shares the mineralogical, geochemical, and other characteristics, of both (see Roth, 2002).

Exceptionally gold-rich mineralization was discovered in 1989, when a company promoted by Murray Pezim, Calpine Resources, intersected **208 metres grading 27.2 g/t gold and 30.2 g/t silver** in diamond drillhole 109. The Eskay Creek mine commenced production in 1994. The ore was initially shipped directly to smelters with no milling or concentrating. A mill was established only in 1998.

Most of the initial reserves at Eskay were defined in the 21B zone, which is hosted in Lower to Middle Jurassic volcanic and sedimentary rocks of the Salmon River formation. The zone forms a lens-shaped body measuring 900m by 300m by 20m thick. The mineralization occurs as a stratabound sheet in carbonaceous mudstones of the Contact Mudstone unit and in feeder veins in the underlying Eskay Rhyolite. Based on mineral associations and continuity of grade, the 21 zone has been divided into two deposits: the 21A and the 21B. These deposits are separated by 140 metres of weak mineralization. Diamond drilling has traced the entire zone for 1.4 km along strike and 250 metres down dip over widths of 5-45 metres.

The exploration success continued. In 1995, drilling intersected the NEX and Hangingwall zones. The NEX lies north of the 21B lens, along the same stratigraphic horizon, and consists of mainly massive sphalerite, tetrahedrite, galena and lesser lead-sulphosalts, with late chalcopyrite stringers crosscutting the lens. The

Hangingwall zone is stratigraphically above the NEX zone, generally above the first basaltic sill, and dominated by pyrite, sphalerite, galena and chalcopyrite.

In 2002, one of two holes drilled into the historic **22 zone**, 2 kilometres south of the mine, yielded **6.2 grams gold over 80.1 metres, including a higher-grade section running 64.1 grams gold over 4.7 metres**. Mineralization encountered in the 22 zone includes both discordant stockworks and stratiform VMS mineralization similar to the 21B zone.

Eskay Rift Setting: The Eskay-Corey Belt

Eskay Creek-type mineralization is a stratabound assemblage of volcanogenic massive sulfide mineralization and stockwork vein systems with local high-grade gold-silver replacement mineralization that was deposited in a shallow, sub-aqueous epithermal hot spring environment. This mineralization is closely related to an assemblage of rift-related volcanic and sedimentary rocks and to controlling fault structures that bound and crosscut the local rift basins. Metallogenic studies by the Mineral Deposit Research Unit (MDRU), and federal and provincial government geological survey branches have determined the Eskay Creek mine sequence is a Lower to Middle Jurassic succession of bi-modal volcanism and clastic sedimentation, termed the Salmon River Formation, a sub-division of the regional Hazelton Group.

Barrett and Sherlock (1996) argue on the basis of lithogeochemistry that the Eskay rhyolite most closely resembles rhyolites erupted at rifted continental margin and are significantly different from the arc related volcanic rocks that compose the rest of the Hazelton Group. The hanging wall basalt unit yields a mainly N-MORB composition. These arguments, together with observed or inferred facies variations in the immediate Eskay Creek area, led Barrett and Sherlock (1996) and Roth (2002) to suggest that the Eskay Creek deposit formed within a roughly north-south trending zone of localized rifting, either in a back-arc or an inter-arc paleotectonic setting, that represents the terminal stage of magmatism within the Hazelton Group.

Work by Kenrich from 2003-06 has further defined the paleotectonic setting of the Eskay Camp, and the important Eskay rift. In the Technical Report for Kenrich-Eskay by McGuigan et al., (2004), that includes contributions by Barrett, the paleotectonic setting of the Eskay rift is interpreted on a camp scale, using data in the public domain (scientific papers, assessment reports and MDRU compilations) and data in the private files of Kenrich-Eskay. Distinctive volcanics and sediments define an **Eskay-Corey belt** that contains all the best Eskay-type deposits and significant discoveries in the Eskay region (see Figures 2a and 2b). **The Eskay Property spans the northern portion of this trend and contains mineralization directly analogous to the Eskay deposits.**

REGIONAL GEOLOGY

The rock types present on the Eskay and Mitchell properties for the most part are representative of the regional geology and for the sake of brevity this is summarized below with an emphasis only on those aspects pertinent to the current exploration program. The properties lie within the Intermontane belt of the Canadian Cordillera. They are underlain by a thick succession of Upper Triassic to Middle Jurassic volcano-sedimentary arc-complex lithologies (Stuhini and Hazelton Groups), Permian and older arc and shelf sequences (Stikine Assemblage) and Middle and Upper Jurassic marine basin sediments (Bowser Lake Group).

Geologists of the British Columbia Geologic Survey and the Geological Survey of Canada have subdivided the Jurassic rocks in the area into the Hazelton Group and the Bowser Lake Group (see Anderson and Thorkelson, 1990) with the volcano-sedimentary rocks on the Eskay (SIB) and Corey properties largely belonging to the Lower to Middle Jurassic Hazelton Group (see Figures 2a and 2b). The regional stratigraphic framework was further refined by Peter Lewis and members of the Mineral Deposit Research Unit in the 1990s and the following description follows Lewis's 1996 scheme.

Triassic Stuhini Group

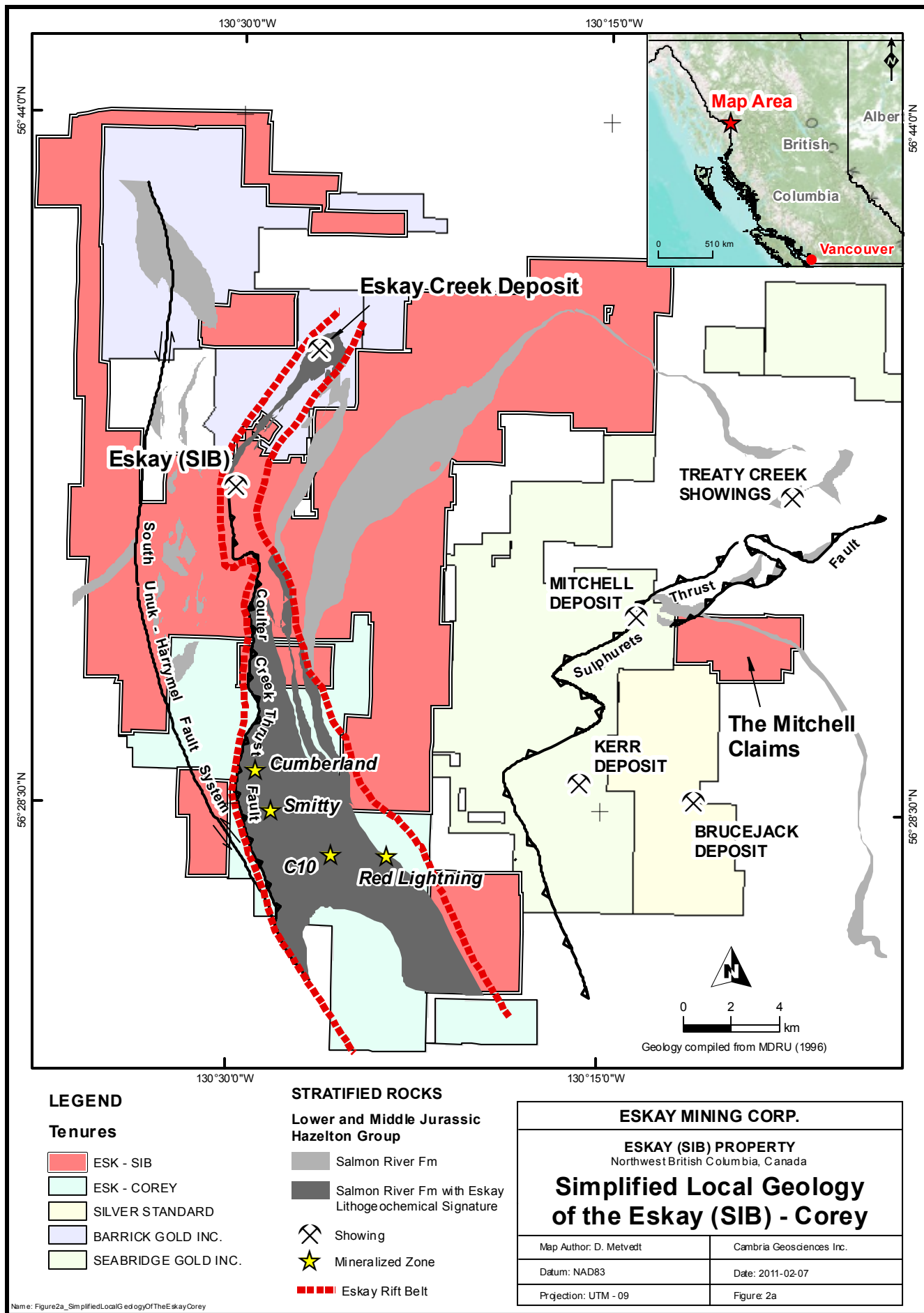
The oldest Mesozoic strata in the region are sedimentary and volcanoclastic rocks of the Triassic Stuhini Group. The Stuhini Group consists of a dominantly sedimentary lower division and a dominantly volcanic and volcanoclastic upper division.

LOWER AND MIDDLE JURASSIC HAZELTON GROUP

The Hazelton Group in northwestern British Columbia records Lower and Middle Jurassic arc volcanism and volcanogenic sedimentation (Alldrick and Britton, 1991).

Jack Formation: Lower Hazelton Group sedimentary strata

Basal Hazelton Group strata typically consist of locally fossiliferous conglomerate, sandstone, and siltstone of the Jack Formation. These rocks are well exposed in the upper Unuk River/Sulphurets area along both limbs of the McTagg anticlinorium and have been traced at least as far south as the Frank Mackie icefield. The most complete and best-exposed sections are located in alpine areas north and south of John Peaks and along the west side of the Jack Glacier, where the unit overlies Stuhini Group strata along an angular unconformity.



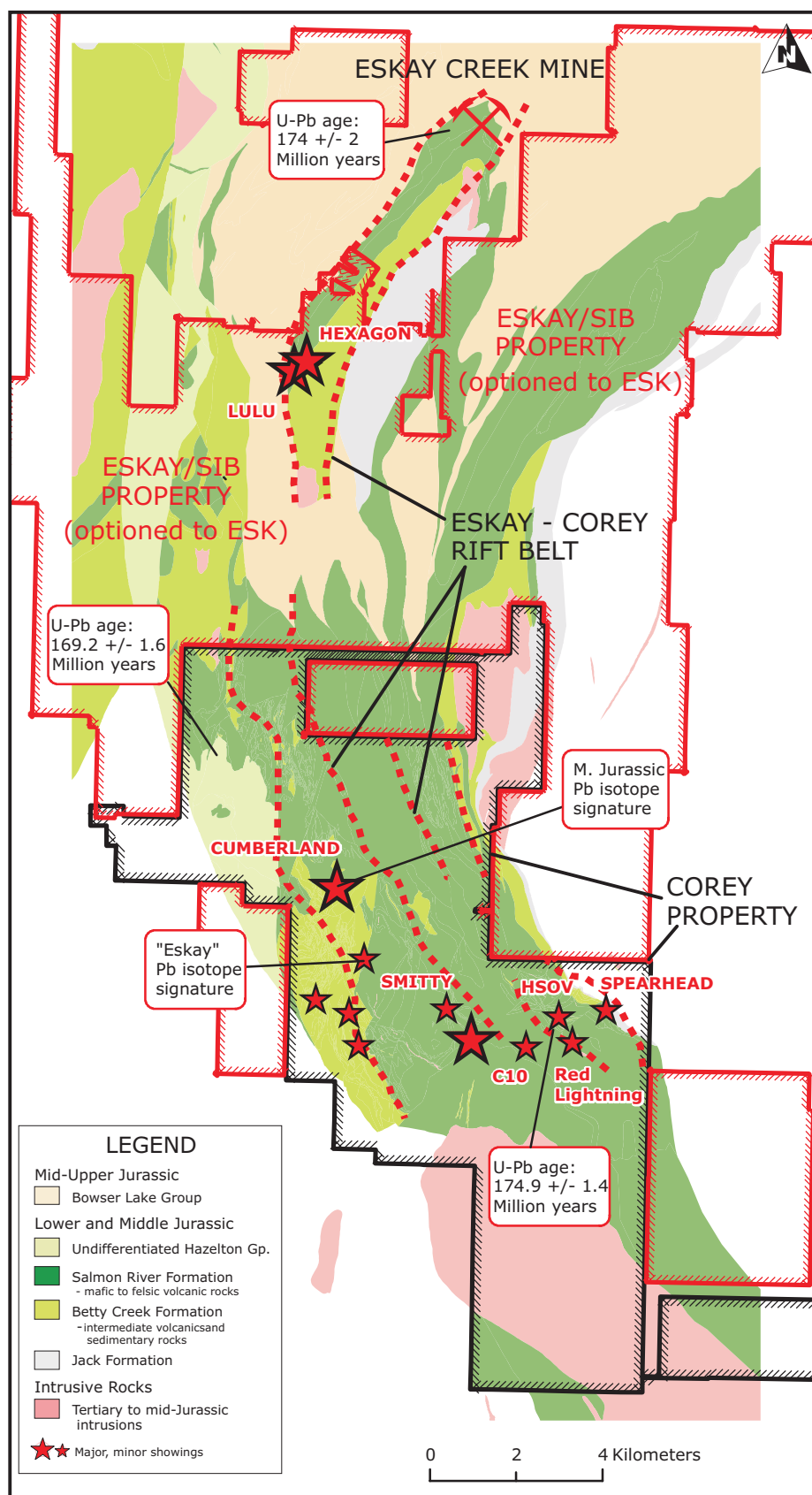


Figure 2b. Simplified local geology of the Eskay-Corey rift belt and mineral showings (limits of rift are shown with red dashes).

Betty Creek Formation: Intermediate composition volcanic and volcanoclastic strata

Lower Jurassic volcanic and volcanoclastic strata have been problematic for workers in the Iskut River area, and stratigraphic nomenclature has been unevenly applied. We assign the entire volcanic and volcanoclastic package from the Jack Formation to a distinct shift to bimodal volcanism in the lower Middle Jurassic, to the Betty Creek Formation intermediate composition volcanic/volcanoclastic sequence. This unit encompasses most of the rocks previously assigned to the Betty Creek and Unuk River Formations, as well as some rocks previously assigned to the Mount Dilworth Formation.

Salmon River Formation: Bimodal volcanic unit

The upper part of the Hazelton Group in the Eskay Creek area comprises dacitic to rhyolitic flows and tuffs, localized interlayered basaltic flows, and intercalated volcanoclastic intervals. This part of the Hazelton Group has attracted the attention of explorationists due to its association with mineralization at Eskay Creek, but at the same time its distribution, internal stratigraphy, and age has often been misunderstood. Previous workers have mapped felsic volcanic components as the Mount Dilworth Formation, and mafic volcanic components as a distinct facies of the Salmon River Formation. However, recent work demonstrates that more than one felsic interval exists in the unit, and that mafic volcanic rocks occur both above and below these felsic intervals (see Lewis, 2001). As such, the term Mount Dilworth Formation is not used herein. Most recently, the Salmon River Formation has been divided into three members: the felsic volcanic-dominated Bruce Glacier Member, the sedimentary Troy Ridge Member and the mafic volcanics of the John Peaks Member (again, see Lewis, 2001). An additional felsic member, the Eskay Rhyolite (see below), has also been identified, but it is generally directly spatially associated with the Eskay Deposit itself and is likely a sub-member of the Bruce Glacier Member.

Bruce Glacier Member: Felsic volcanic rocks are ubiquitous in the Salmon River Formation in the Eskay Creek area. Two felsic members are recognized. Most widespread in its distribution is the Bruce Glacier member, which ranges from a few tens of meters to a few hundred meters in thickness. Lithofacies within the Bruce Glacier member are highly variable both regionally and vertically in a given section. Rocks proximal to extrusive centres include banded flows, massive domes with carapace breccias, autoclastic megabreccias, and block tuffs. Variably welded lapilli to ash tuffs characterize more distal equivalents. Reworked tuffs locally form thick epiclastic accumulations and may infill paleobasins adjacent to extrusive centres.

Eskay Rhyolite: Within and adjacent to the Eskay Creek deposit, a rhyolite with anomalously low titanium content has been separated as a distinct member of the Salmon River Formation, termed the Eskay Rhyolite. Early work concluded the member was distinct from the Bruce Glacier member however the whole rock lithogeochemistry is similar to those parts of the Bruce Glacier member that are proximal to the deposit.

Troy Ridge Member: Lithotypes present in this member include thinly-bedded carbonaceous mudstone, and interbedded turbiditic siltstone/argillite and tuff forming distinctive black and white striped strata ("pajama beds"). These units appear to be relatively abundant on the western flanks of Mount Madge on the Core property.

They commonly form meter to decimeter-scale interbedded with mafic volcanics and, to a lesser extent, felsic volcanics. This is a key unit in the sequence as it likely marks a hiatus, at least locally, in volcanic activity, thus providing an excellent potential environment for Eskay-style massive sulphide formation.

John Peaks Member: Mafic components of the Salmon River Formation are assigned to the John Peaks member. They generally occur above the felsic volcanic rocks, but at Treaty Creek thick sections of mafic flows and breccias lie below felsic welded tuffs. These tuffs are correlated with the Bruce Glacier member. Textures present include massive flows, pillowed flows, broken pillow breccias, and volcanic breccias. The John Peaks Member is generally considered to lie immediately stratigraphically above 'Eskay time'.

MIDDLE JURASSIC BOWSER LAKE GROUP

The cessation of Hazelton Group volcanism in the early Middle Jurassic marks an abrupt shift to siliciclastic sedimentation of the Bowser Lake Group. Bowser Lake Group rocks are widely exposed over a broad region of the northern Cordillera, and concordantly overlap Hazelton Group strata along the northeastern edge of the Eskay Creek project area. They consist primarily of monotonous interstratified thin- to thick-bedded shale, siltstone, wacke, and conglomerate, with the notable absence of a volcanic component. Lowest parts of the sequence contain fossils indicating a Bajocian age, implying little or no gap in deposition from the uppermost Hazelton Group.

INTRUSIONS

Mesozoic intrusive activity in the Stewart-Iskut region occurred in two major intervals: a Late Triassic pulse and an extended period of Early to Middle Jurassic plutonism. MacDonald et al. (1996) propose three major temporal suites of plutonism:

- 1) Late Triassic (228-221 Ma) Stikine Plutonic Suite related to the building of a Late Triassic volcanic arc.
- 2) Early Jurassic (195-190 Ma) Texas Creek Plutonic Suite related to an Early Jurassic volcanic arc that was coeval to the Betty Creek Formation volcanic rocks.
- 3) Early to Middle Jurassic (180-170 Ma) intrusions that are related to the upper division of the Hazelton Group, the Salmon River Formation. These possibly correlate with intrusions of the Three Sisters plutonic suite that occur further west and north.

In the area of the Eskay mine, and on the Eskay property's SIB claim block, mafic dikes and felsic intrusions that are controlled by syn-mineralization faulting are classified with the latest pulse of magmatism. Other intrusions, such as alkali feldspar-plagioclase-hornblende porphyry that are hosted by Betty Creek Formation rocks, are likely related to either the latest pulses of Betty Creek volcanism or to Salmon River volcanism, on the basis of intrusive relationships and composition.

The Eskay Porphyry, which is proximal to the footwall of the 21 Zone at the Eskay mine, is grey-green coloured and plagioclase±K-feldspar±hornblende biotite porphyritic. It is a hypabyssal stock of dacitic to granitic composition and correlates with Early Jurassic magmatism (186.2 Ma, U-Pb [zircon] age, MacDonald, 1992).

STRUCTURAL GEOLOGY

The present distribution of rocks in the Eskay Creek area has been influenced by at least two Mesozoic to Cenozoic deformation events:

Early to Middle Jurassic Deformation

There are several lines of evidence that suggest there was a deformation event that was synchronous with deposition of the Hazelton Group. Certain faults that have been mapped in the region appear to separate blocks of differing volcanic successions. Furthermore, some of these faults have clearly juxtaposed successions of Hazelton Group rocks of differing thicknesses, but do not appear to significantly offset the overlying Bowser Lake Group sedimentary succession. These types of structures are interpreted to be synvolcanic (growth) faults and likely were not active past the last deposition of Hazelton rocks.

The Harrymel Fault is a major brittle structure exposed along the western edge of the project area and is interpreted to pass southward into a broad ductile shear zone referred to as the South Unuk Shear Zone. Kinematic indicators are well exposed in both the brittle and ductile portions of this structure, and consistently show dominantly strike-slip movement with a sinistral sense. U-Pb dating of syntectonic intrusions in the ductile portion of the shear zone indicates that the structure was active in the Middle Jurassic (Lewis, 2001), roughly coincident with or just following cessation of Hazelton Group volcanism.

Cretaceous Contractional Deformation

The Eskay Creek area lies between two regional contractional orogens that were active during Cretaceous time: an extensive westerly-directed system of thrust faulting as along the western side of the Coast Belt, and the east-northeasterly directed Skeena Fold and Thrust Belt of the Bowser Basin. The dominant structures in the project area that relate to these events are major folds and thrust faults.

Contractional structures show a transition from broad open folds in the northern part of the Eskay property to tight folds and thrust faults in the south and on the Corey property. In the north, in the vicinity of the Eskay deposit, thrust faults are rare to non-existent. The distribution of stratigraphic units outlines four major folds; from east to west these are the McTagg anticlinorium, the Unuk River syncline, the Eskay anticline, and the Prout Plateau syncline. Fold scale and geometry varies with stratigraphic level, reflecting the different scale of stratification within the Mesozoic sequence. The well-stratified rocks of the Bowser Lake Group contain abundant open to tight upright folds that are parasitic to major folds while the thicker Hazelton Group rock packages, perhaps with the exception of the interlayered sedimentary members, mainly lack these second order folds.

The widespread development and intensity of the Cretaceous contractional deformation event overprints and obscures earlier-formed structures, and likely reactivated any favorably-oriented pre-existing faults. Both the orientations and relative positions of faults that were active synchronously with Hazelton Group volcanism were strongly modified.

At the SIB claim block on the Eskay Property, this shortening is expressed as folds in the Hazelton and Bowser Lake Group rocks and the Coulter Creek Thrust Fault (CCTF), an important structure along the western boundary of the Eskay property's SIB claim block. It is a gently east-dipping, west-verging fault, thrusting Hazelton Group over Bowser Lake Group strata. The fault has not been observed in outcrop and was first identified by interpreting outcrop mapping and drill core relationships. The magnitude of displacement on the fault has not been accurately determined, but is reported by Lewis (2001) to increase from negligible displacement at the north end of the SIB claim block to several hundred metres of displacement to its south. P.J. McGuigan (pers. comm., 2008), based upon more recent work with Lewis, believes that there is 400 metres of dip-slip displacement on the thrust at the Lulu Zone, plus a minor element of oblique slip movement with an as yet undetermined direction, with dip-slip displacement increasing in magnitude to the south (i.e. a "scissoring" along the thrust plane).

The Eskay-Corey Rift Belt

Barrett and Sherlock (1996) argue on the basis of lithogeochemistry that the Salmon River Formation's Eskay rhyolite most closely resembles rhyolites erupted at rifted continental margin and are significantly different from the arc-related volcanic rocks that compose the rest of the Hazelton Group. This and observed or inferred facies variations in the immediate Eskay Creek area led Barrett and Sherlock (1996) and Roth (2002) to suggest that the Eskay Creek deposit formed within a roughly north-south trending zone of localized rifting. The largely calc-alkaline magmatic affinity Lower to Middle Jurassic Salmon River Formation thus contains a bimodal suite of volcanic rocks of distinctive tholeiitic affinity that define the rift.

Building on this work, the Technical Report by McGuigan et al (2004) for the Company concluded that Eskay-type tholeiitic basalt, and a mixed population of rhyolites (ranging from closely analogous to Eskay Rhyolite to some that are calc-alkaline) occur in a linear, north-south trending belt on the Eskay, SIB and Corey properties. Together they form a distinct Eskay rift sequence and with the accompanying faulting and gold, silver and base metal mineralization form the "Eskay-Corey belt". This belt contains all significant gold and silver occurrences in the Eskay Camp. Calc-alkaline intermediate rocks flank this belt and despite containing time-equivalent members to the Eskay-Corey belt, contain only minor base and precious metal occurrences. This further confirms that the trend of the Eskay rift is the most prospective. Structural repetition by folding and thrust faulting occurred after the formation and mineralization of the Eskay Rift belt. Cretaceous age compressional deformation shortened the strata, forming upright, open to tight folds with axial planes trending NNW. Several important thrust faults of that event include the Coulter Creek Thrust Fault that is now well-defined in the Eskay (SIB) property. The Company has continued with closely spaced lithogeochemical sampling of the volcanic-sedimentary rift succession and focused all exploration work within the Eskay-Corey rift belt.

BACKGROUND TO THE ESKAY (SIB) PROPERTY

The Eskay Creek Mine Deposits

The nearby Eskay Creek deposit was discovered in 1989 by junior mining companies listed on the Vancouver Stock Exchange (now the TSX-Venture Exchange). The current owner, Barrick, closed the mine in early 2008. The mine contains several deposits of exceptionally gold-silver rich polymetallic sulphide and sulfosalt mineralization as volcanogenic and replacement massive sulphide ("VMS"); as debris flow breccias; and as discordant veins and stockworks. The Eskay Creek deposits were formed in a shallow submarine hot-spring environment, and are termed high-sulphidation volcanogenic massive sulphide deposits (Roth, 2002). Like most VMS deposits, they consist of semi-massive to massive, concordant sulphide lenses underlain by discordant stockwork feeder zones. They have diverse geochemical signatures dominated by Au, Ag, Cu and Zn and often accompanied by elevated concentrations of As, Sb, Pb, Te and Hg. Cumulative production at Eskay Creek, until closure in early 2008, was 102.00 tonnes of gold and 4,995.24 tonnes of silver (3,279,415 oz gold, 160,597,110 oz silver) from 2,238,255 tonnes of production milled. The grade of production was an exceptional 45.57 g/t gold and 2,231 g/t silver (1.33 oz/ton gold and 65.1 oz/ton silver) over the mine's life.

Host rocks at the Eskay Creek deposit comprise tholeiitic and transitional rhyolites, tholeiitic basalt and carbonaceous mudstones. Most of the ore grade deposits formed with the "contact mudstone" that lies on the rhyolite and within small basinal depressions. The hanging wall is comprised of basalt flows, mudstones and basalt sills. Eskay stratabound deposits are localized over footwall alteration zones and syn-mineralization faults of northwesterly and northerly trends.

The Eskay/SIB Property Deposits

In the same period of discovery as the Eskay Creek deposit, drilling intersected significant gold-silver mineralization at the Lulu zone, located on the SIB mineral claims. Lulu is hosted by up to 170 metres of the mineralized "Marguerite mudstone", the lowermost mudstone that is interbedded with Salmon River Formation felsic volcanics. The alteration comprises extensive sericite and locally intense pervasive silicification and sodium metasomatism. Below an extensive interval of silicified and albitized felsic strata, drill hole 90-30 intersected 21 metres of black siliceous carbonaceous mudstone. A 14 metre thick interval of the mudstone is mineralized with disseminated pyrite, framboidal pyrite, laminar pyrite and disseminated and fracture-controlled stibnite and sphalerite. Native gold, pyrargyrite and arsenopyrite occur in trace amounts. Historical drilling of the Lulu mudstone assayed 14.4 grams per tonne gold and 1059.5 grams per tonne silver across 14 metres (Owsiacki, 1991). Heritage Explorations Ltd. (now SAS) completed 3 drillholes in the Lulu Zone during 2002 and intersected 11.7 metres grading 19.5 grams per tonne gold and 1,602.9 grams per tonne silver in drillhole 2-113. However, despite these grades and the high gold grades from earlier drilling in 1990-91, Heritage concluded that the mineralization was restricted along strike and at depth.

The Hexagon and Mercury Anomaly Zones were discovered in 2002, during exploration directed by Cambria Geosciences for Heritage Exploration. They are defined by a 4km long, gold-rich multi-element stream sediment anomaly coincident with two areas of strong to intense phyllic alteration cropping out at surface. An interpretation, preferred by McGuigan (2002) is that the veins that host the gold in the Hexagon Zone represent part of a deformed feeder system of an Eskay-type precious metal enriched VMS deposit.

Overall, the stratigraphy encountered on surface in the southwest SIB area is generally equivalent to that found at the Eskay Creek mine, based on litho-geochemistry and the type of stratigraphic units present (McKinley et al., 2009). The major exception is the absence of overlying tholeiitic pillow basalts. However, tholeiitic basaltic intrusive rocks, pillow lavas, and volcanoclastic rocks were encountered below the Coulter Creek Thrust Fault by the deep drill holes EK08-132, 134, and 135 (see below). This suggests that a more complete section of the Salmon River Formation exists in the panel underlying the Coulter Creek Thrust Fault; at surface, the upper part of the Salmon River Formation, including the tholeiitic pillow basalts, has likely been “removed” by fault displacement along the Coulter Creek Thrust Fault and subsequent erosion.

Rhyolitic flows and autoclastic breccias assigned to the Salmon River Formation contain intercalated, lens-like bodies of argillaceous siltstone and finegrained sandstone. Mixtures of brecciated to redeposited sediment and rhyolitic volcanoclastic debris, and mixtures of rhyolite breccias and argillaceous sediment are common in this section of the stratigraphy. The brecciated, mixed, and re-deposited debris suggest eruption of rhyolitic magmas onto and into partially unconsolidated argillaceous sediments.

The rhyolite-rich stratigraphy displays a similar architecture as that found in the sedimentary sequence underlying it. There is a variation in the bedding dip angles from steeply eastward to westward moving from east to west as Coulter Creek is approached. This suggests a thickening wedge of rhyolite and sediment down dip. The felsic volcanic rocks appear to have a transitional affinity at the base of Formation and change rapidly upwards to a tholeiitic affinity (McKinley et al., 2009).

No examples of massive porphyritic rhyolite were encountered in the surface area mapped in 2008. Instead the area containing the Lulu Zone represents a more peripheral facies to the rhyolite domes in the Mackay Adit area and in the 21 Zone area at the Eskay Creek mine (Bartsch, 2001). However, massive porphyritic rhyolite was intersected below the Coulter Creek Thrust Fault in drill holes EK08-132 and 134 indicating that more vent-proximal facies exist at depth. Combined with the significant alteration and mineralization in the footwall of the Coulter Creek Thrust Fault, this footwall appears to be significantly more prospective than the rocks in the hanging wall of the thrust. These vent proximal and mineralized rocks are open to the north and south as well as downdip and provide a much more extensive exploration target than do the equivalent rocks exposed at surface above and to the east of the Coulter Creek Thrust Fault.

Although the lower contact of the Salmon River Formation has not yet been intersected by drilling below the Coulter Creek Thrust Fault, the thickness of the units encountered in the 2008 drilling and the presence of tholeiitic basalt beneath the Coulter Creek Thrust Fault suggest that a complete section of Eskay Rift rocks exists beneath this fault. Precious and base metal-enriched stringer sulphide

mineralization intersected in Hole EK08-134 (see below) clearly highlights that an extension of the Lulu Zone mineralization, and/or entirely new zones of Eskay-style mineralization, may exist in this newly discovered panel of rocks.

Historical Exploration

Consolidated Silver Butte (SIB) and its predecessor companies were active on the SIB from the 1930s until the early 1990s. Later, Heritage Exploration (now St. Andrew Goldfields, SAS) applied a systematic and multidisciplinary approach to its exploration at the core SIB property (including the Lulu and Hexagon zones) in the period 2001 to 2003. Those programs involved aggressive drilling of targets developed by geological mapping, geochemical sampling (stream sediments), lithogeochemical sampling (rocks), and airborne geophysics (AeroTEM II). In aggregate, the historical exploration programs completed a total of 132 diamond drill holes, comprising 19,417 meters of diamond drilling. Most of the historical drilling was targeting the mudstone horizon at the Lulu zone. Historical interpretation of the results concluded that the Coulter Creek Thrust Fault limited the exploration of the Lulu to the south, and at depth.

The Company optioned the Property in May 2008. The Company's 2008 exploration program comprised geological mapping and lithogeochemical sampling, along with 4 drillholes for 2,333.6 metres of drilling. Geological mapping was conducted on the surface over the SIB claims in order to tie the SIB lithologies into the mapping scheme of the more advanced geological work from the 2003-08 Corey programs. That early summer 2008 work confirmed the validity of the drill targets lying in the footwall of the Coulter Creek Thrust Fault, an untested region of the SIB claims. Three of the four diamond drill holes of the 2008 program targeted the footwall to the Coulter Creek Thrust Fault, each intersecting Salmon River formation rocks of the Eskay-type geological signature (see below for details).

The prospective Salmon River formation rocks on the SIB claims are cut off by the Coulter Creek Thrust Fault along strike to the south and also down dip, thereby limiting the prospective area for VMS exploration. All the historical drilling to shallow depths was conducted in a panel of volcanic and sedimentary rocks lying on the hanging wall of the Coulter Creek Thrust Fault. However, fault-displaced targets identified in 2002 for Heritage by McGuigan and Lewis are located in the footwall of the Coulter Creek Thrust Fault and remained undrilled until the Company's 2008 program.

Drillhole EK08-133 was a Lulu Zone confirmation hole and also tested the immediate along strike extension of the zone. The hole is collared in an Eskay-type tholeiitic rhyolite flow breccia that passes into a 15.5 metre interval of highly faulted carbonaceous and finely pyritic mudstone. A 10-metre core interval from 55.7 to 65.7 metres depth returned grades of 9.0 g/t Au, 405 g/t Ag, 0.2% Zn, 0.3% As and 2.9% Sb. This includes a 2.3 metre drilled interval (55.7-58.0 metres) of finely laminated to clastic pale to dark grey massive and semi-massive sulphides (likely stibnite)/sulphosalts and mudstone. This higher grade interval returned 15.9 g/t Au, 1299 g/t Ag, 0.5% Zn, 0.4% As and 7.8% Sb. The notable enrichment in Sb indicates that stibnite likely forms a significant part of the zone.

The 2008 diamond drilling program was targeted at "Lulu horizon" mineralization within stratigraphically equivalent mudstone in the footwall of the Coulter Creek

Thrust Fault, beneath the overthrust Betty Creek formation rocks. Historical diamond drilling on the SIB property was relatively shallow in depth, and did not test the footwall of the Coulter Creek Thrust to any significant depth. Displacement on the Coulter Creek Thrust Fault at the Lulu Zone was estimated at less than 500 metres in earlier work by McGuigan and Lewis, placing targets readily within feasible drilling depths with a light-weight drill rig. Drillholes EK08-132, 134 and 135 were planned to locate the fault displaced segments of the Lulu Zone stratigraphy in the footwall of the Coulter Creek Thrust Fault. All three diamond drill holes penetrated the Coulter Creek Thrust Fault, at depths ranging from 450 to 560m downhole. All drillholes intersected Salmon River formation mudstones, rhyolites and basalts, below the Coulter Creek Thrust Fault. In combination with surface geological mapping by Sebert, the 2008 diamond drilling defined a much refined and very prospective geology.

The most significant mineralization was returned from hole EK08-134. That hole passed through the Coulter Creek Thrust Fault at 458 metres depth. The interval between 488.2 and 513.6 metres is a pale grey-green Eskay-type tholeiitic rhyolite intrusive into mudstone, and in part flow banded. Notably, thin, anastomosing quartz-polymetallic sulphide veins cut this unit. These locally carry up to 15% sphalerite-galena-chalcopryrite-pyrite. Additionally, there are thicker, laminated "stockwork" style quartz-polymetallic sulphide veins with up to 5% sphalerite-galena-chalcopryrite. This extensive veined interval encompasses a **25.4 metre thick drilled interval (488.2 to 513.6 metres) with length-weighted average grades of 2.12 g/t Au, 4 g/t Ag, 0.17% Zn, and 0.13% Pb** plus anomalous As and Sb. The presence of Eskay mine equivalent stratigraphy in hole EK08-134 (tholeiitic basalt and low TiO₂ rhyolite in proximity to mudstone), plus associated hydrothermal alteration and VMS style stockwork mineralization indicate excellent potential for massive sulphide mineralization at this contact both along strike and up and down dip.

Detailed geological mapping complemented by a comprehensive lithogeochemical survey and an airborne geophysical survey has resulted in a very well defined geological template for the Eskay (SIB) Property. This work has clearly demonstrated that the Eskay rift rocks (the Salmon River Formation) trend southward from the Eskay Creek Mine onto the SIB Property.

2010 EXPLORATION ON THE ESKAY PROPERTY

Cambria Geosciences Inc. was retained in 2010 by the Company to supervise and conduct a field exploration program, commencing in mid July and ending at the start of October following a period of planning in late June and early July. The 2010 field program comprised:

- 3856.7 metres of diamond drilling from 5 drillholes (an additional 44.2 metres were drilled in hole "138A" which had to be abandoned and restarted as Hole 138)
- Geological mapping and lithogeochemical sampling

A comprehensive, systematic program of lithogeochemical sampling of outcrop and drillcore was initiated at the Eskay (SIB) property by Cambria in the 2008 exploration season, following its successful application to the Corey Property. This sampling has proven to be an invaluable aid in the differentiation of the major volcanic and intrusive rocks underlying the Eskay (SIB) Property using the immobile element techniques described by Barrett and Maclean, 1999. Christopher Sebert P.Eng. has summarized and interpreted the 2008 lithogeochemical data in an internal report for Cambria (Sebert, 2008a & b). Here, the schema erected by Sebert for the geologically contiguous Corey property has in part been used to identify rock type and magmatic affinity.

The 2010 lithogeochemical data for the Eskay property are presented in full in Appendix B of this report. They are also briefly discussed in the drilling and mapping summaries below. As an adjunct to the geochemical and assay data provided by Acme Analytical Laboratories, Thermo Scientific's Niton portable X-ray Fluorescence (XRF) analyzer was used extensively during the field season to provide an invaluable "first pass" identification of rock type and assessment of metal concentrations.

2010 Diamond Drilling

Appendix A provides summary geology logs of the 2010 diamond drilling and Appendix B presents geochemical and assay data. As in 2008, the 2010 diamond drilling on the Eskay Property has largely concentrated on the extension of the Lulu Zone host rocks on the footwall side of the Coulter Creek Thrust Fault. The Lulu Zone has the same geochemical and geological characteristics as the precious metal-rich Eskay Creek deposit, and equivalent stratigraphy elsewhere on the SIB claims is thus highly prospective for the discovery of "Eskay Creek type" gold-rich massive sulphides. However, the density of historical drilling along strike from the zone and down its dip extent, plus the existence of the Coulter Creek Thrust Fault leaves little room for the discovery of shallow extensions to the zone. Nonetheless, scope does remain for the discovery of "Lulu type" mineralization within stratigraphically equivalent mudstone in the footwall of the Coulter Creek Thrust Fault.

The 2010 drilling commenced on July 26 and finished on October 1. Five drillholes were completed for a total 3856.7 metres of drilling (An additional 44.2 metres were drilled in hole "138A" which had to be abandoned and restarted as Hole 138) (see

Figure 3 and Table 1; sections for this drilling, Figures 4-7, are located in the pocket at the rear of the report).

Table 1. 2010 diamond drilling on the Eskay Property

Hole ID	Easting (NAD83)	Northing (NAD83)	Elev (m)	Dip	Azimuth	Depth (m)
EK10-136	408090	6273393	1072	-74	297	778.8
EK10-137	408094	6273588	1119	-69	300	776.8
EK10-138*	408193	6273100	1010	-64	292	847.3
EK10-139	407354	6273366	907	-55	112	637.0
EK10-140	408181	6272945	1022	-63	292	816.8
Total 5 holes for 3856.7m						

BOLD=DOWNHOLE SURVEY CARRIED OUT

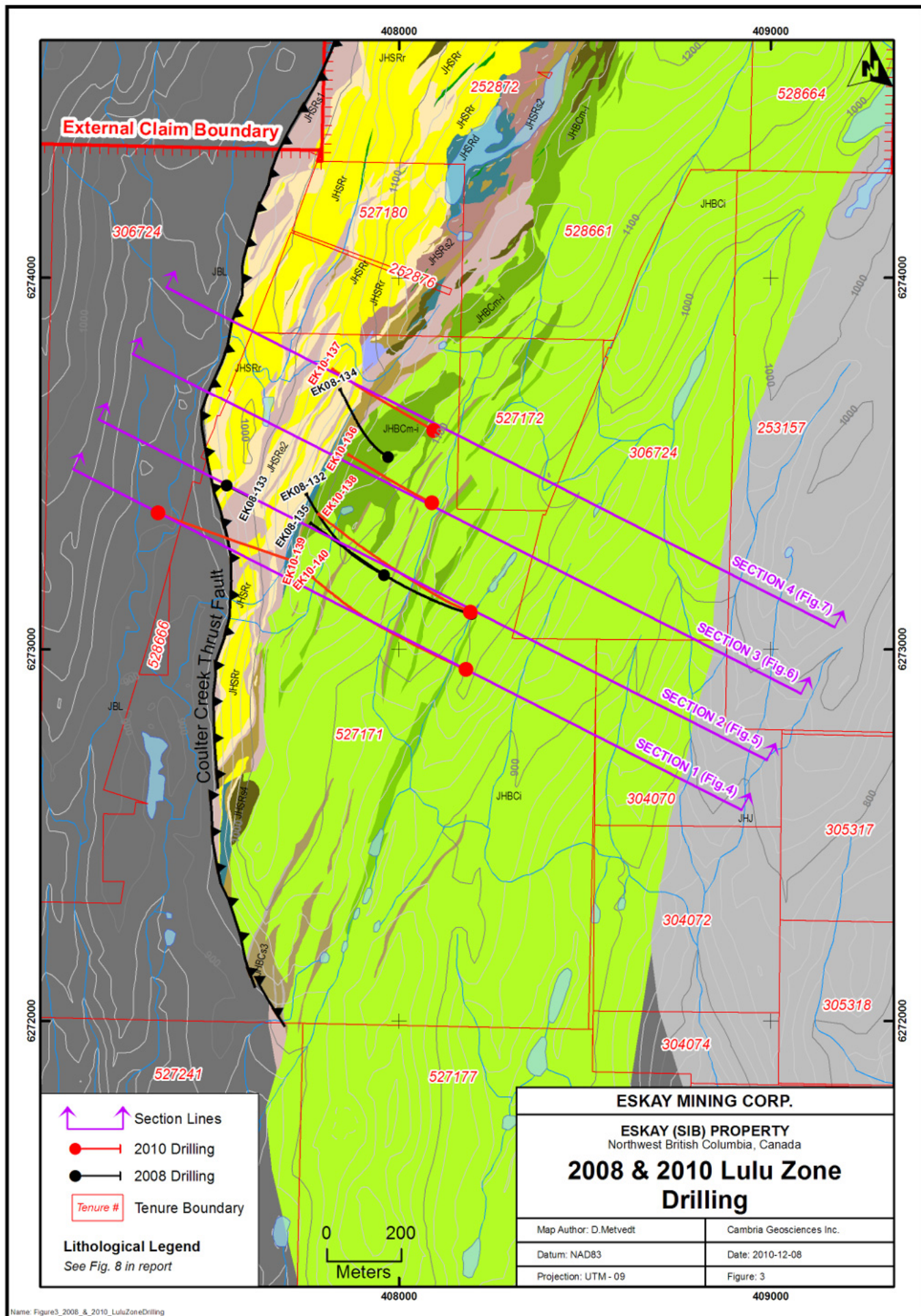
*NOTE: an additional 44.2 metres were drilled in hole "138A" which had to be abandoned and restarted as Hole 138

Drillhole EK10-136

This hole was designed to follow up on the intersection of gold-enriched stockwork sulphide mineralization in Hole 134. It tests the Eskay stratigraphy in the footwall of the Coulter Creek Thrust Fault located about 100 metres south of Hole 134 (see Figure 7). This hole intersected an unusually thick package of Betty Creek Formation rocks in the hangingwall of the CCTF (see Figure 6). True Salmon River Formation rocks, basalts in this case, were not intersected until a depth of 600 metres. One possible explanation for this is that there is a splay in the Coulter Creek Thrust Fault in this location. In other words, there are two branches of the fault that have "thickened" the intersection of Betty Creek rocks. The presence of multiple faults in this part of the drillhole also adds weight to this theory. The lack of occurrence of rhyolite between the fault and the basalt suggests that the hole has pierced rocks higher up in the Lulu Zone section (Note: basalts are not observed at surface in the vicinity of the Lulu Zone likely because they sit higher in the succession than the mineralization and have been faulted and/or eroded away from the surface exposures).

Drillhole EK10-137

This hole was designed to test the stockwork in rhyolite and the adjacent stratigraphy about 100 metres north of Hole EK08-134 (see Figure 3). The Coulter Creek Thrust Fault was intersected slightly deeper in the hole than expected.



Interestingly, a 12 metre interval of rhyolite with weak stockwork veining was intersected in the hangingwall to the CCTF (see Figure 7). This might simply be a wedge of the same Salmon River rhyolites mapped at surface. Assays have been received for this stockwork-mineralized interval. **A 10.0 metre intercept from 463.0 to 473.0 metres depth returned assays of 0.65 g/t Au, 0.26% Pb, 0.38% Zn and 6 g/t Ag. This includes a 0.5 metre intercept from 472.5 to 473.0 metres depth that returned 2.35 g/t Au, 0.97% Pb, 1.89% Zn and 25 g/t Ag.** The mineralization is very similar in style to that intersected in Hole EK08-134, including anomalous As and Sb. This extends the zone of stockwork-mineralized rhyolite an additional 75 metres to the north. The location of the stockwork above the Coulter Creek Thrust Fault in this hole suggests that the mineralization in Hole 137 and in Hole 134 both occur in the fault's hangingwall block.

A typical package of volcanoclastic sandstone and argillite occurs in the footwall panel of the Coulter Creek Thrust Fault, overlain by over 100 metres of Eskay Rhyolite. No basalts were intersected and thus there is no reliable stratigraphic marker in this interval, but it is clearly a good section of Eskay Rift rocks. No significant mineralization was intersected in this section.

Drillhole EK10-138

This hole was designed to test the Lulu Zone host rocks in the footwall panel of the Coulter Creek Thrust Fault down dip of 2008 holes 132 and 135 (see Figure 3). This hole had to be restarted due to problems encountered with the overburden in the top of the hole and shifting of the casing. As such, the original hole was abandoned after 44.2 metres, restarted and drilled to depth. Unfortunately, despite the use of the stabilized core barrel, this hole experienced greater than normal flattening and as such, the hole did not intercept as deeply within the Salmon River Formation as was hoped. However, valuable geological information was still obtained which has been very important in the interpretation of the orientation of the rock units in the footwall panel to the CCTF.

The Coulter Creek Thrust Fault was intersected at a depth of 545.1 metres depth (see Figure 5). A sequence of interlayered Salmon River Formation rhyolite and basalt was intersected below the fault. Below the CCTF, a 2-metre interval contained 5% pyrite-sphalerite-galena stringers within a sheared and veined argillite unit. Within this, a **0.9 metre interval from 556.9 to 557.8 metres depth returned assay grades of 0.06g/t Au, 0.2% Pb, 0.5% Zn and 5 g/t Ag.** Deeper in the hole within the basalt flow breccias, an interval from 799.4 to 803.4 metres depth contained 15% matrix-filling pyrite, but no significant base or precious metal values were detected. Some encouraging intervals of strongly chlorite-altered rhyolite were also intersected in the lowest part of the drillhole.

Drillhole EK10-139

This hole was designed to test for a possible left-lateral displacement of the Lulu Zone-equivalent stratigraphy across the Coulter Creek Thrust Fault. The hole was designed to test approximately 150 metres southwest of the Lulu Zone (see Figure 3). In contrast to the other deep holes in this program, Hole 139 was designed to drill from west to east and to not pass through the CCTF. Rather, the hole was

designed to pass through the younger Bowser Lake Group sedimentary rocks and pass into the underlying Salmon River Formation.

The contact between the Bowser Lake Group and the Salmon River Formation was intersected at a hole depth of 399.4 metres (see Figure 4). This contact is rarely, if ever, seen at surface (it is inferred from outcrop mapping), and it has rarely been drilled on the SIB claims and never this far south. It is very important to know the exact location of this contact as it provides the upper limit to the Salmon River formation and thus the uppermost limit to where Eskay-like mineralization might be expected to occur. Thus, it provides an upper geological “bracket”, the lowermost “bracket” being the contact between the Salmon River Formation and the underlying Betty Creek Formation. This lower bracket is yet to be intersected in the footwall panel to the CCTF and should continue to be tested in future programs. Eskay-type rhyolites which can be correlated with hole 132 to the north were intersected through the rest of the hole down to the final depth of 636.7 metres.

The drilling of this hole was greatly complicated due to high water pressures encountered around 475 metres depth. The water pressure created extreme artesian conditions and a fountain of water at surface and made it very difficult to lower and lock the core tube. In an attempt to counteract this, the NQ rods were left in the ground and the hole was down-sized to BQ core. The drillers were able to advance the hole down to 636.7 metres depth, but the fluctuating water pressures made it increasingly unsafe to operate. In fact, in one instance, the pressures were so great that the water forced the core tube and water swivel out of the rods at a force great enough to cause significant injury if a worker had been in their path. As such, the decision was made to stop the hole before it reached its target depth.

Drillhole EK10-140

Similar to Hole 139, this Hole 140 was designed to test for left lateral offset of the Lulu Zone stratigraphy across the Coulter Creek Thrust Fault. However, this hole was drilled from east to west and was designed to cross the CCTF and drill a deeper portion of the Salmon River Formation than in hole 139 (see Figure 3). The CCTF was intersected at 522.2 metres depth (see Figure 4). Below the CCTF, from depths 530.5 to 598.2 metres, the hole intersected a mixed package of volcanoclastic rock and argillite. This is very similar to the units mapped at surface that mark the transition between the Betty Creek Formation and the Salmon River Formation. This demonstrates that this hole may have drilled very close to that contact, as was the intention for the drillhole. This is very important information because, along with the upper Salmon River contact intersected in hole 139, we now know that we have a succession of Salmon River Formation that is at least 500 metres thick. The hole continued to target depth and intersected a mixture of basalt and Eskay rhyolite. Deteriorating weather conditions precipitated the decision to not extend this hole past the design depth.

Geological Mapping

In 2010 Christopher Sebert continued the geological mapping and lithogeochemical sampling campaign he started in 2008 (Sebert, 2008b), concentrating on the area between the Lulu Zone and the northern end of the SIB claim block. He focused on the facies relationships within the thick Eskay Rhyolite package and the structural

relationships in the upper part of that sequence and the transition between it and the overlying Bowser Lake Group. He has now mapped most of the Salmon River rocks exposed at surface on the SIB claims as well as the adjacent lithologies. He last carried out reconnaissance mapping and sampling on the company's Mitchell claims to the east of Seabridge's Iron Cap Zone. The results of his work are summarized below.

AREA COVERED, OBJECTIVES, AND PRESENTATION

Geologic mapping was performed by the author between July 23 and September 29, 2010 in the northern portion of the SIB Claim Block (Map 1, in pocket). This work augmented previous mapping performed in 2008 in the southwest part of the claim block.

The primary objectives of the 2010 mapping program were the following:

1. Perform detailed outcrop mapping in combination with assessment of existing drill holes in the northern part of the claim block in the area south and southwest of the McKay adit. This area contains extensive zones of silicic and pyritic alteration. Also, Bartsch (2001) reports the presence of a rhyolite flow dome in the area of the McKay adit. The mineralization at the Eskay Creek mine occurs in proximity to a massive flow dome complex composed of coherent rhyolitic lavas flanked by rhyolitic breccias, volcanoclastic rhyolitic, and argillite. Hydrothermal venting can be promoted by the heat flow generated from the emplacement of more massive, voluminous bodies of rhyolite. Mapping and geologic interpretation in this section of the property holds the possibility of providing additional future drill targets on blind sulphide mineralization.
2. Attempt to locate and trace the northern extension of the Coulter Creek thrust fault. Also, try to better define the spatial relations between the Coulter Creek thrust and Argillite Creek faults, and assess the effect of these structures on the location and orientation of economically prospective, Eskay-equivalent stratigraphy.
3. Complete mapping of the upper (Salmon River/Eskay rhyolite) Hazelton stratigraphy from SIB Lake to McKay adit area. Emphasis to be placed on its facies, architecture, alteration, and contained structural features. This serves to extend the 2008 mapping work into the northern part of the claim block including the strong alteration found around Battleship Knoll and SIB adit.

Several secondary objectives include those below:

1. Performing a compilation of the attitudes, character, paragenesis and displacement for prominent cross and axis parallel fault structures on the SIB ridge. Alteration along faults to assess the potential structural role in controlling hydrothermal activity within the footwall stratigraphy was seen as important feature potentially related to the Lulu Zone or other blind sulphide mineralization. Also, better understanding of fault attitude, paragenesis, and associated alteration might aid interpreting the geology in the 2010 drill holes exploring the lower plate of the Coulter Creek thrust fault.

2. Selective outcrop mapping targeted to assess the stratigraphy along the trace of the Coulter Creek thrust fault southward from the SIB ridge. This work would attempt to better define the location of the Coulter Creek thrust fault and the degree of offset on rock units. Additional work in this part of the property could aid in assessing the thickness of Bowser Lake Group sediment overlying the prospective Salmon River stratigraphy.

STRATIGRAPHY AND LITHOGEOCHEMISTRY

General Stratigraphy of the SIB Ridge Area

The area examined consists of a succession of volcanic and sedimentary rocks, which generally young from east to west. Most of the rocks examined on the SIB property during the 2008 and 2010 season were interpreted as belonging to the Bowser Lake, Salmon River, or Betty Creek Formation. The Mount Dilworth Formation label applied by previous workers for Eskay-equivalent rhyolitic rocks is dispensed with and these rocks are interpreted as part of the Salmon River Formation. This re-interpretation is in keeping with the revisions of unit labels embarked on by MDRU in the early 1990's (see Lewis et al., 2001).

The stratigraphic sequence can be broadly summarized (from older to younger) in terms of four units. The first three are composite units interpreted as part of the mid- to early-Jurassic Hazelton Group containing variable combinations of volcanic, volcanoclastic, and sedimentary layers of differing composition. The mid-Jurassic Bowser Lake Group is the fourth unit.

- **Composite Unit 1** is composed of intermediate and minor mafic tuffaceous volcanoclastic units, and of massive to brecciated, pyroxene-amphibole-feldspar-porphyrific, intermediate intrusive rock. These rocks are presently interpreted to belong to the Betty Creek Formation.

The volcanoclastic rocks consist of water-lain lapilli and ash tuff, and of tuffaceous wacke sandstones. Rare worm burrows in tuffaceous wacke sandstone suggest that the clastic rocks were deposited in relatively shallow water below wave base.

Lithogeochemical analyses indicate that the intermediate volcanoclastic rocks vary from andesitic to dacitic composition. Mafic members are generally basaltic-andesite. Samples of the intermediate porphyry possess major and trace-element chemistry that is andesite or dacite equivalent. The volcanoclastic rocks possess trace-element characteristics that suggest they are of transitional to mild-calc-alkaline magmatic affinity and were erupted in an evolved arc setting. The intermediate intrusive rocks were labeled as transitional but otherwise are chemically quite alike the intermediate volcanoclastic rocks.

- **Composite Unit 2** represents a sedimentary-dominated sequence. It is presently seen as the lower stratigraphy of the Salmon River Formation and constitutes a transition zone lying between older volcanoclastic and intrusive rocks interpreted as the Betty Creek Formation and the younger rhyolite-dominated stratigraphy of the Salmon River Formation.

This unit is composed of pebble-cobble conglomerate, dark graphitic argillaceous siltstone and sandstone, and grey to greenish wacke sandstone. Individual layers of these sedimentary rocks tend to be lensey and, with the exception of some siltstone units, are generally poorly bedded. In places, conglomeritic layers display cross-cutting, angular-unconformable basal contacts with finer-grained sedimentary rocks. Both east and west dipping bedding is present. The angular discordant contacts, the presence of conglomeritic rocks, the discontinuity of bedding, and dip angle changes suggest a high energy depositional environment was present, likely along a slope, which may have steepened over time, producing a fan of sediments.

The overall thickness of this sedimentary sequence is highly variable. To the east of the Lulu Zone these sedimentary rocks (as exposed in plan view) measure ~30m in width; to the north of drill pad D (at ~ 6274200N) they measure ~200 m in width. In the Gap Area the sedimentary sequence approaches 300 m in breadth.

Eroded rounded clasts of intermediate volcanoclastic in the conglomerate units are similar in composition to the older rocks lying below and were well lithified at time of erosion. Also, the conglomeritic and sandstone units in the lower portion of the sequence contain rusty, pyritic pebbles and cobbles, which appear to have been derived from the altered felsic-to intermediate rocks found in the upper portions of Composite Unit 1 in the Betty Creek stratigraphy. The presence of these eroded and altered clasts suggests there was a significant hiatus between the emplacement of the Betty Creek rocks and the deposition of the conglomeritic rocks; and that the alteration so prominent in the SIB ridge east of Sib and Gap lakes pre-dates the eruption of low-Ti rhyolite and Eskay Creek mineralization.

Layers of grey-green intermediate ash tuff, lapilli-ash tuff and tuffaceous sandstone are found near the top of the sedimentary sequence just below the contact with rhyolitic rocks. These units vary from andesitic to dacitic composition and have the chemical characteristics of rocks of transitional magmatic affinity. They may be equivalent to tuffaceous rocks outcropping in the HSOV and Spearhead Area on the Corey Property.

- **Composite Unit 3** is equivalent to the Salmon River Formation. It is composed of rhyolitic flow rocks, breccias, and tuffaceous volcanoclastic rocks. Several successive intercalated layers of variably banded rhyolitic flows, autoclastic breccias, and epiclastic rhyolite volcanoclastic rocks appear to be present. Mixtures of rhyolite breccia and carbonaceous sediment, and brecciated and disturbed, re-deposited carbonaceous sediment containing rhyolitic detritus are common. Lesser volumes of rhyolitic tuff and rhyolite-rich tuffaceous debris were also encountered at the base of the rhyolite-rich stratigraphy; these rocks represent early phases of rhyolitic volcanism.

The rhyolitic volcanic rocks possess the low-Ti and transitional to tholeiitic geochemical signature of the rhyolitic rocks associated with the massive sulphide mineralization at the Eskay Creek Mine.

The rhyolite-rich stratigraphy displays a similar architecture as that found in the sedimentary sequence underlying it. There is a variation in the bedding dip angles from steeply eastward to westward moving from east to west down slope towards

Coulter Creek. This suggests a thickening wedge of rhyolite and sediment down dip. The brecciated-rhyolite and sediment mixtures, and disturbed to re-deposited argillaceous sediment suggest the eruption of rhyolitic magmas onto and into partially unconsolidated argillaceous sediments along a slope. The discontinuity of the Lulu sulphide mineralization is likely partially due to the unstable character of the depositional environment.

Examples of basaltic and basaltic-andesite sub volcanic intrusive rocks were encountered in the northern corner of the claim block. Minor mafic dike rocks were also mapped further south in the western portion of the SIB ridge. These rocks possess tholeiitic geochemical signatures and are the intrusive equivalents of pillowed mafic flow rocks found near the Eskay Creek mine.

- **Bowser Lake Group** sedimentary rocks were examined along the western toe of the SIB ridge and in the valley of Coulter Creek. Most outcrops of these rocks lie west and below the plane of the Coulter Creek Thrust. Where examined, they consist of thin bedded, variably graphitic silt and sandstone. Very minor thin layers of fine-grained tuffaceous sediment were noted as well. Load casts and cross bedding in thin-bedded silt and sandstone provide indicators of bedding attitude.

Rock Units and Lithogeochemical Overview

The following subsections provide an itemized description (from older to younger) of the individual rock units that make up the composite units described above.

Lithogeochemical sampling accompanied the mapping in order to help define the different volcanic rock units as to their magmatic affinity and confirm their place within the stratigraphic column. During 2010 lithogeochemical sampling was augmented by the use of a Niton portable XRF, which providing additional spot analyses of rock specimens in the field providing quicker access to Zr-Ti ratios and aiding in daily rock identification.

Composite Unit 1 (Betty Creek Formation?)

These older rocks have been interpreted as being part of the Betty Creek Formation. They are largely of volcanoclastic origin and constitute the water-lain tuffaceous products of successive magmatic events. Individual layers range in composition from basaltic andesite to dacite. Minor thin layers of graphitic siltstone and more heterogeneous sandstone represent hiatuses in the deposition of the tuffaceous volcanoclastic rocks. Lesser, but significant, volumes of massive intermediate to felsic intrusive and porphyritic brecciated rocks occur in the upper portion of the unit.

Green Intermediate Volcanoclastic Rocks (JHBCi)

These rocks are variably tuffaceous, banded, and represent weakly reworked epiclastic crystal ash, lapilli ash, and re-deposited sand-sized lithic and vitric detritus. Individual layers are generally unbedded internally and range in thickness from several metres to 10's of metres. A range of composition and variable degrees of heterogeneity are present in the individual volcanoclastic layers encountered. More tuffaceous beds tend to display a torn veil-type texture with an anastomosed combination of flattened vitric lapilli and ash. Layers of fine- to coarse-grained, banded but un-bedded volcanic sandstone, or tuffaceous wacke, are more granular in

texture, and contain re-worked tuffaceous and lithic detritus, and may be weakly graphitic locally. Heterogeneity in these latter rocks is more obvious, indicated by varicoloured fragments and less often by darker, fine rip ups of argillaceous sediment. Rare examples of worm burrows were observed in a sandstone layer in the upper part of the volcanoclastic rich sequence, which indicates that these units were deposited during hiatuses in volcanic activity and that the water depth was relatively shallow.

Lithogeochemical results indicate that the composition of these intermediate volcanoclastic rocks varies from andesitic to dacitic. The andesitic members tend to possess Zr/TiO₂ between 120 and 150, while the dacitic varieties have Zr/TiO₂ values between 140 and 200 (Figure 2). Most samples analyzed have La/Yb $\geq$ 6 (average ~ 6.5), and Th/Yb $>$ 1.3 ranging to $>$ 2, which is typical of volcanic rocks of calc-alkaline magmatic affinity. The Zr:Y ratios of samples tend to be between 4 and 7; this is like that found in volcanic rocks of transitional affinity. Chondrite-normalized REE patterns display moderate amounts of LREE enrichment and possess very slight concave profiles. Given the tendency to La/Yb $>$ 6, Th/Yb $>$ 1.3 and the slightly concave profiles most samples of these rocks can be labeled as being of mildly calc-alkaline affinity.

Dark Green Mafic Volcanoclastic Rocks (JHBCm)

Mafic volcanoclastic rocks mainly composed of lapilli tuff were mapped on surface on the eastern side of the SIB ridge in 2008. In 2010, no further delineation of these rocks in the older stratigraphy was attempted as they are likely not an important factor for massive sulphide exploration.

These mafic volcanoclastic rocks are intercalated within the intermediate volcanoclastic rocks described above. Occurrences of mafic lava breccia were encountered in core from the 2008 diamond drill holes. The brecciated rocks may be spatially associated with the mafic tuffaceous volcanoclastic rocks that outcrop on surface.

Lithogeochemical analyses of these older mafic rocks indicate that they are nearly all of basaltic andesite composition. Zr/TiO₂ is between 72 and 134, and SiO₂ contents vary between 51 and 58%. These mafic rocks possess La/Yb ranging from 5.3 to 11.5, and Th/Yb $>$ 0.65. Zr/Y is between 2.9 and 13.7. These results are not definitive of either transitional or calc-alkaline affinity but rather overlap both classifications. Individual samples have been labeled as either calc-alkaline or transitional weighing the values of La/Yb and Zr/Y. However, all these samples probably represent the product of one magmatic episode and chondrite-normalized REE patterns display moderate LREE enrichment and are weakly concave in shape. The latter feature hints at a calc-alkaline affinity and at present the authors favour this interpretation.

Porphyritic Massive and Brecciated Intrusive Rock (JHBCi)

Massive, pyroxene-amphibole-feldspar-porphyritic intermediate to felsic intrusive rock is present in the upper portion of Composite Unit 1 near its western boundary with the transition zone sedimentary rocks. This rock type is common in the area

around drill pad D, and stretches northeastward past the east side of Sib Lake and onward to the northern boundary of the property.

Sections of Battleship knoll are also made-up of this intrusive phase. Large volumes of relatively monomictic, brecciated porphyritic intrusive rock were encountered in the area from the collar of DDH89-12 & 13, northward along the contact between overlying, younger, sedimentary rocks.

Lithogeochemical analyses of the balance of the intrusive rock samples indicate that this unit is of intermediate to felsic composition and contains SiO₂ concentrations between 55 and 68% and Zr/TiO₂ varies from 126 to 239. Most examples are of andesite-equivalent composition with a significant minority of dacitic chemistry. La:Yb ratios vary between 4.5 and 10.6. Zr:Y ratios range from 3.9 to 10. These ratios imply a transitional to calc-alkaline affinity. Th:Yb ratios are generally > 1.1 – commensurate with calc-alkaline intrusive rocks. Most of the chondrite-normalized REE patterns from these rocks possess slightly more kinked profiles compared with the host volcanoclastic rocks. Also they are slightly less enriched in LREE than the intermediate volcanoclastic rocks. Individual units can be labeled as transitional or mildly calc-alkaline.

Fine and Coarse-Grained Intercalated Sedimentary Rocks – Transition Zone Composite Unit 2

Cobble and Boulder Conglomerate (JHSRs4)

Lenses of cobble and boulder-rich conglomerate are present within this sedimentary package, notably in its lower portions close to the contact with older volcanic and intrusive rocks of Composite Unit 1. The conglomeritic rocks are dominated by rounded to sub rounded clasts of green intermediate volcanoclastic and porphyritic volcanic or intrusive rock. Local examples lying south of Battleship Knoll just west of the SIB adit contain a significant component of angular intrusive clasts and have the appearance of talus breccias. The matrix of the conglomeratic rocks is composed of fine sandstone, which may have a significant dark, muddy component in places but generally is largely of intermediate volcanic derivation. Lithogeochemical analysis of green intermediate volcanoclastic cobbles yields major and trace-element chemistry similar to the mildly calc-alkaline rocks of Composite Unit 1. Quartz-sericite-pyrite-altered cobbles occur in within conglomerates in the vicinity of the SIB Adit and Battleship Knoll. The alteration in the clasts is similar to that found in the prominent rusty-weathered outcrops found in this area and suggests that the alteration there is older than the Salmon River Formation rocks making up Composite Unit 2 and 3.

Fine to Coarse-Grained Wacke Sandstone (JHSRs3)

Layers of green to grey or brown, fine-to coarse-grained sandstone are located in the lower section of the sediment-rich stratigraphy. These rocks tend to be largely composed of grains eroded from intermediate intrusive and volcanoclastic rocks similar to those discussed above. Dark graphitic, argillaceous grains are present in places and these represent rip-ups of finer-grained muddy sediment. Layers of wacke sandstone containing rounded cobbles and boulders of intermediate volcanoclastic and intrusive rock are present locally in proximity to more cobble-dominated conglomerate layers. Some boulders observed in the field reach, or exceed, several metres in diameter.

Dark Argillaceous Silt and Sandstone (JHSRs1 & JHSRs2)

Layers of graphitic, fine-grained sediment, ranging from siltstone to fine-grained sandstone are intercalated in the sedimentary sequence at various levels. These rocks are thinly bedded in places and represent periods of relatively low energy deposition. Beds of these rocks may be discontinuous in places reflecting their scouring-out by more coarse-grained sedimentary material.

Intermediate Tuff Layers (JHSRd)

Layers of grey-green intermediate ash tuff, lapilli-ash tuff and tuffaceous sandstone are found near the top of the sedimentary sequence just below the contact with the rhyolitic rocks of Composite Unit 3. These units are weakly heterolithic containing other sedimentary rip-up clasts, and display various degrees of reworking from minimal to epiclastic bedded examples. Local layers of green volcanic pebble to cobble conglomerate with a sandy tuffaceous matrix are associated with this unit, and form unconformable contacts with older, finer-grained silt and sandstones.

Zr/TiO₂ in lithogeochemical samples ranges between 143 and 363. Most of these rocks are of dacitic composition with minor andesitic constituents. La/Yb ranges from 3.5 to 6.4 and Zr/Y ranges between 3.2 and 5.7, which are values commensurate with that found in other transitional volcanic rocks. The chondrite-normalized REE patterns of these samples display a kinked shape – flat through the HREE to MREE, and inclined upward through the LREE. This shape is typical of transitional volcanic rocks. Th/Yb is generally >0.65, which is a calc-alkaline trait, however the Th:Yb ratios in these rocks (nearly all <1.3) is noticeably lower on average than those found in the older intermediate volcanic rocks making up Composite Unit 1.

Rhyolite-rich Upper Stratigraphy – Composite Unit 3

Rhyolite Flow Rocks and Rhyolite Breccia (JHSRr)

Variably banded to brecciated rhyolitic flow rock are the dominant lithologies encountered in the upper portion of the Hazelton stratigraphy mapped in the western portions of the SIB ridge.

The rhyolitic lava rocks tend to be sparsely quartz-feldspar-porphyritic or more often aphanitic. Banding in lava flows varies from well developed to hardly discernible (Photo Plate 1). Rhyolite breccias are mostly of the autoclastic variety - products of a combination of processes including autobrecciation and quench brecciation (Photo Plate 2). A portion are autobreccias, which are relatively matrix-poor with larger, variably banded angular to lensey clasts. Others are more typical of hyaloclastites with sharp angular to cusped clasts partially supported by a finer-grained granular matrix. Most of the rhyolite volcanic rocks mapped have seen at least weak levels of chlorite+/-sericite+/-quartz alteration, which in most locales is likely the result of sea-floor metamorphism. Alteration intensity increases in tenor in the immediate area of the Lulu zone and on strike to the northeast, where pyritization is also present in places. The rhyolitic rocks sampled tend to be elevated in SiO₂ (~79%) and possess high Zr:TiO₂ ratios exceeding 2000. This reflects the low-Ti geochemistry of these rocks (generally <0.1% TiO₂). La/Yb is <6.5 for all the



Photo Plate 1. Weakly flow banded rhyolite lava.



Photo Plate 2. Rhyolite autoclastic breccia. Prominent flow banding is visible in partially interlocking, angular to cusped clasts.

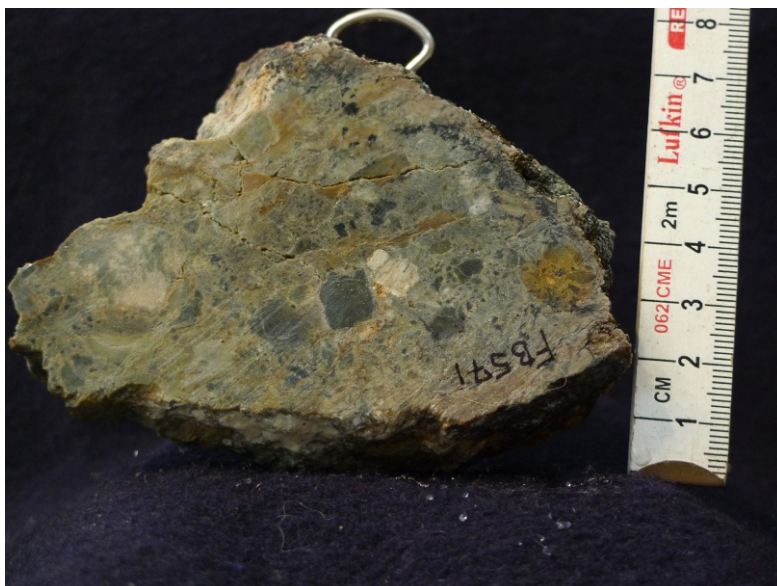


Photo Plate 3. Epiclastic rhyolite breccia or volcanic conglomerate. Matrix-supported angular to subangular clasts display varying coloration and alteration styles. Pale bleached, darker chlorite-altered, and rusty oxidized fragments are prominent.

rhyolite samples taken and Zr/Y is less than 4, which indicates that nearly all of these rocks can be classed as being of tholeiitic affinity.

It has been proposed by Barrett and Sherlock (1996) that the rhyolitic rocks at the Eskay Creek mine are the products of partial melting of the lower crust. This hypothesis is supported by the low-Ti chemistry, a result of the refractory nature of most Ti bearing phases.

Rhyolitic Tuff-Rich Breccia, Felsic Sandstone, and Rhyolite-Rich Conglomerate (JHSRe2)

Minor volumes of rhyolite tuff-rich breccia, rhyolitic tuffaceous sandstone or epiclastic ash, and rhyolite-rich volcanic conglomerate were encountered near the basal contact of the rhyolite-rich stratigraphy representing Composite Unit 3. The tuff-rich breccia possesses a light grey matrix is rich in rhyolitic ash and supports a variety of sub rounded to angular lithic clasts including pale, banded rhyolitic tuffaceous and lava clasts, and a variety of grey to greenish intermediate clasts. This unit represents heterogeneous tuffaceous debris potentially emplaced as mass flows in the early stages of the rhyolitic volcanic cycle that succeeded Composite Unit 2.

Felsic tuffaceous sandstone is fine-grained and light in colour, made-up of reworked rhyolitic crystal ash. It is banded, but generally not bedded. Coarser, weakly heterolithic felsic sandstone is composed of granular rhyolitic spall.

Rhyolite-rich felsic volcanic conglomerate is dominated by granular to cobble-size rhyolitic detritus. It displays heterogeneity in the form of varicoloured, differently altered rhyolite clasts (Photo Plate 3) and by the inclusion of minor volumes of other volcanic and sedimentary detritus. Select outcrops were observed that contained crude bedding composed of alternating metre-scale layers of relatively finer pebbly versus more cobble-rich ones.

Dark Argillaceous Mudstone, Siltstone, and Sandstone (JHSRs1 & JHSRs2)

Layers of graphitic mudstone, argillaceous siltstone, and argillaceous fine-grained sandstone are intercalated with the rhyolitic flow rocks.

The argillaceous sediments tend to be composed of graphitic siltstone and very fine-grained sandstone and represent the products of relatively more deep-water deposition. These rocks are bedded in places but are more often found to possess disrupted textures consisting of re-deposited lumps or the sediment is totally remixed with varying amounts of included rhyolite fragments. These disrupted units allude to the eruption of rhyolite into and over what was soft sediment.

Mixtures of Rhyolite Breccia and Sediment (JHSRe2)

The western portion of the SIB ridge is composed of multiple layers of rhyolite flows and autoclastic breccia that include intercalated mixtures of epiclastic rhyolite breccia and dark, graphitic mudstone, siltstone and sandstone. These latter mixtures display a wide variation in the proportion of included sediment versus rhyolitic spall from place to place (Photo Plate 4 to 6).

These mixed units are the products of episodic rhyolite eruption and emplacement over carbonaceous argillic sediment deposited in more quiescent periods.

Basaltic Intrusive Rocks (JHSRm)

Small (10 to 100 m scale) basaltic plugs and dike rocks intrude the rhyolite-rich stratigraphy of Composite Unit 3 and sedimentary rocks of Composite Unit 2. The mafic intrusive rocks are generally fine-grained and variably pyroxene-feldspar-porphyrific. Alteration is generally weak consisting of chloritization related to seafloor and later metamorphism. Some examples of mafic intrusive rocks, specifically those intruding the northern portion of the property, exhibit moderate levels bleaching by sericitization, mild silicification, and pyritization. These particular rocks had been mapped as intermediate intrusive by previous workers but were found to possess a mafic, tholeiitic geochemical signature.

These rocks have La:Yb ratios <4 and Zr:Y ratios <4 – and a tholeiitic geochemical signature. They represent sub-volcanic intrusive equivalents to the basaltic flow rocks encountered near the Eskay Creek mine. Most examples sampled on the surface of the SIB ridge possess Zr:TiO₂ ratios that are indicative of a basaltic andesite composition and display a high-Ti geochemical signature. Most samples contain >1.5% TiO₂. This is different from the average basalt chemistry found on the Corey property to the south, which tends to have lower more NMORB-like TiO₂ contents and are more primitive in chemistry (i.e. lower SiO₂ and higher MgO).

Bowser Lake Group Sedimentary Rocks (JBL)

Bowser Lake Group rocks outcrop on the toe of the western slope of the SIB ridge and in the valley of Coulter Creek. These rocks were not extensively mapped in 2008 or 2010, but where encountered generally consist of thin bedded, variably graphitic silt and sandstone. Beds of heterolithic pebble conglomerate and medium to coarse-grained wacke sandstone were encountered near the toe of the SIB ridge in the northern section of the map area. Very minor thin layers of fine-grained tuffaceous sediment were noted as well. Load casts and cross bedding in thin-bedded silt and sandstone provide indicators of bedding attitude.

Facies and Stratigraphic Architecture of the SIB Ridge

The following provides observations and interpretations regarding facies and contact relationships for the major geological units covered by the 2008 and 2010 mapping program.

Depositional Environment of Older Footwall Rocks

These calc-alkaline intermediate (and lesser volumes of transitional mafic) tuffaceous volcanoclastic rocks represent water-lain lapilli and ash tuffs. These rocks were laid down below wave base but still in a relatively shallow water environment given the presence of rare worm burrows. The volcanic sandstones layers represent hiatuses between volcanic eruptions, where eroded pyroclastic material was deposited rapidly given the general absence of internal bedding in these units. Deposition may have taken place on a gradually sloping shelf just offshore from several volcanic centres.

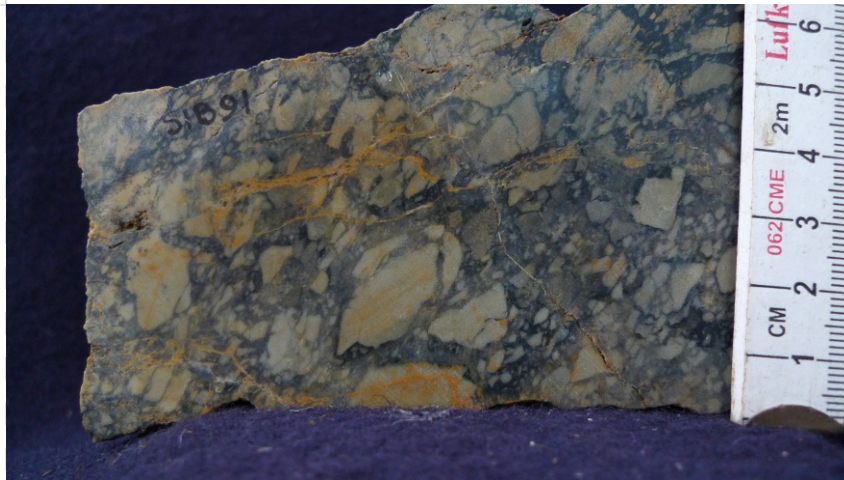


Photo Plate 4. Mixed rhyolite autoclastite and dark, carbonaceous siltstone.



Photo Plate 5. Mixed rhyolite autoclastic debris and dark argillaceous silt to fine-grained sandstone.

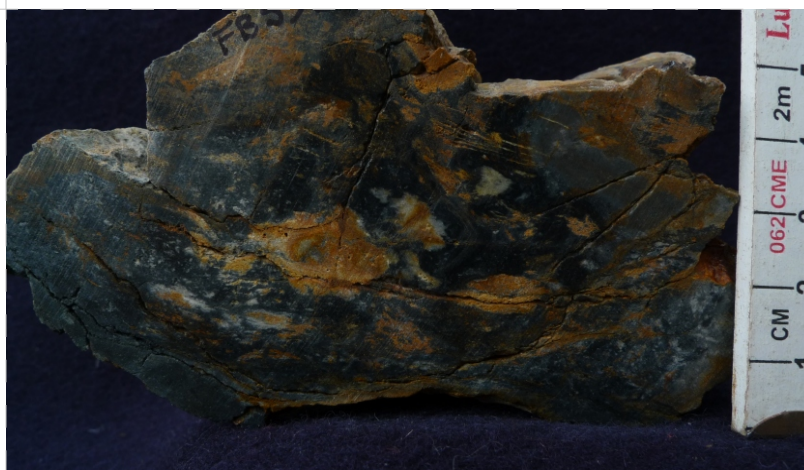


Photo Plate 6. Mixed dark carbonaceous mud to siltstone with a lesser volume of rhyolite breccia spall.

Intermediate porphyritic intrusive rocks intruded the intermediate volcanoclastic rocks after their deposition. These rocks tend to be massive deeper in the footwall but display brecciation in areas located close to the contact with overlying sedimentary rocks. Outcrops of brecciated intrusive include those that are basically monomictic, composed of interlocking angular to sub angular pebbles and cobbles, which display the signs of incipient fracturing and granulation in places. The absence of significant heterogeneity or fine debris in these units implies that they are primary intrusive breccias rather than sedimentary units. Mixtures of sediment and intrusive breccias are also present near and at the contact with overlying fine-grained argillaceous silt and sandstone. These units tend to lack bedding and consist of granular, porphyritic intrusive with varying amounts of grey fine-grained argillaceous sediment. Angular sediment fragments are present and these imply that the sediment was at least partially consolidated before its disruption. These occurrences of mixed granular breccia may represent intrusion into soft sediment, or are later local erosional products of mass wasting.

Nature of the Contact between older Betty Creek and Lower Salmon River Formation Rocks

The contact between older green intermediate volcanoclastic rocks of Composite Unit 1 (interpreted as Betty Creek Formation) and younger sedimentary and rhyolitic rocks lying to the west can be interpreted to be largely of a sedimentary, stratigraphic nature. Deposition of argillaceous and heterogeneous sedimentary rocks of Composite Unit 2 succeeded that of older intermediate volcanoclastic ones and the material within these relatively younger sedimentary rocks was partially derived from the older volcanic and intrusive rocks of Composite Unit 1.

The attitude of the sedimentary strata within lying above the contact with Betty Creek volcanoclastic rocks generally ranges from 010° to 060° azimuth, with a tendency of azimuths to shift more easterly as one moves northward on the SIB ridge. Dip angles range from steeply eastward (-60°) to steeply westward (-75°). Individual sedimentary layers tend to be lensey. Angular unconformable contacts occur within the sedimentary rocks deposited above intermediate volcanoclastic rocks of the Betty Creek Formation. Certain conglomeritic units, containing rounded green volcanoclastic-derived cobbles and pebbles, have been observed as cutting finer-grained sedimentary layers. One such occurrence of conglomerate lies to the west and on top of a siltstone layer, which it truncates; the siltstone unit dips steeply east at a similar attitude to the volcanoclastic rocks of the Betty Creek Formation. In contrast, certain conglomeritic rocks have been found to display shallow attitudes although as a rule they tend to dip very steeply eastward to steeply westward. This variation from steep eastward to westward dipping strata is also present within the rhyolitic rocks and sediments of Composite Unit 3 (Salmon River Formation) examined north of the Lulu Zone.

Several oriented samples of laminated argillaceous silt and sandstone collected contain load casts that suggest bedding tops are west-facing. The angular unconformable contacts between conglomeritic layers and siltstone layers described above also suggest a younging trend to the west. Therefore, barring complications from folding, east-dipping beds in the stratigraphy tend to be overturned and west-dipping ones are upright.

Given the presence of angular unconformable contacts, the change in bedding attitudes, lensey bedding, and the presence of large volumes of coarser-grained

sediments, it appears that the younger sedimentary and volcanic strata of the Salmon River Formation were deposited on top of the older rocks of Composite Unit 1 (Betty Creek?) on a paleo-slope, which may possibly have become steeper over time as a result of tectonic activity. Also, there may have been a significant hiatus between the cessation of deposition of older, Betty Creek, volcanoclastic rocks and deposition of Salmon River sedimentary rocks. The lithic nature of green volcanoclastic cobbles in the conglomerate layers mentioned above suggest that the older intermediate volcanoclastic rocks were well lithified by the time of their erosion and re-deposition.

A semi-continuous band of andesitic to dacitic lapilli to ash tuff and accompanying tuffaceous sandstone and conglomerate is located at the top of Composite Unit 2 at the base of the low-Ti, Eskay-type, rhyolitic volcanic rocks. It displays a range of contact attitudes with older silt and sandstone lying beneath ranging from steeply east-dipping to steeply west-dipping. Minor examples of shallow (-20°) contact attitudes are also present. It is semi-conformably to unconformably overlain by the rhyolite flows and breccias on steep east to west-dipping and locally on shallow, sub-horizontal contacts.

Nature of Salmon River Stratigraphy – Composite Unit 3

The rhyolitic rocks occur as flows, flow breccias, and autoclastic breccias. Sedimentary rocks included within the Salmon River stratigraphy consist of carbonaceous siltstone to mudstone, felsic volcanic sandstone to conglomerate, intermediate (dacitic) tuffs and tuffaceous sandstone.

Brecciated rhyolitic rocks are more prevalent than massive flow rocks in the southwestern area examined in 2008 and also in most of the middle and northern portion mapped in 2010. Carbonaceous mudstone and siltstone, felsic sandstone and volcanic conglomerate occur at the periphery and within the rhyolitic flow and brecciated rocks. Both sharp and gradational bedding contacts between rhyolitic and sediment-rich layers are present. Individual rhyolite flow and breccia units are discontinuous and are intercalated with mixtures of rhyolitic granular to pebble breccia and argillaceous sediment. These mixed units are most common adjacent to several sediment-rich faulted bands forming north to northeast-trending (005° to 045°) gulleys and along the western slope into the valley of Coulter Creek. These mixtures of brecciated rhyolite and disrupted fine-grained argillaceous sediment represent lava rubble and ripped-up sedimentary detritus forming carapaces and basal blankets to the flow rocks.

An example in point, the sulphide of the Lulu Zone is contained in Composite Unit 3, in a significant, mappable band of argillaceous silt and sandstone. This extends from the 2008 Lulu drill pad area (at ~6273400N) north northeastward. The sediment-rich band is recessive and forms a prominent lineament, which can be traced northward (to ~6274200N). Coherent bedding in the sedimentary rock is uncommon. Instead, bedding is discontinuous, or is wholly disrupted and brecciation and remixing of what was unlithified sediment has occurred.

The general, strike directions of sedimentary and rhyolitic volcanic units in Composite Unit 3 is northeasterly, with azimuths between 015° and 055°. In the southwestern area of the SIB ridge layering tends to become more northerly-striking as one moves westward.

Within the rhyolite-dominated stratigraphy, dip angles vary widely from -60° E to -35° W but are generally sub-vertical in the area of the Lulu Zone. In the northern section of the claim block the rhyolitic rocks tend to display steep to shallow northwestward dipping layering and banding between -80° to -25° (Photo Plate 7 and 8). Generally, there is a tendency towards more shallow west-dipping rhyolite flows and brecciated rhyolite near the western toe of the SIB ridge on the flank of Coulter Creek as opposed to steeply east-dipping strata close to the contact of rhyolitic rocks with older sedimentary and tuffaceous rocks to the east. Rhyolitic flows, brecciated rhyolite and rhyolite sediment mixtures tend to give way to bedded carbonaceous mudstone and siltstone at the western toe of the SIB ridge. Bedding in these rocks exhibit a wide range of dip angles from shallow to sub-vertical east-dipping and overturned.

Facies Analysis

The disturbed to re-deposited nature of the sedimentary rocks and mixtures of rhyolite and argillaceous sediments suggest the rhyolitic volcanic rocks were emplaced in an unstable setting.

Observation of the dip angles of rhyolite flows, and rhyolite-rich clastic rocks in 2010 reveals a combination of shallow and steep dips adjacent to north to northeast trending gulleys in the western part of the SIB ridge (Photo Plate 9 and 10). The north to northeast trending gulleys and sediment rich bands cutting the rhyolite-rich sequence do not simply represent intercalated beds of argillaceous sediments in the rhyolitic volcanic rocks but rather represent fault structures along which extension occurred in response to the formation of the Eskay rift. Rhyolitic volcanism and sedimentation was controlled by these structures and the shallow versus steep dips in the rhyolite-rich stratigraphy can be seen as emplacement of flows and clastic rocks along and over a series of stepped faults. Subsequent regional tectonic events have rotated the original bedding attitudes, and fault contacts tending to steepen and (in places) overturn them.

In terms of massive sulphide exploration potential, Bartsch (2001b) mapped and interpreted the volcanic facies of the rhyolite-rich stratigraphy in the vicinity of the Mackay Adit and in the area of the Eskay Creek mine. He identified the presence of flow domes and feeder dikes in these locations and then described the more peripheral volcanic facies consisting of flow brecciated rhyolite and black matrix breccias, where rhyolite breccias and argillaceous sediment were intermixed. No sizeable outcrops of massive rhyolite flow domes were encountered on surface in the southwest SIB Area during 2008. Instead the rhyolite-rich stratigraphy in the area of the Lulu Zone appears to be equivalent to Bartsch's "Rhyolite outer flow dome facies", which is typified by intercalated rhyolite flows and autobreccia, and occurrences of black (sediment) matrix rhyolite-rich breccias.

In 2010, more cohesive rhyolite stratigraphy composed of banded to locally massive rhyolite flows and intercalated breccias were observed in the northern corner of the SIB claim block north of 6275500N (UTM) and Gap Lake. Unlike the rhyolite-rich stratigraphy further to the southwest, this occurrence of rhyolite did not contain substantial occurrences of intercalated sedimentary layers. Significant volumes of mixed rhyolite spall and argillaceous silt were only encountered on the northwest-facing slope of this rhyolite mass toward the headwaters of Coulter Creek. In addition this area of the SIB claims contains larger, more voluminous examples of



Photo Plate 7. Salmon River Formation rhyolitic stratigraphy consists of alternating rhyolite flows and intercalated debris of rhyolite breccia spall and reworked argillaceous silt to mudstone. Photo is of an outcrop in the northern section of the claim block near the contact between the Hazelton and Bowser Lake Group. Bedding of the rhyolite and mixed debris here is moderate to shallow northwest dipping.



Photo Plate 8. Another example of Salmon River Formation rhyolitic stratigraphy with north-west dipping bedding. Layers consist of rhyolite breccia with varying amounts of included, reworked argillaceous sediment. Location is in the northern portion of claim block; Coulter Creek lies below and to left.



Photo Plate 9. Brecciated rhyolite flows display sub-vertical dipping banding and contacts on the south-eastern wall of a northeast (~ 045) trending fault controlled gulley. This structure is joined by a north-northeast-trending (020) fault in the centre of the photo.

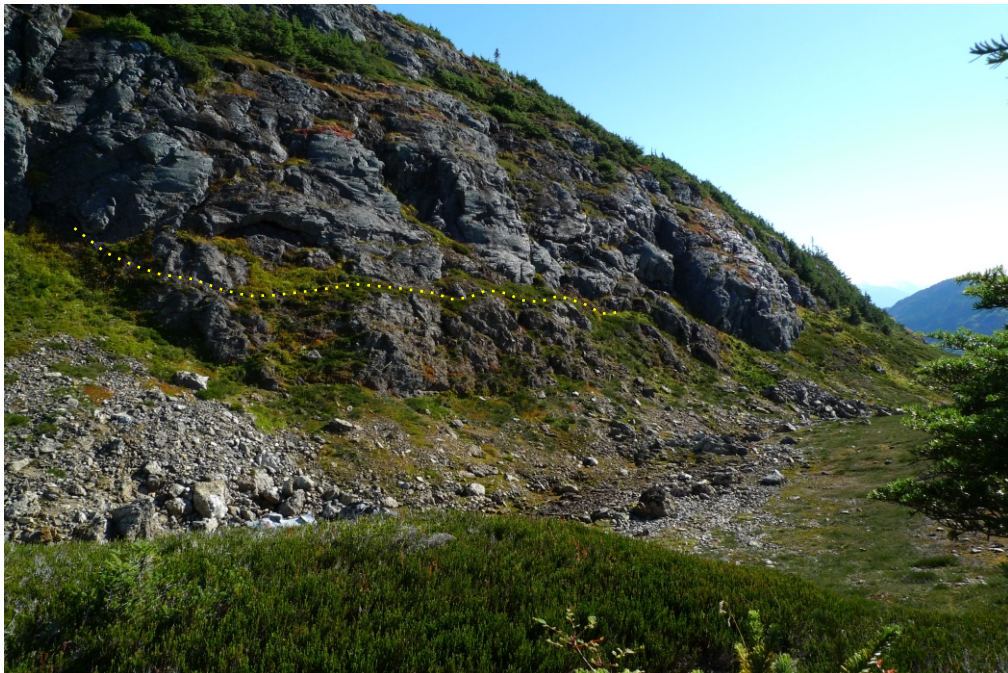


Photo Plate 10. Rhyolite flows overlie a layer of mixed rhyolitic spall and argillaceous sediment on the southeastern margin of a north-northeast-trending (035) faulted gulley. The basal contact (dotted yellow line) of the rhyolite flow rock dips shallowly to the west.

mafic intrusive rocks, and their presence suggests that this section of the property was more proximal to a magmatic centre.

Note on the Alteration Around the SIB Adit and Battleship Knoll

A band of sericite-pyrite +/- silica +/- K-spar alteration occurs along the crest of the SIB ridge. This alteration extends from the southwestern part of the property northeastward on the east side of SIB and Gap lakes and across the northern boundary of the claim group. It is almost wholly contained within older intermediate to felsic intrusive and volcanoclastic rocks of Composite Unit 1 presently interpreted as Betty Creek Formation their contact with overlying sedimentary strata of Composite Unit 2 interpreted as being the lower part of the Salmon River Formation. Alteration appears to be partially controlled by certain northeast and north-northeast-striking faults and is generally most intense in the vicinity of massive to brecciated felsic to intermediate intrusive rocks. Good examples of intense alteration in intrusive rocks occur in and around the Sib Adit and Battleship Knoll.

Altered cobbles of intermediate to felsic volcanoclastic and intrusive rock were observed within heterogeneous conglomeritic rocks in the lower part of Composite Unit 2. This suggests that erosion of pre-existing alteration has occurred and that the alteration pre-dates emplacement of Salmon River Formation Rocks.

Mapping of the Battleship Knoll Area reveals the knoll itself to be largely composed of strongly altered felsic (dacitic-equivalent) intrusive rock. It possesses similar geochemical characteristics as intermediate to felsic rocks lying to the east across the valley hosting Sib and Gap Lakes. It is interpreted as a topographic remnant of the older footwall rocks, possibly a large foundered block, which is partially surrounded by younger unaltered tuffaceous rocks and sediments belonging to the lower Salmon River stratigraphy.

Structural Geology

The structural data collected in 2008 and 2010 indicates that the SIB Ridge was subjected to multiple deformational events, which resulted in the shearing, faulting, and rotation of volcanic and sedimentary strata.

General Summary of Structural Fabric

Bedding in Hazelton Group rocks has largely been rotated to steeply east and west-dipping attitudes with east-dipping bedding overturned.

The large volumes of volcanoclastic rocks comprising the older calc-alkaline stratigraphy of Composite Unit 1 (Betty Creek Formation?) are generally strongly foliated sub parallel to bedding. This strong bedding sub parallel foliation is labeled S1 and has an average orientation of ~040°/67°E. This foliation is also present in the overlying sedimentary rocks and in the rhyolitic rocks of the Salmon River Formation. The average strike of S1 is similar to that in the older stratigraphy, but its dip tends to be slightly steeper. It is generally far less penetrative in character in the more competent rhyolitic rocks of Composite unit 3 (Salmon River Formation).

An elongation lineation (L1) related to stretching is well developed along S1 within the older volcanoclastic rocks of Composite Unit 1 and to an extent within sedimentary units found in Composite Unit 2. This feature is probably related to an episode of shearing parallel to the S1 foliation plane. The elongation lineation generally rakes steeply to the south in the plane of the S1 foliation with an average orientation of 64° - 155°. This implies that (in present terms of reference) a north-northwest directed principal stress vector was present. Examples of steeply north raking and shallower southward raking elongation lineation are present, but not common. The S1 foliation is also slightly undulating and shallowly folded, with estimated limb orientations of ~057°/71° and ~016°/67°. This undulation is the result of the superposition of later deformational events.

Examples of foliation boudinage occur in the strongly foliated eastern volcanic stratigraphy of Composite Unit 1. The boudinage is oriented in conformance with a northerly to north-northwesterly shear direction. Some examples could have been produced by a principal stress oriented roughly parallel to that suggested by the L1 elongation lineation (~335°); other occurrences by a more northerly directed shear event.

Numerous S1 sub-parallel faults are present in Hazelton Group rocks especially in the older units in Composite Unit 1. These structures tend to steeper dips than the S1 foliation, and represent a more brittle phase of deformation. They are discussed in more detail in the section below.

A later, generally steeply east and west-dipping foliation or spaced cleavage is present in the Salmon River and Betty Creek rocks. It is weaker and less penetrative and tends to be localized in less competent sedimentary and tuffaceous volcanoclastic units. This foliation - S2 - tends to strike more northerly than the earlier (S1) foliation at a given location, and is generally more steeply dipping. The average orientation of S2 within Hazelton Group rocks is 027°/80°E. Somewhat in similar fashion to S1 it displays a range of strike orientations (between 008° and 045°). This range is the product of a rotating stress field and influence of the earlier S1 foliation, which tends to strike more northeast.

The sedimentary rocks of the Bowser Lake Group are noticeably less foliated than those of the Hazelton Group but are folded and possess a spaced axial planar cleavage. The cleavage is related to the Cretaceous-age Skeena Fold and Thrust Event. It is equivalent to the S2 foliation found in the Hazelton Group rocks mentioned above. In the area just to the west of the SIB property, Bowser Lake Group rocks display open to tight folding at a wavelength on a scale of 20 to 100 m.

Summary of Fault Structures

Several families of fault structures are present in the mapped area. Certain fault structures were identified that controlled the eruption of the low-Ti rhyolitic rocks of the Salmon River Formation closely associated with mineralization at the Eskay Creek mine.

The fault sets are summarized in order of older to younger (as interpreted to date). Kinematic indicators present are discussed and used to interpret movement direction and history.

Faults, striking northwest (290° to 325°) and generally dipping moderately steeply northeastward, were observed cutting older intermediate volcanic rocks of Composite Unit 1 northeast of Gap Lake. Others were observed cutting the contact between composite unit 3 (Salmon River) rhyolitic volcanic rocks and younger carbonaceous mud and siltstones, or contact argillite, in the northern part of the claim block. Another example cuts Salmon River and older footwall stratigraphy near the Lulu Zone.

Steeply raking striae and slickensides indicate a phase of dip slip or normal movement. Strike-slip movement also occurred on some of these structures and is evidenced by shallow raking striae and ancillary shears.

A portion of the northwest-striking structures exercised control on rhyolitic volcanism. This is evidenced by the indent of rhyolitic volcanic rocks into older sedimentary strata in the western margin of the SIB ridge along a 300 degree-trending gulley located west of Gap Lake at ~6275350N (UTM). The thickness of sedimentary strata overlying the Betty Creek Formation rocks increases abruptly immediately to the northeast, across the northwest-trending cross fault running past drill pad D. This implies that this structure is older and affected deposition in the area. However, the northwestern extension of this particular fault structure appears to not produce significant offset in the upper sections of the younger rhyolite-rich stratigraphy, which suggests that most movement along it occurred prior to the eruption of the rhyolitic rocks.

In the east, the northwest-striking structures have been cut and shifted by north-northeast to northeast-striking shears and faults contained in the intermediate volcanic rocks of Composite Unit 1 (interpreted as Betty Creek Formation). This further reinforces the interpretation of these faults being early structures. However, a minority of northwest-striking faults crosscut the northeast-trending structures. These exceptions may have seen later re-activation by the stress regime responsible for the long-lived Jurassic-age Harrymel fault structure and/or by the later Cretaceous age Skeena event.

Older Pre-Skeena Bedding and Ridge Sub-Parallel Faults

Northeast to north-northeast trending faults are present cutting the older intermediate and mafic volcanic rocks of Composite Unit 1 (Betty Creek Formation?) on the eastern side of the SIB ridge. Similar structures cut the younger stratigraphy of Composite Units 2 and 3 (Salmon River Formation). These faults form an intersecting network or mosaic and in places provide bounds between rhyolitic and sedimentary rocks. Field observations indicate they exercised a control on rhyolite emplacement.

This group of faults displays a range of strike orientations, from northeast to more north-northeast. Plotting poles to planes on a stereo net two sets of fault structures are present. One set clusters around an average strike of 036°, which is close to the average attitude of the S1 foliation. A second less populous group clusters around an attitude of 012°. The northeast-trending shear and fault structures are developed sub-parallel to bedding and S1 foliation; dip attitudes are generally somewhat

steeper than bedding. In the west, within the younger rhyolite-dominated stratigraphy, these structures tend to display steep westerly dips.

Kinematic indicators located within and in proximity to this set of shears and faults include kink bands and crenulation and drag folds, striae, lunate cracks, poorly formed ancillary shears.

The general orientation of most of the axes of kink bands and crenulation folding is a shallow plunge to the northeast or southwest. Their axes sit approximately within a plane with a similar attitude to the average S1 foliation. This geometry implies a movement direction on these older shears and faults was sub-parallel to the dip of the bedding with movement of the eastern wall upwards and generally slightly northward in relation to the western wall. Approximating the direction of principal stress and movement direction as being perpendicular to the crenulation fold and kink band axes gives a range of potential orientations from $\sim 327^\circ$ to 292° (as projected in plan). This orientation is partially in line with the steeply south-raking mineral elongation observed in the S1 foliation plane ($\sim 335^\circ$).

Striae, lunate cracks, and minor occurrences of ancillary shears were also observed along fault planes. Several clusters of striae occur on stereo net, of which the most populous group is located close to the average fault plane great circle ($036^\circ/78^\circ$) and are sub horizontal in attitude. These striae record an episode of strike-slip, which was of a sinistral nature as indicated by the geometry of ancillary shears, lunate cracks, and steps in striae. Estimating the principal stress direction, assuming it acted at about 30° to the fault plane, yields a roughly northward to north-northwest directed stress vector. This would coincide with the potential stress field required to produce the sinistral movement observed on the large regional-scale Harrymel fault structure reported by Lewis (2001).

Several other groupings of fault striae are apparent on stereonet plots. One set plunges sub vertically in the fault plane indicating a phase of dip slip movement on some fault structures. Another minor set of striae (oriented $65\text{--}122^\circ$) is sub parallel to the more northwesterly projected stress vector inferred from the crenulation fold axes ($\sim 292^\circ$) and to the mineral elongation observed in the plane of the S1 foliation.

Rare shallow plunging slickensides suggest there was a phase of dextral strike slip on select northeast-trending faults. This may possibly have been a phase of re-activation produced by later more west-directed compression related to the Skeena Fold and Thrust event and Coulter Creek Thrust Fault.

The amount of slip along the northeast and north-northeast-trending faults is difficult to determine due to the absence of definitive marker beds. Mapping of recessive lineaments interpreted as expressions of older north-northwest trending faults suggests that movement is limited to a magnitude of <30 m on those northeast-striking faults located west of drill pad A. However, the magnitude of slip on wider structures such as the Mercury Anomaly Lineament, located on the eastern flank of the SIB ridge could potentially be of greater magnitude.

In 2008 after the examination of the southern third of the SIB ridge it was concluded that the northeast-striking fault structures supplanted the north-northeast-trending ones. After the 2010 mapping this assessment can no longer be supported. A portion of northerly-trending structures are cut by northeast-trending ones, but in 2010

more examples were found which are the opposite; where north-northeast striking faults tend to offset northeast-trending ones. There is a tendency for the former to eventually curve into the latter.

The northeast to north-northeast striking shears and faults were long lived structures. Bartsch (2001b) and Lewis (2001) describe northeast –striking and eastward-dipping structures in the area of the Eskay Creek mine, which display normal movement. Mapping the SIB ridge it is apparent that a portion of these structures provided a control on the rhyolitic volcanism. Evidence is allowing that part of the northeast and north-northeast-trending faults formed early as a result of transtensional extension along the Eskay rift. The contrasting north-northeast versus more northeast-striking fault sets may in part reflect an early geometry produced by normal extensional faulting (northeast structures) linked by relay ramping (north-northeast). The bedding sub-parallel attitude of the northeast-trending fault and shear structures as observed within the older rocks of Composite Unit 1 (Betty Creek Formation?) suggests that they formed after these older rocks had already been rotated and perhaps partially deformed.

Most of the observed kinematic indicators described above are likely later products of more ductile-brittle deformation post-dating the Eskay rift and associated extensional faulting.

The shallowly dipping striae combined with lunate cracks and ancillary shears record a phase of sinistral strike slip movement on the northeast and north-northeast-trending faults. They are interpreted to have been produced by a north to north-northwest directed principal stress vector and related to the regional scale Harrymel fault, which experienced sinistral strike-slip movement in Jurassic time.

The elongation lineation within the S1 foliation observed within the older rocks of Composite Unit 1 and to a lesser degree in the sediments of Composite Unit 2 is suggestive of a more ductile style of deformation. This is also true of the crenulation folding and foliation boudinage observed along northeast-trending fault structures. Such a ductile phase of deformation need have occurred with the rocks reasonably confined by burial. This period of deformation and fault movement was post Salmon River Formation emplacement in age. It may have been related to a compressional stress regime contemporaneous with continuing movement on the regional-scale Harrymel fault or less likely by a later northwest-oriented stress regime related to the Cretaceous-age Skeena Fold and Thrust event.

Coulter Creek Thrust Fault and Cretaceous Skeena Event Deformation

The Cretaceous-age Skeena Fold and Thrust event was superimposed on earlier faulting and deformational events. It rotated and folded the bedding in the volcanic and sedimentary stratigraphy of the Hazelton Group and rotated and re-activated earlier shears and faults. Sedimentary rocks of the Bowser Lake Group were extensively folded and thrust faults were emplaced in Jurassic and older stratigraphy.

The S2 cleavage and the orientation of fold axes in the Bowser Lake Group can be used to provide an estimate of the principal stress direction that produced the Skeena deformation. Mapping in 2008 and 2010 has augmented the cleavage and fold orientation data available. Fold axial planar cleavage directions in the Bowser Lake Group display a range of orientations from roughly north-south (~345°) to

about 040° with steep to sub vertical dips. The S2 cleavage in Hazelton Group rocks is oriented between 008° and 45° dipping steeply eastward at >75°.

Some of the tendency toward more northeasterly-strike orientations (>025° azimuth) in S2 within the Hazelton Group may locally have been influenced by the earlier S1 foliation. However, this phenomenon is unlikely to have occurred in the Bowser Lake Group rocks, which display a range of between 345° and 035° in the strike direction of axial planar cleavage and fold axes. Therefore, it must be concluded that the principal stress producing the Skeena event deformation in the area of the SIB property varied from a northwest (~310°) to a more west-northwest (275°) orientation.

The Coulter Creek thrust fault is related to the Cretaceous-age Skeena Fold and Thrust Event (Lewis, 2001). Its surface trace is generally recessive and it is inferred to lie somewhere between Coulter Creek and the lower west-facing slope of the SIB Ridge where older, volcanic-dominated Salmon River Formation rocks contact younger Bowser Lake Group sedimentary rocks. Only minor weak thrust splays have been found in outcrop along the bank of Coulter Creek; however, deeper diamond drill holes emplaced in 2002, 2008, and 2010 have intersected a significant shallow dipping thrust plane.

Projections of the fault trace on plan and on section using drill hole intercepts suggests a curvilinear surface trace - trending southeastward in the southern part of the mapped area and north northeastward north of the Lulu area. The structure is presently interpreted to have a shallow eastward to southeastward dip of between 25 and 40°. This shallow dip is in marked contrast to the steeply dipping faults and shears found on the eastern slope of the SIB ridge described above.

No evidence of the Coulter Creek thrust fault was encountered in the upper valley of Coulter Creek located at the northwestern periphery of the SIB claims. Lewis (2001) depicts it being truncated by the Argillite Creek in the area of 6275000N in the valley of Coulter Creek.

In the southwest portion of the property the Coulter Creek thrust fault has cut stratigraphy and moved older Betty Creek and Salmon River Formation rocks against and over younger Bowser Lake Group rocks. Deep drill holes emplaced during the 2008 exploration season show that altered rhyolite-rich Salmon River rocks occur in the lower thrust plate beneath older Salmon River sedimentary rocks and older intermediate volcanic rocks of the Betty Creek Formation.

Estimates of the movement direction along the thrust fault were attempted during the 2008 field season. It was inferred that folding of the Bowser Lake sedimentary rocks was largely produced by the Cretaceous-age Skeena deformation. The estimate of movement direction was based on the orientation of folding and the axial planar cleavage in the younger Bowser Lake Group sedimentary rocks lying near the Lulu Zone. The orientation of the axial planar cleavage and fold axes in this area are approximately north-south and the data suggests the Coulter Creek thrust moved in a 275° direction in this area.

North-northwest striking (340° to 350°) fault structures tend to produce sinistral offsets on both the Betty Creek and Salmon River stratigraphy, and cut the pre-Skeena bedding sub-parallel shears and faults previously described, including the Hexagon Zone lineament. Striae on fault planes tend to plunge shallowly to sub horizontally northward. Dip attitude is generally sub-vertical to steep northeast. Examples of these structures occur just north and near the south end of Gap Lake; and between Gap and SIB Lake on the south flank of Battleship Knoll.

The continuity and geometry of these faults varies along strike. A portion appear to curve into early, pre-existing 290° to 310° striking structures or into the more north-northeasterly trending sub portion of pre-Skeena ridge sub parallel faults. Displacements (in the horizontal plane) of the stratigraphy on these late northwest-striking faults, estimated from map pattern, are on the order of 20 to 50 m.

East-northeast striking (045° to 060°) faults were interpreted in the southern portion of the SIB ridge in 2008. These structures produce offsets of the Salmon River to Betty Creek contact and of lineaments related to the older north-northeast and northeast-striking faults and shears described above. The traces of these faults across topography are fairly linear but slightly concave towards the south, which suggests that they dip steeply southeastward.

Displacements on these faults are difficult to gauge. One such structure appears to displace contacts dextrally on the order of ~10 to 40 m. The orientation of these structures is similar to that of later cross faults located in the Eva Creek-Spearhead Area on the Corey Property. This latter structure is known to have displacements on the order of at least 0.5 km.

Argillite Creek Fault

The Argillite Creek fault consists of a major moderately to steeply west-dipping structure that has been mapped by Lewis (2001), Bartsch (2001b), and other workers in the vicinity of the Eskay Creek mine and interpreted southwestwards into the northern portion of the SIB claims.

Examination of the upper valley and head of Coulter Creek revealed several sets of moderately to steeply east and west-dipping splay faults with striae indicating dip-slip movement. The east-dipping structures can be interpreted as ancillary conjugate faults formed along the main west-dipping structure. In all, however, outcrop exposure in the valley of Coulter Creek is limited and an exact definition of the trace of the Argillite Creek fault and determination of attitude remained elusive.

Examination of argillaceous sedimentary units on the northwest-facing slope of the SIB ridge on the east bank of Coulter Creek reveals a trend toward steeply west-dipping to sub vertical bedding. This is what would be expected by downward dragging from a west-dipping normal fault.

Presently, the timing of movement along the Argillite Creek fault is suspected to be late or post-Skeena in age. Post-cretaceous-age extensional episodes have occurred in northwest B.C., notably in the Tertiary as cited by Logan et al. (2000).

DISCUSSION AND SUMMARY OF RESULTS

Mapping during the 2008 and 2010 field seasons resulted in the following:

1. Deposition of older intermediate and mafic tuffaceous volcanoclastic rocks (Composite Unit 1) took place below wave base but in relatively shallow water as attested by the presence of worm burrows in tuffaceous sandstone layers.
2. The older rocks of Composite Unit 1 are interpreted to be part of the Betty Creek Formation. However, no definitive fossil or geochronological data has been found on the SIB property to support this. This leaves the possibility that these rocks may be significantly older than now thought.
3. The transition zone sedimentary rocks (Composite Unit 2), which overlie older intermediate to mafic volcanoclastic rocks of the Betty Creek Formation are made-up of cobble and boulder conglomerate, dark carbonaceous argillaceous siltstone and sandstone, and grey to greenish wacke sandstone. Individual sedimentary beds are lensey, and, with the exception of graphitic siltstone, tend to lack bedding internally. Angular unconformable contacts between pebble and cobble conglomerate layers and finer-grained siltstone and sandstone, and the presence of boulder conglomerate with eroded, rounded, intermediate volcanoclastic clasts up to several metres across suggest that sedimentation took place on a slope. Eroded rounded clasts of intermediate volcanoclastic in the conglomerate units are similar in composition to the older rocks lying below and were well lithified at time of erosion. The presence of these eroded clasts suggests there was a significant hiatus between the emplacement of the Betty Creek rocks and the conglomeritic rocks.
4. Altered cobbles found within conglomeritic rocks of Composite Unit 2 are equivalent to altered older intrusive and volcanoclastic rocks found on Battleship Knoll and near the SIB Adit. This suggests that the alteration in these rocks pre-dates the emplacement of the Salmon River Formation.
5. Both east and west dipping, discontinuous bedding is present within the sedimentary sequence lying above Composite Unit 1 (Betty Creek Formation?). This geometry is likely the result of sedimentation taking place along a slope, which may have steepened over time, producing a fan of sediments, which thickens down dip and likely northeastward as well.
6. Low-Ti rhyolitic flows and autoclastic breccias within composite unit 3 are assigned to the Salmon River Formation. They contain intercalated, lensey layers of argillaceous siltstone and fine-grained sandstone. Mixtures of brecciated rhyolite and disrupted fine-grained argillaceous sediment accompanied the flow rocks and represent lava rubble and ripped-up sedimentary detritus forming carapaces and basal blankets to the flow rocks. The brecciated, mixed, and re-deposited debris suggest eruption of rhyolitic magmas onto and into partially unconsolidated argillaceous sediments. Teardrop-shaped flow-fronts suggest that rhyolitic rocks were extruded and advanced west to northwestward (in present terms) down a faulted paleo-slope. The discontinuous nature of the Lulu sulphide mineralization is probably due in part to the unstable character of the depositional environment.

7. Low-Ti Rhyolitic volcanic rocks display shallow to steeply west-dipping, and northerly to east-northeast striking flow contacts. The attitudes of the flow contacts within the rhyolite stratigraphy mimic the ancient paleo-topography. Attitudes of rhyolitic flow rocks are steeply dipping in the vicinity of fault traces and possess shallower dips more peripherally to these fault structures, which is suggestive of rhyolite emplacement into a rifted set of stepped, blocks. This poses further constraints on the down-dip continuity of any potential sulphide mineralization.
8. The rhyolite-rich stratigraphy displays a similar architecture as that found in the sedimentary sequence underlying it. There is a variation in the bedding dip angles from steeply eastward to westward moving from east to west down slope towards Coulter Creek. This suggests a thickening wedge of rhyolite and sediment down dip.
9. Overall, the stratigraphy encountered on surface in the southwest SIB area is generally equivalent to that found at the Eskay Creek mine, based on lithogeochemistry and the type of stratigraphic units present. The major exception is the absence of overlying tholeiitic pillow basalts. However, tholeiitic basaltic intrusive rocks, pillow lavas, volcanoclastic rocks were encountered below the Coulter Creek thrust fault in deep drill holes.
10. Three sets of faults appear to control the deposition and possibly extrusion of the low-Ti rhyolitic rocks: northeast or ridge axis sub-parallel ($\sim 036^\circ$); north-northeast ($\sim 010^\circ$); and northwest (290 to 310°).
11. Relatively thick sequences of intercalated rhyolite flows and breccias make-up the western brow of the SIB ridge lying north of the Lulu Zone. Individual layers of rhyolite flow and breccia units are generally <10 m wide. Massive rhyolitic sub-volcanic intrusive rocks occur locally and are a minor component of the stratigraphy exposed on surface. Larger volumes of rhyolitic flow rocks with thicker individual flow units are present in the extreme northern end of the claim group and the scale of occurrence there is of a scale that fits Bartsch's (2001b) the description of a rhyolite flow dome complex.
12. Larger examples of tholeiitic mafic sub volcanic intrusive rocks occur in the northern section of the SIB property within the area containing thicker rhyolitic flows described above. This suggests that this area is more proximal to a magmatic centre.

Summary of Structures and Structural Interpretations

The following observations and interpretations were made as a result of the mapping during the 2008 and 2010 exploration program:

1. The Hazelton Group rocks contain a well developed bedding-sub parallel tectonic foliation (S1), which has an average attitude of $\sim 040^\circ/67^\circ\text{E}$. The foliation tends to be steeper than bedding found in the older volcanoclastic rocks of Composite Unit 1 located in the eastern part of the map area.

2. Older fault structures oriented northwest cut the stratigraphy. They display striae indicating dip-slip and may mark abrupt changes in facies and lithology. These structures have partially controlled sedimentation and defined the contact of rhyolitic volcanic rocks in places. These faults are cut by later ridge sub parallel northeast-trending faults and north-northeast-striking faults.
3. An overlapping series of older northeast to north-northeast-striking faults and shears cut but partially sub parallel the S1 foliation and bedding in the older Hazelton Formation rocks. Two orientations of faults and shears are present averaging: $036^{\circ}/78^{\circ}$ and $012^{\circ}/75^{\circ}$. The emplacement of rhyolitic volcanic rocks of the Salmon River Formation (Composite Unit 3) was in part controlled by these fault structures, which may have originated in a transtensional regime as linked extensional faults.
4. The northeast and north-northeast oriented structures tend to displace the north to north-northwest oriented faults (described above) sinistrally, usually for distances of <30 m. Kinematic indicators suggest that several phases of movement, ranging from sinistral strike-slip to reverse dip-slip, occurred along these shears and faults. The principal stress direction that produced the sinistral strike slip along these structures likely ranged in orientation between west-northwest to northwest. This stress regime may have coincided with that which produced strike slip motion on the major regional Harrymel fault lying further to the west.
5. The S1 foliation in the Hazelton Group rocks is overprinted by a weaker tectonic foliation (S2), which has an average attitude of $027^{\circ}/80^{\circ}\text{E}$ but ranges in strike between $\sim 008^{\circ}$ and $\sim 045^{\circ}$. The spaced cleavage in younger Bowser Lake Group rocks, lying in the valley of Coulter Creek is likely contemporaneous with S2. It is oriented between $\sim 345^{\circ}$ and $\sim 040^{\circ}$ and on average dips sub vertically to steeply east. The S2 foliation and spaced cleavage in the Bowser Lake rocks was produced by the Cretaceous-age Skeena Fold and Thrust event.
6. The range in strike directions of the S2 cleavage reflects a principal stress direction that varied from northwest to west-northwest.
7. The Cretaceous-age Skeena event rotated and partially overturned older fault structures and bedding within the Hazelton and Bowser Lake Group.
8. The Coulter Creek thrust fault is a shallow east-dipping structure, which is part of the Skeena event and cuts the Hazelton Group and Bowser Lake Group stratigraphy in the southwest SIB area. It juxtaposes the older rhyolitic rocks of the Salmon River Formation against the sediments of the Bowser Lake Group in the southwestern part of the SIB property but does not extend into the northern part of the property.
9. The direction of movement on the Coulter Creek thrust fault was taken to be perpendicular to the axial planar cleavage (S2) in Hazelton Group rocks and the cleavage and fold axes developed in the Bowser Lake Group sedimentary rocks. Using the orientation data gathered in the Lulu Zone area a movement direction of 275° was estimated.

10. Later north-northwest striking faults displace Hazelton and Bowser Lake Group rocks sinistrally. These structures were formed in part along older, pre-existing northwest-striking structures. Shallow striae suggest a phase of strike slip, which could have been caused by the same stress regime as that responsible for the Skeena event.

2010 EXPLORATION ON THE MITCHELL PROPERTY

Objectives and Area Covered

A geological reconnaissance was performed between August 29th and September 9th 2010 of the western portion of the Mitchell Glacier claim block. This section of the property abuts Silver Standard Resources Iron Cap property, which is presently being actively explored and contains extensive gossan after copper-gold-bearing disseminated and veinlet-style mineralization. Given the current exploration activity on adjacent ground it was decided to investigate the possible potential for economic copper-gold mineralization of the western part of the Mitchell Glacier property. Reconnaissance mapping and sampling work was performed by the author, D. Carstens, and K. Doyon. Mapping and sampling was conducted in the area roughly bounded by 426500 to 428000 E; and 6267900 to 6269900 N UTM (Map 2, in pocket).

Geology of the Western Portion of the Mitchell Glacier Claims

Outcrop exposures are generally extensive at the upper levels (above 1600 m elevation) of the area examined. The lower slopes down the edge of Mitchell Glacier are generally talus covered. The property is underlain by volcanic, intrusive, and sedimentary rocks of the early- to mid-Jurassic-aged Hazelton Group.

Stratigraphy

Bedrock in the area of interest consists of three main rock types: felsic and intermediate tuffs; fine-grained, argillaceous sediments; and medium-grained intermediate hornblende-feldspar-porphyritic intrusive rocks. Several series of tuffaceous rocks are present and consist of intermediate lapilli, lapilli-ash, and epiclastic tuff breccia to lapillistone, with locally intercalated felsic lapilli-ash tuff.

In the area examined, an upper series of tuffaceous rocks occupies the higher elevations of the east-west-oriented ridge to the north of Mitchell Glacier (Map 2, in pocket). It is composed of grey-green locally maroon coloured, felsic lapilli tuff; grey rhyolitic lapilli-ash tuff (Photo Plate 11; and epiclastic, heterogeneous, felsic to intermediate, tuffaceous and lithic volcanic pebble sandstones (Photo Plate 12), lapillistones, and breccias. These upper rocks display north-northwest to north-northeast-striking (025° to 350°) bedding. Local layers of ignimbritic and welded tuff were noted. Dip angles vary with a range of shallow (25°) to moderately steep (70°)

east or west-dipping bedding orientations. Cross-bedding and angular disconformable contacts are common between individual tuff layers in the upper sequence. Minor examples of metre-scale lenses of rhyolitic to dacitic porphyritic volcanic or sub volcanic intrusive rock are contained in the upper series of tuffaceous rocks.

Niton portable XRF analyses of felsic tuffaceous and lava rocks in the upper tuffaceous series indicate that these rocks possess high Zr/TiO₂ ratios (> 1000 to >3000; see Table 2) not dissimilar from those found in the Eskay rhyolites.

To the south and southeast at lower elevations above the Mitchell Glacier, a second older series of tuffaceous rocks occurs. These rocks are largely composed of intermediate (andesitic to dacitic) lapilli tuff and tuffaceous volcanic pebble sandstone. These rocks tend to exhibit northeast-striking (025° to 065°) and steeply east or west dip angles. Intermediate to felsic hornblende-feldspar porphyritic intrusive rocks (Photo Plate 13) form northeast-striking, steeply-dipping, finger-like, sill bodies up to ~50 m or more wide within the lower series of tuffaceous rocks.

Argillaceous sediments composed of muddy siltstone underlie the southwestern portion of the map area. Intercalated lenses of this rock type are also found within the lower, second series of intermediate tuffaceous rocks.

Table 2. Outcrop sample descriptions – Mitchell Property

Station No.	UTM Location E	UTM Location N	Sample No. *	Sample Analysis	Description	Zr/TiO ₂ Niton
GL1	427067	6267625		Hand Spec.	Niton Felsic lapilli ash tuff	523
GL2	426838	6267751		Hand Spec.	Niton Green felsic epiclastic tuff breccia	461
GL3	426833	6267762		Hand Spec.	Niton Felsic ash tuff and lapilli ash tuff	389
GL4	426823	6267819		Hand Spec.	Niton Maroon, hematitic felsic volcanic conglomerate	558
GL5	427007	6267655		Hand Spec.	Niton Grey felsic ash tuff	593
GL6	427296	6267528		Hand Spec.	Niton Biotite-feldspar-porphyritic felsic volcanic	3735
GL7	427333	6267545		Hand Spec.	Niton Biotite-feldspar-porphyritic felsic volcanic	2591
GL8	427416	6268539		Hand Spec.	Niton Green-grey intermediate lapilli tuff.	1209
GL9	427408	6267491		Hand Spec.	Niton Grey-green, banded, lapilli tuff	1140
GL10	427160	6266637		Hand Spec.	Niton Biotite-hornblende-feldspar-phyric intermediate intrusive	285
GL11	427246	6266645	611716	Grab	ICP + Niton Light grey-green, altered, intermediate lapilli ash tuff	174
GL12	427284	6266639		Hand Spec.	Niton Hornblende -feldspar-porphyritic, intermediate intrusive	307
GL13	427376	6266916		Hand Spec.	Niton Intermediate volcanic sandstone	232
GL14	427352	6266894	611719	Chip 1.5 m	ICP + Niton Skarnified argillite	209
GL15	427223	6266974	611717	Chip 6 m	ICP Altered intermediate lapilli ash tuff	N.D.
GL16	427210	6266934	611718	Chip 8 m	ICP Altered intermediate lapilli ash tuff	N.D.
GL17	427176	6266876	611720	Chip 10 m	ICP Altered intermediate lapilli ash tuff	N.D.
GL18	427171	6266881		Hand Spec.	Niton Intermediate epiclastic lapilli ash tuff	248

N.D. = NOT DONE

* see Appendix B for analytical results



Photo Plate 11. Macro photo of hand sample GL5 taken from weakly banded, felsic lapilli ash tuff from upper series tuffaceous stratigraphy.



Photo Plate 12. Macro photo of banded, intermediate tuffaceous lapilli and pebble sandstone from the lower series tuffaceous stratigraphy.



Photo Plate 13. Medium-grained hornblende-feldspar-porphyrific intermediate intrusive rock. These rocks form sill like bodies within the lower series tuffaceous stratigraphy. Sample displays weak chlorite alteration.

Structure

Original bedding attitude has been locally dragged by a series of steeply east- and west -dipping (70° to 90°), northeast-striking (035° to 060°) shear zones. Rotated lithic grains, internal foliations, and local pinnate fractures suggest left-hand strike slip has occurred on these structures.

A tectonic foliation was noted with attitudes that somewhat mimic the attitude of the shear northeast-striking structures. The strike of the foliation ranges from northerly (~000°) to northeast (065°); dip angles are generally steep (between 70° and 90°) east and west.

Numerous shallow to moderately steeply dipping quartz-rich tension veins were observed in the tuffaceous and intrusive lithologies. The orientation of these structures is roughly in line with an east-west to northwest-oriented compression and these may have formed in response to the Cretaceous Skeena fold and thrust event.

Alteration and Mineralization

Several rusty-weathered bands ranging from metre-scale to decameter widths are present within the eastern and southern portion of the area examined (Photo Plate 14). These zones occur within and alongside a portion of the northeast-striking shear zones previously described. The alteration tends to become minimal at higher elevation along the east-west oriented ridge-top above 1800 m elevation.

Alteration in the rusty-weathered bands consists of pale bleached sericitized and pyritized tuffaceous (Photo Plate 15) and locally intrusive rock. It is localized and patchy in nature, varying in intensity within the rusty weathered bands. Alteration intensity is controlled by ancillary shear fractures, foliation, and dragged bedding found within and beside the shear zones. Minor veinlets and silicification occur in places within the rusty weathered bands; only rare chalcopyrite was observed in places.

One example of skarnified (pyroxene and carbonate-rich) argillaceous sediment was observed near the contact with porphyritic intermediate intrusive rock (Photo Plate 16). Minor local pyritization accompanies the skarnification.

Other alteration types observed in (and in proximity) to shear structures and in tuffaceous and intrusive rocks were hematization and chloritization. Several strongly hematized, maroon-coloured, epiclastic tuffaceous breccia and lapillistone layers are present. It is unclear as to the origin of this alteration whether it be early hydrothermal or more syngenetic. Later chloritization was observed to crosscut the hematization in places and was also noted in porphyritic intrusive rocks and adjacent tuffaceous rocks.

No substantial gossanous alteration was observed in the western part of the property in proximity to the western border and the Silver Standard-controlled ground. In this area the upper elevations (>1700 m) consist of both green and maroon coloured tuffs and epiclastic rocks of the upper series. The initial mapping suggests that this section is largely underlain by argillaceous sedimentary rock.



Photo Plate 14. Rusty – stained sericite-pyrite- +/- quartz-altered bands are the typical alteration encountered on the Mitchell Glacier property. Alteration is developed locally along ancilliary fractures and dragged bedding within decameter-wide shear zones striking northeast.

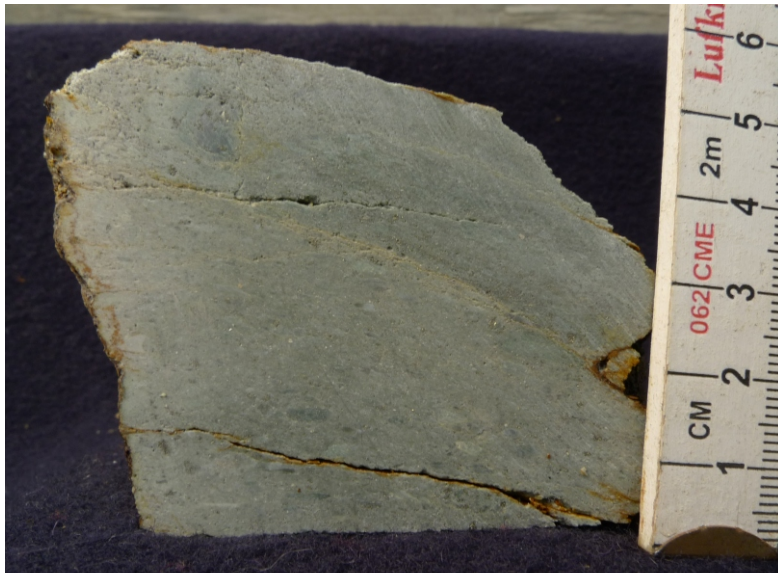


Photo Plate 15. Macro photo of GL11 sample material. Former banded epiclastic lapilli-ash tuff is altered by sericite and ~5 to 10% disseminated fine-grained pyrite. This style of alteration and mineralization is also typical of that found at the GL15 to GL17 sample sites.



Photo Plate 16. Skarn-altered, pyroxene (augite?)-carbonate-rich, former argillaceous siltstone is found along contacts of porphyritic intermediate intrusive rocks. Weak pyritization is present locally within the skarn zones. Photo is of sample GL14.

Sulphide mineralization observed in the rusty-weathered bands was largely pyritic in nature. Several chip samples were taken across the largest rusty-weathered, pyritized zone located approximately at 427200E; 626700N UTM.

Geochemical results were anomalous but insignificant for both precious metals and base metals. Values for arsenic and antimony are also anomalous.

Summary and Conclusions

Only limited sericitic (+/- quartz) alteration and pyritization were encountered in this initial mapping and sampling foray. Base and precious metal contents in the largest alteration zone identified are anomalous but economically insignificant.

The alteration and pyritization is spatially co-incidental with several northeast-striking shear zones. It is best developed at the relatively lower elevations on the south-facing slopes above the glacier.

The extensive gossan found on the adjacent Silver Standard Ltd. ground to the west appears to be limited by the large Bruce Jack Fault at the western periphery of the Mitchell Glacier property (see Map 2, in pocket). This structure has down-dropped the stratigraphy located to the east versus that located in the west. The pyritization found on the Mitchell Glacier claims may represent a higher more peripheral phase of mineralization than that found to the west of the Bruce Jack fault on Silver Standard Ltd.'s ground.

CONCLUSIONS

It is apparent that the Eskay Rift succession in the footwall to the Coulter Creek Thrust Fault is not an exact duplicate of the succession at the Eskay Creek mine. Rather, it appears that there may be rhyolites above and below the basalts. Thus, the basalt provides a potentially important marker unit to one's position in the rock sequence. Cambria suspect that the basalts also lie stratigraphically above the Lulu Zone position. Further to this, holes 139 & 140 (drilled from opposite directions on the same section less than 125 metres south of the Lulu Zone) reveal that the Salmon River Formation in the CCTF's footwall is at least 500 metres thick, a far greater thickness than is apparent from outcrop in the fault's hangingwall.

The Niton portable XRF device has allowed geologists to obtain real-time analytical data for base metals and many major and trace elements. This has proven particularly useful in differentiating the various volcanic rock units and has been a great assistance to the geologists. The ratios of Zr to Ti are readily determined using Niton data and are very useful in breaking out Betty Creek andesites and mafic rocks from Salmon River basalts and rhyolites.

RECOMMENDATIONS

A continuing co-ordinated, systematic program of diamond drilling, geological mapping, lithogeochemical sampling and geophysical data modeling are recommended. The stratigraphic mapping focus that was closely followed over the past two exploration seasons on the SIB claims should be continued as an adjunct to the lithogeochemical sampling program.

A downhole IP survey should be carried out as early as possible in order to characterize the geophysical response of the Lulu Zone mineralization and identify like-responding anomalies for future drill testing.

Drilling should focus on testing the Eskay Rift rocks beneath the Coulter Creek Thrust Fault. Given that the basalts encountered by the 2010 drilling in the footwall of the CCTF most likely lie stratigraphically above the Lulu Zone position, those parts of the footwall panel that are downsection from the basalts are priority targets and should be tested for Lulu Zone-equivalent mineralization.

The laterally extensive zone of Eskay-type rhyolite-hosted stockwork mineralization in holes EK08-134 and EK10-137 should be followed up and down dip in the wedge of the Salmon River Formation rhyolites mapped at surface. Again, a downhole IP survey should aid this.

In order to properly integrate the geologic mapping data and provide a more rigorous 3-D interpretation of the architecture and structural paragenesis of the SIB ridge it is necessary to perform a systematic re-interpretation of the geology in cross-section and plan view. Better differentiation of individual volcanic and sedimentary units on surface using lithogeochemistry and more consistent facies analysis has opened the possibility of re-interpreting older drill data and building a more accurate geologic interpretation of the geology in three dimensions. A more accurate geologic model would serve to provide additional constraints on future geophysical studies and deep drilling efforts.

Respectfully submitted,





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CERTIFICATE OF AUTHOR

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I, Sean D. McKinley, M.Sc., P.Geo. am a Professional Geoscientist residing in Burnaby, British Columbia, and do hereby certify that:

- I am a "qualified person" as defined in National Instrument 43-101: Standards of Disclosure for Mineral Projects ("NI 43-101") and my qualifications include the following:
 - I graduated from Queen's University, Kingston, Ontario in 1992 with a B. Sc. (Honours) degree in Geology.
 - I graduated from the University of British Columbia, Vancouver, B.C. in 1996 with an M. Sc. degree in Geology.
 - I am a Professional Geoscientist (P. Geo.) registered in the Association of Professional Engineers and Geoscientists of British Columbia, member# 28226, and have been a member in good standing since 2003.
 - From 1993 to the present, I have been actively engaged as a geologist in mineral exploration and geological research in British Columbia and Europe.
- I have worked on the Eskay (SIB) project on behalf of Eskay Mining Corp. as a consulting geologist from October 2008 to present, and have been responsible for compilation and interpretation of historical data and collaborating with the other professionals to establish the recommended exploration plan for the property.

Dated December 14, 2010.

Sean D. McKinley, M.Sc., P. Geo.

APPENDIX A

Eskay Property 2010 Drillhole Summary Geology Logs

Diamond Drill Hole Log

Company: Eskay Mining Corp.

Project: Eskay

Drillhole No.: EK10-136

Prospect: SIB

Start Date: 7/26/2010

Logged by: S. Tennant

End Date: 8/5/2010

Logged by:

Collar Azimuth: 297

UTM East (NAD83): 408090

Collar Dip: -74

UTM North (NAD83): 6273393

Hole Depth (m): 778.8

Elevation (m): 1072

Collar Location:

Drilling Contractor: Falcon Drilling

Collar Survey Type: Handheld GPS

Drill Model: Falcon-3000

Downhole Survey Type: Icefield Tool

Core Size: NQ

Comment: Targets stockwork veining in EK08-134 along strike and down dip.

Project: Eskay
Drill Hole ID: EK10-136

Drill Hole ID: EK10-136

[illegible]

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data													Survey					
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.
			40																	295.7	73.1	
	[1.9 - 73.4 m] Intermediate tuff/ash tuff (INTF). Domainal alteration gives a coarsely clastic look but sub-lapilli to rare lapilli sized degraded crystals and lithics, some vesiculated, dominate. Sheared and veined throughout		20																	295.5	73.2	
																					296.2	72.9
																					295.8	72.9
																					296.2	72.5
																					296.8	72.6

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data												Survey						
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.
			40 20																			
	[1.9 - 73.4 m] Intermediate tuff/ash tuff (INTF). Domainal alteration gives a coarsely clastic look but sub-lapilli to rare lapilli sized degraded crystals and lithics, some vesiculated, dominate. Sheared and veined throughout																					
	[73.4 - 122.7 m] Mafic tuff/ash tuff (MFTF). Sheared mafic sub-volcanic intrusion with multiple apothoses in an intermediate tuff screen. Slightly chilled and amygdular marings suggest emplacement into a "hot" screen. Chlorite-rich intervals occur, possibly representing altered mafic hyaloclastite.																					

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data														Survey					
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.	Dip
95	[73.4 - 122.7 m] Mafic tuff/ash tuff (MFTE). Sheared mafic sub-volcanic intrusion with multiple apothoses in an intermediate tuff screen. Slightly chilled and amygdular marings suggest emplacement into a "hot" screen. Chlorite-rich intervals occur, possibly representing altered mafic hyaloclastite.		40																				
			20																		296.8	71.9	
																						298.1	72
																						298.4	72
																						298.1	71.8
	[122.7 - 202.5 m] Mafic breccia (MFBX). Sheared mafic breccia/hyaloclastite(?). Micro-vesicular lapilli common in a chlorite & pyrite-rich, likely formerly glassy, matrix. Horizons of intense chloritisation and pyritisation occur intermittently, likely representing formerly vitriclast-rich portions.																						
																					297.6	71.8	
125																							

Project: Eskay
Drill Hole ID: EK10-136

Drill Hole ID: EK10-136

[illegible]

Project: Eskay
Drill Hole ID: EK10-136

Drill Hole ID: EK10-136

[illegible]

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data												Survey						
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.
			40 20																			
190	[122.7 - 202.5 m] Mafic breccia (MFBX). Sheared mafic breccia/hyaloclastite(?). Micro-vesicular lapilli common in a chlorite & pyrite-rich, likely formerly glassy, matrix. Horizons of intense chloritisation and pyritisation occur intermittently, likely representing formerly vitriclast-rich portions.	[190.7 - 197.1 m] Mafic intrusive (MFIV)																		298.8	71.3	
195																						
200																					298.5	71.3
205	[202.5 - 210.6 m] Sediment - mixed (SDMX). Very sheared mix of mudstone carrying lapillus to cobble sized mafic detritus and chloritised hyaloclastite with striking domainl silicification and pyritisation.																					
210																					300.2	71.3
215	[210.6 - 211.1 m] Structure - fault (STFT). Gouge and broken core with associated alteration in surrounding rock. Undulatin gouge seams are at a low angle to core axis. [211.1 - 369 m] dacite intrusive (DCIV). Patchily silicified, massive unit with rare clastic (in-situ brecciated?) horizons. Is this an intrusion?																			299.6	71.2	
																				298.9	71.3	

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data														Survey				
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.
	[211.1 - 369 m] dacite intrusive (DCIV). Patchily silicified, massive unit with rare clastic (in-situ brecciated?) horizons. Is this an intrusion?		40 20																			
																					298.9	71
																					299.3	70.8
																					298.8	70.9
																					300	70.7
																					299.9	70.7

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data													Survey					
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.
			40 20																			
	[211.1 - 369 m] dacite intrusive (DCIV). Patchily silicified, massive unit with rare clastic (in-situ brecciated?) horizons. Is this an intrusion?																					

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data														Survey				
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.
315	[211.1 - 369 m] dacite intrusive (DCIV). Patchily silicified, massive unit with rare clastic (in-situ brecciated?) horizons. Is this an intrusion?		40																	299.2	70.5	
320																				299.4	70.4	

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data													Survey					
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.
			40																			
			20																			

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data														Survey				
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.
	[370.6 - 485.2 m] dacite intrusive (DCIV). Same as 211.1-369.0m		40 20																			

Project: Eskay
Drill Hole ID: EK10-136

Drill Hole ID: EK10-136

[illegible]

Project: Eskay
Drill Hole ID: EK10-136

Drill Hole ID: EK10-136

[illegible]

Project: Eskay
Drill Hole ID: EK10-136

Drill Hole ID: EK10-136

[illegible]

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data													Survey					
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.
			40 20																			
	[485.2 - 552.9 m] Intermediate intrusive (INIV). Sheared, intermittently silicified in-situ brecciated intermediate intrusion																					
																					300.9	69.5
540																						
545																						
550																					302	69.8
555	[552.9 - 555.9 m] Structure - fault (STFT). Crushed quartz veins and gouge define fault in argillite																					
	[555.9 - 561.8 m] volcanic - mixed (VOMX). Argillite loaded with poorly to moderately ordered sub m scale horizons of epidote-chlorite rich intermediate (?) debris and possible mud rip-ups																				300.8	69.8
560																						
	[561.8 - 572 m] Intermediate tuff/ash tuff (INTF). Intermediate tuff; generally massive looking grey-green crystal tuff with more pumice (?) rich and chloritised hyaloclastite rich portions																					
565																					301.2	69.5

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data														Survey				
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.
630	[626 - 635 m] Sediment - mudstone (SDMD). Argillite with rare sub-cm pebbly beds and silty trubidite laminations		20																		302.5	68.8
635																						
640	[635 - 650.3 m] Basalt intrusive (BSIV). Sub-volcanic basalt intrusion becoming increasingly vesiculated/amygdular with chlorite-rich altered hyaloclastite rich horizons downhole cutt by rare exsolved po stringers																					
645																						
650	[650.3 - 682.1 m] Basalt pillowed (BSPL). Gradational change to vesiculated and quench brecciated basalt unit with rare pillow fragments and much chloritised hyaloclastite																					
655																						
660																						

Project: Eskay
Drill Hole ID: EK10-136

Drill Hole ID: EK10-136

[illegible]

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data													Survey					
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.
	[682.1 – 778.8 m] Basalt intrusive (BSIV). Texturally variable basalt intrusion with coarsely spherulitic horizons, then massive and fine grained, followed by glomerophyric then in-situ brecciated/coarsely vesiculated and brecciated downhole		20																			
			40																			
695																					302.1	68.4
700																					301.6	68.6
705																					301	68.3
710																					300.9	68.3
715																						
720																					301.6	68.5

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data														Survey					
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.	Dip
			40 ----- 20																				
	[682.1 – 778.8 m] Basalt intrusive (BSIV). Texturally variable basalt intrusion with coarsely spherulitic horizons, then massive and fine grained, followed by glomerophyric then in-situ brecciated/coarsely vesiculated and brecciated downhole																						
725																							
730																					301.1	68.7	
735																							
740																					301.8	68.4	
745																							
																						300.7	68.7
																						301.9	68.4
																				302.1	68.4		

Project: Eskay
Drill Hole ID: EK10-136

Drill Hole ID: EK10-136

[illegible]

Diamond Drill Hole Log

Company: Eskay Mining Corp.

Project: Eskay

Drillhole No.: EK10-137

Prospect: SIB

Start Date: 8/5/2010

Logged by: S. Tennant

End Date: 8/17/2010

Logged by:

Collar Azimuth: 300

UTM East (NAD83): 408094

Collar Dip: -69

UTM North (NAD83): 6273588

Hole Depth (m): 776.8

Elevation (m): 1119

Collar Location:

Drilling Contractor: Falcon Drilling

Collar Survey Type: Handheld GPS

Drill Model: Falcon 3000

Downhole Survey Type: Icefield Tool

Core Size: NQ

Comment: Northerly stepout targeting stockwork intersected in drill hole EK08-134.

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data															Survey			
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.
			40 20																		300.1 300.1	69.4 69.4
	[0 - 1.5 m] Drillhole casing (DHCS)																					
	[1.5 - 166.1 m] Intermediate lapilli tuff (INLT). Variably altered, sheared and veined grey-green to dark green volcanoclastic with granular to washed out fine and coarse lapilli-rich intervals and possible carb ax FSp(?) hrzs. Some pseudoclastic "coarsening" of unit as indicated usual shear induced fabric defined by ax ~ 50 degrees TCA and rare "windows" through ax eg. 55.7m																					

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data														Survey				
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.
	[1.5 - 166.1 m] Intermediate lapilli tuff (INLT). Variably altered, sheared and veined grey-green to dark green volcanoclastic with granular to washed out fine and coarse lapilli-rich intervals and possible carb ax FSP(?) hrzs. Some pseudoclastic "coarsening" of unit as indicated usual shear induced fabric defined by ax ~ 50 degrees TCA and rare "windows" through ax eg. 55.7m		40																			
			20																			

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data													Survey		
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)
	[1.5 - 166.1 m] Intermediate lapilli tuff (INLT). Variably altered, sheared and veined grey-green to dark green volcanoclastic with granular to washed out fine and coarse lapilli-rich intervals and possible carb ax FSP(?) hrzs. Some pseudoclastic "coarsening" of unit as indicated usual shear induced fabric defined by ax ~ 50 degrees TCA and rare "windows" through ax eg. 55.7m	[50.6 - 73.2 m] Intermediate crystal tuff (INXT)	40 20 -----																

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data													Survey							
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.	Dip	
	[182.9 - 462.4 m] Sediment - siltstone (SDSI). Black silty argillite, fining downhole with variable shear and nonshear related QZC veining, overall overprint of wk chl alteration. Also with much sub-ordinate volcanic-derived pale grey. Graded, sands and more disordered pebbly mixed sediment. Some AX, possibly associated with shearing and veining, including interesting fg py stringers.	[189.5 - 190.7 m] Sediment breccia (SDBX)	40																					
190			20																					
195																						299.5	67.4	
200																						299	67.5	
205																						299.3	67.3	
210																						299.4	67.3	
215																						299.1	67.7	

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data													Survey							
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.	Dip	
	[182.9 - 462.4 m] Sediment - siltstone (SDSi). Black silty argillite, fining downhole with variable shear and nonshear related QZC veining, overall overprint of wk chl alteration. Also with much sub-ordinate volcanic-derived pale grey. Graded, sands and more disordered pebbly mixed sediment. Some AX, possibly associated with shearing and veining, including interesting fg py stringers.		20 40																					

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data														Survey				
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.
	[182.9 - 462.4 m] Sediment - siltstone (SDSI). Black silty argillite, fining downhole with variable shear and nonshear related QZC veining, overall overprint of wk chl alteration. Also with much sub-ordinate volcanic-derived pale grey. Graded, sands and more disordered pebbly mixed sediment. Some AX, possibly associated with shearing and veining, including interesting fg py stringers.		40 20																			
																					300.4	67.5
255																						
260																					299.2	67.7
265																						
270																					300.6	67.4
275																					299.4	67.6
280																					300.8	67.3
																					300.8	67.3

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data													Survey					
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.
	[182.9 - 462.4 m] Sediment - siltstone (SDSI). Black silty argillite, fining downhole with variable shear and nonshear related QZC veining, overall overprint of wk chl alteration. Also with much sub-ordinate volcanic-derived pale grey. Graded, sands and more disordered pebbly mixed sediment. Some AX, possibly associated with shearing and veining, including interesting fg py stringers.		40																			
285			20																			
290																					300.9	67.3
295			[295.4 - 296.7 m] volcanic - sandstone (VOSA)																		300.6	67.2
			[296.7 - 297.6 m] Sediment - mixed (SDMX)																			
300																					300.2	67.4
305																						
310		[308.1 - 309.1 m] volcanic - sandstone (VOSA)																		299.8	67.4	
				611605	313.4	314.4	1		20.5		18.7		45		46.4	0.5		161.2		2.1	300.5	67.4
				611606	314.4	315.4	1		19.9		40.2		26		13.1	0.8		112.3		6		

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data															Survey									
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.	Dip					
315	[182.9 - 462.4 m] Sediment - siltstone (SDSI). Black silty argillite, fining downhole with variable shear and nonshear related QZC veining, overall overprint of wk chl alteration. Also with much sub-ordinate volcanic-derived pale grey. Graded, sands and more disordered pebbly mixed sediment. Some AX, possibly associated with shearing and veining, including interesting fg py stringers.			611606	314.4	315.4	1		19.9		40.2		26		13.1		0.8	112.3		6								
320																						300.2	67.4					
325																							299.3	67.2				
330																								299.3	67			
335																									299.3	66.9		
340																										299.3	66.9	
345																											300.3	66.9

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data														Survey					
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.	Dip
	[182.9 - 462.4 m] Sediment - siltstone (SDSI). Black silty argillite, fining downhole with variable shear and nonshear related QZC veining, overall overprint of wk chl alteration. Also with much sub-ordinate volcanic-derived pale grey. Graded, sands and more disordered pebbly mixed sediment. Some AX, possibly associated with shearing and veining, including interesting fg py stringers.		40																				
— 350																					301	66.8	
— 355																					300	66.8	
— 360																							
— 365																					300.8	66.6	
— 370																					300.5	66.4	
— 375			[374.6 - 378.8 m] Sediment - sandstone (SDSA)																		299.7	66.2	

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data													Survey					
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.
	[182.9 - 462.4 m] Sediment - siltstone (SDSI). Black silty argillite, fining downhole with variable shear and nonshear related QZC veining, overall overprint of wk chl alteration. Also with much sub-ordinate volcanic-derived pale grey. Graded, sands and more disordered pebbly mixed sediment. Some AX, possibly associated with shearing and veining, including interesting fg py stringers.	[374.6 - 378.8 m] Sediment - sandstone (SDSA)	40																			
		[378.8 - 379 m] Sediment - sandstone (SDSA)	20																			
380																					299.7	66.1
			[379 - 386.3 m] Sediment - sandstone (SDSA)																		300.4	66.2
385																						
390																						
395			[386.3 - 403 m] Sediment - sandstone (SDSA)																		300.9	65.9
400																					299.8	65.9
405																					300.6	65.9

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data													Survey						
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.	Dip
410	[182.9 - 462.4 m] Sediment - siltstone (SDSi). Black silty argillite, fining downhole with variable shear and nonshear related QZC veining, overall overprint of wk chl alteration. Also with much sub-ordinate volcanic-derived pale grey. Graded, sands and more disordered pebbly mixed sediment. Some AX, possibly associated with shearing and veining, including interesting fg py stringers.		40																				
			20																		299.7	65.7	
																						301.1	65.7
																						300.7	65.6
																						300.8	65.5
																						301.1	65.3

Project: Eskay
Drill Hole ID: EK10-137

Drill Hole ID: EK10-137

[illegible]

Project: Eskay

Drill Hole ID: EK10-137

[illegible]

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data														Survey				
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.
			40																			
			20																			

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data													Survey					
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.
			40 20																			
	[544.5 - 569.1 m] volcanic - sandstone (VOSA). As in 516.5m -> fine grained > 554.3m with "dirty" look and py of 508.7m.																					
	[569.1 - 572.8 m] Structure - fault (STFT). Serious fault zone in argillite and v. fine sandstone. Much deformed QC veins and cataclasite particularly 571.7-572.2m, at ~70 degrees TCA.																					
	[572.8 - 588.1 m] volcanic tuff/ash tuff (VOTF). Complex interval of sheared, variably altered felsic? ash tuffs, tuffs and rare lapilli tuffs with argillite lam in "lower" portion (>~583.7m). Bedding is consistently ~45 degrees TCA but this may be shear-induced fabric more obvious bedding > 583.7 is at 75-80 degrees TCA.	[572.8 - 573.9 m] Felsic tuff/ash tuff (ESTF)																				

Project: Eskay

Drill Hole ID: EK10-137

[illegible]

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data														Survey						
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.	Dip	
630	[611 – 728.5 m] Rhyolite flow (RHFL). Autoclastic rhyolite with sections of flowbanded autoclasts, bands to aligned spherulites, and sparse 2-3mm quartz eye sections. The quartz eyes are more prominent between 615.3m to 638.2m. Also, 1-2m sections of core speckled white, possibly the result of 1mm sized phenocrysts. 637.2–693.5m, overall the unit gets lighter in colour from med grey to light green and grey.		40 20																					
635																								
640																								
645																								
650																								
655																								
660																								

Project: Eskay
Drill Hole ID: EK10-137

Drill Hole ID: EK10-137

[illegible]

Project: Eskay

Drill Hole ID: EK10-137

[illegible]

Project: Eskay
Drill Hole ID: EK10-137

Drill Hole ID: EK10-137

[illegible]

Project: Eskay
Drill Hole ID: EK10-137

Drill Hole ID: EK10-137

[illegible]

Diamond Drill Hole Log

Company: Eskay Mining Corp.

Project: Eskay

Drillhole No.: EK10-138

Prospect: SIB

Start Date: 8/17/2010

Logged by: A. Ramsay

End Date: 8/31/2010

Logged by:

Collar Azimuth: 292

Collar Location:

UTM East (NAD83): 408193

Collar Dip: -64

UTM North (NAD83): 6273100

Hole Depth (m): 847.3

Elevation (m): 1010

Drilling Contractor: Falcon Drilling

Collar Survey Type: Handheld GPS

Drill Model: Falcon 3000

Downhole Survey Type: Icefield Tool

Core Size: NQ

Comment: Eastern extension of the Lulu Zone in the footwall of the Coulter Creek Thrust on same section as 2008 drilling, azimuth and dip adjusted to allow for hole deviation.

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data												Survey						
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.
			40																	292.4	64	
	[1.5 - 42.6 m] Intermediate lapilli tuff (INLT). Intermediate lapilli tuff with crystal rich sections, an apparent coarsening/pseudofragmental appearance due to chl/sericite alt. Lapilli are typically elongate, few mm thin and 0.5-2cm long. Unit is sheared.		20																	292.7	63.9	
																				292.8	63.9	
																				292.7	63.8	
																				293.3	63.8	
	[42.6 - 50.4 m] Mafic crystal tuff (MFXT). Mafic crystal tuff, sheared with mm sized possible feldspar lath crystal and minor lapilli. Gradational upper and lower contacts.																			292.5	63.8	
																				292.9	63.7	
	[50.4 - 72.7 m] Intermediate lapilli tuff (INLT). Same as 1.5-42.6m; med to dark green, possibly less sericite than chl. More visible lapilli, few mm to 3mm in size.																			293.1	63.6	
																				293.5	63.6	
																				293.2	63.5	
																				293.1	63.4	

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data													Survey					
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.
			40 20																			
	[50.4 - 72.7 m] Intermediate lapilli tuff (INLT). Same as 1.5-42.6m; med to dark green, possibly less sericite than chl. More visible lapilli, few mm to 3mm in size.																					293.6 / 63.4
																						294.3 / 63.3

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data													Survey				
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)
95	[89 - 97.1 m] volcanic - sandstone (VOSA). Green, medium grained volcanic sandstone, fining downhole with sections of coarse fragments (conglomerate). Sand is 1-2mm with 2mm sized possible feldspar crystals.	[91.9 - 95.7 m] volcanic conglomerate (VOCG)	20																	296.3	62.1
100	[97.1 - 106.7 m] Intermediate tuff/ash tuff (INTF). Beige to light green sheared fine grained intermediate tuff. Lower contact missing in lost core.	[102.3 - 103.2 m] Intermediate lapilli tuff (INLT) [103.4 - 103.8 m] volcanic - sandstone (VOSA)																		295.5	62.2
105	[106.7 - 120.4 m] volcanic - sandstone (VOSA). Volcanic sandstone, sheared and some crenulation cleavage in sections. Gradational lower contact.																			294.7	62.3
110																					
115																				295.2	62.2
120	[120.4 - 237.9 m] Mafic - massive (MFMS). Dark to medium green mafic massive, fine to medium grained with 1-5cm shards? or possible domainal alteration. Strong chl and minor sericite/carbonate alt. Also, zones with 1-3cm white possible feldspar crystals in both matrix and shards?. Minor 4-10m sections with amygdulies.																			294.6	62.2
125																				295.3	62.2

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data														Survey				
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.
	[120.4 - 237.9 m] Mafic - massive (MFMS). Dark to medium green mafic massive, fine to medium grained with 1-5cm shards? or possible domainal alteration. Strong chl and minor sericite/carbonate alt. Also, zones with 1-3cm white possible feldspar crystals in both matrix and shards?. Minor 4-10m sections with amygdules.		40 ----- 20																			
																					296.6	61.8
																					295.7	61.6
																					296.4	61.5
																					296.5	61.4
																					296.6	61

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data														Survey								
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.	Dip			
	[120.4 - 237.9 m] Mafic - massive (MFMS). Dark to medium green mafic massive, fine to medium grained with 1-5cm shards? or possible domainal alteration. Strong chl and minor sericite/carbonate alt. Also, zones with 1-3cm white possible feldspar crystals in both matrix and shards?. Minor 4-10m sections with amygdules.		40																							

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data														Survey					
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.	Dip
	[120.4 - 237.9 m] Mafic - massive (MFMS). Dark to medium green mafic massive, fine to medium grained with 1-5cm shards? or possible domainal alteration. Strong chl and minor sericite/carbonate alt. Also, zones with 1-3cm white possible feldspar crystals in both matrix and shards?. Minor 4-10m sections with amygdules.		40 20																				
190																							
195																							
200																					299.2	58.3	
205																					300.1	57.8	
210																					299.1	57.4	
215																					301.5	56.9	
																					299.7	56.8	

Project: Eskay
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[illegible]

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data												Survey								
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.	Dip	
315	[281.9 - 374.1 m] Mafic breccia (MFBX). Mottled dark grey and beige mafic breccia with 0.5-1.0m subrounded to subangular, poorly sorted, mafic clasts/fragments. The fragments have some micro amygdules and are mainly loosely packed with lesser areas of moderate packing. Matrix is fine dark grey/green chloritic. Also, towards the bottom of the unit the appearance of few cm to 1m thick massive lobes with chill margins or possible shattered boundaries. Upper contact hidden by veining.		40 ----- 20 ----- 0																					
																								</

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data														Survey						
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.	Dip	
			40 20	611615	377.1	378.1	1		72.6		38		14		0.5		1.6		622.3		24.1			
	[374.1 - 378.1 m] Sediment - mudstone (SDMD). Black fine grained argillite, sheared with minor brecciations. Upper contact sheared with above unit ~55 degrees TCA and sharp lower contact at 75 degrees TCA. [378.1 - 425.1 m] Mafic - massive (MFMF). Light to medium grey green, fine to medium grained mafic massive with signs of shearing. Possible 1-4mm chl amygdules elongate and aligned with shearing. Sarp lower contact at 55 degrees TCA with 1cm thick veins at low angle TCA at contact.																							
			[378.1 - 396.4 m] Mafic breccia (MFBX)																					

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data												Survey							
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.	Dip
410	[378.1 - 425.1 m] Mafic - massive (MFMS). Light to medium grey green, fine to medium grained mafic massive with signs of shearing. Possible 1-4mm chl amygdules elongate and aligned with shearing. Sarp lower contact at 55 degrees TCA with 1cm thick veins at low angle TCA at contact.		40 20 -----																				
415																							
420																							
425																							
430	[425.1 - 436.5 m] Altered volcanoclastic (AXVC). Grey, silica and sericite altered volcanoclastic with few mm to 2cm, subangular to subrounded, weakly chl clasts, poorly sorted, and matrix supported unit. Sheared unit as clasts aligned with shearing. Also, few 10cm intervals of SDMD; argillite; interbedded in unit.	[425.1 - 426.5 m] Sediment - mixed (SDMX)																					
435																							
440	[436.5 - 480.9 m] Intermediate tuff/ash tuff (INTF). Pale green, fine grained, sheared with strong-mod chlorite/sericite. 1-5cm shards? or dominal alteration. Also, altered 1-3mm white abraided feldspar(?) lathes and fine scoria fragments, 5% concentration. Towards bottom of unit appearance of fine grained, massive lobes/epophases 0.5-1.0m in size. Lower contact sheared.	[432.9 - 433.7 m] Mafic - massive (MFMS)																					

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Drill Hole ID: EK10-138

[illegible]

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data														Survey				
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.
	[480.9 - 545.1 m] Intermediate - massive (INMS). Green, fine grained, massive with patches of ghostly insitu breccia enhanced by stronger chl and epidote alt. Below 537.7m, unit is more brecciated and an increase in the amount of shear related veining.		40																			
			20																			
505																						
510																				304.4	52	
515																				304.2	51.8	
520																				304.4	51.7	
525																						
530																				303.3	51.7	
																				305.2	51.4	

Project: Eskay
Drill Hole ID: EK10-138

[illegible]

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data												Survey					
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)
630	[628.3 - 634.3 m] Felsic tuff/ash tuff (FSTF). Light green, felsic tuff, fine grained ash and rare few mm to 1cm lapilli. Also, minor sections of coarser medium grained tuff with rare lapilli. Towards lower contact 10cm sections of argillite and fine sediments.	[629.9 - 631.7 m] Mafic - sandstone (MFSA)	40 20																		
635	[634.3 - 639.8 m] Mafic lapilli tuff (MFLT). Green mafic lapilli tuff with chl and carbonate alteration. Lapilli are possible dark glass(?) fragments 4-7mm sized and sparse. Also, minor qtz eyes. Sharp upper and lower contacts.	[637.6 - 638.8 m] Felsic lapilli tuff (FSLT)																			
640	[639.8 - 644.2 m] Felsic lapilli tuff (FSLT). Light green, felsic lapilli tuff, fine grained matrix with few mm to 2cm sized, generally elongate pumic and other compositions lapilli, 10-20%. Towards lower contact becomes more tuffaceous.	[641.4 - 642.9 m] Mafic - sandstone (MFSA)																			
645	[644.2 - 650.6 m] Felsic - sandstone (FSSA). Light grey, felsic pebbley sandstone with subangular to subrounded, moderate sorting and generally matrix supported with few mm to 1cm sized clasts. Unit fines downhole with fewer 1cm clasts and more sand 1-2mm grains. Also, a few 1cm thing argillit beds towards contact, see sub lithology description.																				
650	[650.6 - 651.3 m] Structure - fault (STFT). Catclasite fault with gouge and breccia zones. Also, qtz-carbonate veining. [651.3 - 664.9 m] Felsic - sandstone (FSSA). Same as 644.2-650.6m. Sharp upper and lower contacts. Upper contact is associated with a veing hiding contact and lower contact is at high angle TCA.	[648.5 - 650.6 m] Sediment - sandstone (SDSA) [651.3 - 652 m] Sediment - mudstone (SDMD)																			
655																					
660																					

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data													Survey					
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.
	[651.3 - 664.9 m] Felsic - sandstone (FSSA). Same as 644.2-650.6m. Sharp upper and lower contacts. Upper contact is associated with a veing hiding contact and lower contact is at high angle TCA.		00																			
			02																			
665	[664.9 - 679.4 m] Felsic tuff/ash tuff (FSTF). Light grey, fine grained bedded with possible laded bedding sections. Also, minor zones where beds are sheared. Sublitho contact is sharp and loaded indicating younging downhole, therefore unit is possibly overturned. Also towards bottom of unit there are zones of more lapilli rich tuff.																					
670		[669.2 - 670.2 m] Felsic lapilli tuff (FSLT)		611616	669.9	670.9	1		15.5		8.7		107		9.1	0.2			0.8			0.6
				611617	670.9	671.9	1		36.3		15.6		111		1.7	0.1			0.9			0.9
				611618	671.9	672.9	1		45.2		45.8		116		1.3	0.3			0.9			0.8
				611619	672.9	673.9	1		33.3		11.9		96		4	0.1			0.5			0.4
				611620	673.9	674.9	1		35.2		8.8		131		2.7	0.1			0.5			0.7
675				611621	674.9	675.9	1		35.1		42.2		146		4.7	0.2			19.3			0.9
				611622	675.9	676.9	1		33.3		16.5		149		2.3	0.1			11.1			0.7
				611623	676.9	677.9	1		37.2		12.8		152		5.1	0.1			2.8			0.7
680	[679.4 - 680 m] Structure - fault (STFT). Crumbly, minor gouge fault with 10-20cm of argillite. [680 - 683.7 m] Mafic - massive (MFMS). Grey brown, fine grained, massive, mafic.																					
685	[683.7 - 684.1 m] Structure - fault (STFT). Crumbly, minor gouge and veining with sections of argillite. [684.1 - 716.6 m] Mafic pillowed (MFPL). Light to medium brown green mafic pillowed flow with zones of pillowed breccia. Noticable pillows with dark green altered selvages starts at 688.6m and continues downhole. Also, sections of spherulites and chl/qtz(?) amygdulues. Mineralization is associated with the dark green pillow selvages and within martris of the breccia.																					
690																						

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data													Survey					
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.
	[716.6 - 771.2 m] Mafic - massive (MFMS). Mottled green and tan massive mafic with zones of insitu breccia and fine breccia along the more massive lobe margins. Breccias are enhanced by alteration along fractures. Also, zones of chl and qtz amygdules found within fragments and some massive portions. In addition, zones with spherulites and devitrification "colliflower" texture. Within the top 9 meters of core is mottled green and white with coarse grained texture, possible cumulate mafic with a minor shear zone.		40																			
				20																		

Project: Eskay

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[illegible]

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data														Survey				
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.
			40																			
	[818 - 824.6 m] Rhyolite breccia (RHBX). Same as 809.8-816.1m		20																			
				</																		

Diamond Drill Hole Log

Company: Eskay Mining Corp.

Project: Eskay

Drillhole No.: EK10-139

Prospect: SIB

Start Date: 9/1/2010

End Date: 9/18/2010

Logged by: A. Ramsay

Logged by:

Collar Azimuth: 112

Collar Dip: -55

Hole Depth (m): 637

Collar Location:

UTM East (NAD83): 407354

UTM North (NAD83): 6273366

Elevation (m): 907

Drilling Contractor: Falcon Drilling

Drill Model: Falcon 3000

Core Size: NQ

Collar Survey Type: Handheld GPS

Downhole Survey Type: Icefield Tool

Comment: Along-strike step-out from 2008-2010 "Lulu Extension" drilling

Project: Eskay
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Drill Hole ID: EK10-139

[illegible]

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data														Survey				
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.
			20																			
			40																			

Project: Eskay
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Drill Hole ID: EK10-139

[illegible]

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data														Survey					
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.	Dip
	[62.4 - 253.8 m] Sediment - siltstone (SDS1). Same as 2.2-61.7m with less sandy units and possibly finer grained in sections towards the top of unit. Few meter sections of blocky broken core with minor gouge typically 5-10m apart, such as 70.0-81.4m blocky broken core with minor gouge. 66.5-111.0m variable bedding angles from 25-80 degrees TCA due to shearing, folding and soft sediment deformation. Unit has less definitive sedimentary structures. Also, from 213.0 to the bottom of the unit there is an increase of interbedded 10-40cm sections of medium grained sandstones units approx. 20-80cm apart.		40																				
			20																				
130																							
135																							
140																							
145																							
150																							
155																							

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data														Survey					
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.	Dip
	[62.4 - 253.8 m] Sediment - siltstone (SDS1). Same as 2.2-61.7m with less sandy units and possibly finer grained in sections towards the top of unit. Few meter sections of blocky broken core with minor gouge typically 5-10m apart, such as 70.0-81.4m blocky broken core with minor gouge. 66.5-111.0m variable bedding angles from 25-80 degrees TCA due to shearing, folding and soft sediment deformation. Unit has less definitive sedimentary structures. Also, from 213.0 to the bottom of the unit there is an increase of interbedded 10-40cm sections of medium grained sandstones units approx. 20-80cm apart.		40																				
			20																				

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data														Survey					
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.	Dip
	[62.4 - 253.8 m] Sediment - siltstone (SDS1). Same as 2.2-61.7m with less sandy units and possibly finer grained in sections towards the top of unit. Few meter sections of blocky broken core with minor gouge typically 5-10m apart, such as 70.0-81.4m blocky broken core with minor gouge. 66.5-111.0m variable bedding angles from 25-80 degrees TCA due to shearing, folding and soft sediment deformation. Unit has less definitive sedimentary structures. Also, from 213.0 to the bottom of the unit there is an increase of interbedded 10-40cm sections of medium grained sandstones units approx. 20-80cm apart.		40																				
			20																				
190																							
195																							
200																							
205																							
210																							
215																							

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data														Survey					
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.	Dip
	[62.4 - 253.8 m] Sediment - siltstone (SDS1). Same as 2.2-61.7m with less sandy units and possibly finer grained in sections towards the top of unit. Few meter sections of blocky broken core with minor gouge typically 5-10m apart, such as 70.0-81.4m blocky broken core with minor gouge. 66.5-111.0m variable bedding angles from 25-80 degrees TCA due to shearing, folding and soft sediment deformation. Unit has less definitive sedimentary structures. Also, from 213.0 to the bottom of the unit there is an increase of interbedded 10-40cm sections of medium grained sandstones units approx. 20-80cm apart.		40																				
			20																				
— 225																							
— 230																							
— 235																							
— 240																							
— 245																							
— 250																							

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data												Survey						
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.
	[62.4 - 253.8 m] Sediment - siltstone (SDS1). Same as 2.2-61.7m with less sandy units and possibly finer grained in sections towards the top of unit. Few meter sections of blocky broken core with minor gouge typically 5-10m apart, such as 70.0-81.4m blocky broken core with minor gouge. 66.5-111.0m variable bedding angles from 25-80 degrees TCA due to shearing, folding and soft sediment deformation. Unit has less definitive sedimentary structures. Also, from 213.0 to the bottom of the unit there is an increase of interbedded 10-40cm sections of medium grained sandstones units approx. 20-80cm apart. [253.8 - 288.1 m] Sediment - sandstone (SDSA). Grey interbedded, medium grained sandstone with pebbly sandstone and siltstone (argillite) interbeds. Also between 267.7-284.4m there are 40-60cm sections of silty matrix supported fragmental with subangular to subrounded fragments/clasts of variable compositions; mafic to sandstone. These units may be the result of slumping in a slurry of debris flows. The sandstone sections tend to be medium grained, strong to moderately sorted with localized graded bedding such as fining upwards sequences and some planar beds associated with siltstone units. Upper and lower contacts are gradational.		40																			
			20																			
255																						
260																						
265																						
270																						
275																						
280																						

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data														Survey						
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.	Dip	
315	[303.1 - 315.1 m] Sediment - siltstone (SDSI). Same as 2.2-61.7m, dark and light grey, bedded with fine scale slumping, soft sediment deformation, and thin 5-10cm sparse medium grained sandstone beds. [315.1 - 330.6 m] Sediment - sandstone (SDSA). Grey medium grained, pebbly sandstone with 20-30cm sections of coarse clast rich sandstone polymite and 10cm sections of siltstone. Variable non definitive grading, over length of unit. Some of the more conglomerate like units are matrix supported with the SDSI unit possibly due to slumping and debris flows.		40 20 -----																					
320																								
325																								
330	[330.6 - 396 m] Sediment - siltstone (SDSI). Same as 2.2-61.7m, dark grey and light grey bands with less sandy units downhole and sections of more mud rich units with no visible bedding. Also, towards the bottom of the unit bedding is variable due to slumping, folding, and soft sediment deformation. Upper contact is hidden by veining and lower contact is gradational. Between 375.6-377.4m zones of 10-40cm with silty/muddy, mafic fragments surrounded by silty matrix. Fragments are subangular to angular up to 1cm size and matrix supported. Also, between 388.4-389.2m 10-20cm sections similar to above of few mm to 2cm sized fragments with variable composition, some clasts are massive py.																							
335																								
340																								
345																								

Project: Eskay

Drill Hole ID: EK10-139

[illegible]

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data														Survey				
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.
— 410	[405.4 - 412.4 m] Mafic intrusive (MFIV). Green brown, medium grained massive mafic intrusive with zone of white to beige spherulites 2-6mm in size giving a coarser grained texture. Visible fine grained leucoxene and possible biotite. Upper and lower contacts are marked by fine grained chill margins and insitu brecciation. [412.4 - 455 m] Rhyolite intrusive (RHIV). Same as 399.4-405.4m. Lower contact marked by brecciation.		40 20																			
— 415																						
— 420																						
— 425																						
— 430																						
— 435																						
— 440																						

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data													Survey						
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.	Dip
	[412.4 - 455 m] Rhyolite intrusive (RHIV). Same as 399.4-405.4m. Lower contact marked by brecciation.		40 20 -----																				
445																							
450																							
455	[455 - 460.4 m] Mafic porphyry (MFPH). Grey feldspar phyrlic mafic porphyry with 10-50cm sections of insitu and crackle breccia, enhanced by alteration and veining. White feldspar phenocrysts are anhedral, 2-6mm in size and approx. 5%. The brecciation is more prominent towards the upper and lower contacts, in addition, rock is finer grained and less porphyritic near contacts indicating chill margins.																						
460	[460.4 - 465.2 m] Rhyolite intrusive (RHIV). Same as 399.4-405.4m with altered and vein rich zones surrounding contacts with mafic intrusion. Also, less to no quartz phenocrysts.																						
465	[465.2 - 467.4 m] Mafic intrusive (MFIV). Similar to 455.0-460.4m, but no phenocrysts and possible fine, chl amygdulles. Fine grained chill margins with sharp irregular contacts. Believed to be the same as above MFPH, but a thinner interval. [467.4 - 574 m] . Similar to 399.4-405.4m, contains brecciated zones, minor local flow banding, and zones of devitrification/spherulites, more prominent below 557.0m. [467.4 - 636.7 m] Rhyolite intrusive (RHIV). Similar to 399.4-405.4m, contains brecciated zones, minor local flow banding, and zones of devitrification/spherulites, typically found towards the lower portion of the unit. Also, more massive rhyolite sections are associated with the mafic subunit around the contacts, possibly due to contact metamorphism.	[464.1 - 464.3 m] Mafic intrusive (MFIV)																					
470		[471 - 471.4 m] Mafic intrusive (MFIV)																					

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data														Survey						
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.	Dip	
	[467.4 - 574 m] . Similar to 399.4-405.4m, contains brecciated zones, minor local flow banding, and zones of devitrification/spherulites, more prominent below 557.0m. [467.4 - 636.7 m] Rhyolite intrusive (RHIV). Similar to 399.4-405.4m, contains brecciated zones, minor local flow banding, and zones of devitrification/spherulites, typically found towards the lower portion of the unit. Also, more massive rhyolite sections are associated with the mafic subunit around the contacts, possibly due to contact metamorphism.		40																					
			20																					

Project: Eskay
Drill Hole ID: EK10-139

Drill Hole ID: EK10-139

[illegible]

Project: Eskay
Drill Hole ID: EK10-139

Drill Hole ID: EK10-139

[illegible]

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data														Survey					
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.	Dip
630	[467.4 - 636.7 m] Rhyolite intrusive (RHIV). Similar to 399.4-405.4m, contains brecciated zones, minor local flow banding, and zones of devitrification/spherulites, typically found towards the lower portion of the unit. Also, more massive rhyolite sections are associated with the mafic subunit around the contacts, possibly due to contact metamorphism.		40 20 -----																				
																						108.6	51.6
635																							
640																							
645																							
650																							
655																							
660																							

Diamond Drill Hole Log

Company: Eskay Mining Corp.

Project: Eskay

Drillhole No.: EK10-140

Prospect: SIB

Start Date: 9/19/2010

End Date: 10/1/2010

Logged by: A. Ramsay

Logged by:

Collar Azimuth: 292

Collar Dip: -63

Hole Depth (m): 816.8

Collar Location:

UTM East (NAD83): 408181

UTM North (NAD83): 6272945

Elevation (m): 1022

Drilling Contractor: Falcon Drilling

Drill Model: Falcon 3000

Core Size: NQ

Collar Survey Type: Handheld GPS

Downhole Survey Type: Icefield Tool

Comment: Along-strike extension of Lulu Zone equivalent stratigraphy, taking into account ~150m south strike-slip displacement of footwall on Coulter Creek Thrust

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data														Survey				
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.
			40 20																		292.2 292.2	63.9 63.9
	[0 - 1.2 m] Drillhole casing (DHCS). Overburden, casing.																					
	[1.2 - 34.5 m] Intermediate lapilli tuff (INLT). Light to medium green intermediate lapilli tuff to intermediate tuff.. Also contains zones with a pseudoclastic/shardy appearance due to domainal alteration. Overall unit is shear with visible foliation and local crenulation cleavage. 1-2mm possible lapilli elongate and aligned with shearing/foliation. Lower gradational contact.																				296.1 296.1	60.7 60.7
																					292.2 292.2	63.9 63.9
																					293.2 293.2	63.5 63.5
																					294.1 294.1	63.1 63.1

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data													Survey							
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.	Dip	
	[47.6 - 196.1 m] Intermediate lapilli tuff (INLT). Light to medium green, same as 1.2-34.5m, but with zones of intermediate crystal tuff. Crystals are typically white, lath shaped, 1-3mm feldspar crystals approx. 10-20% concentration. Also, some visible local few mm thin elongate pumice lapilli. Compositional variation from intermediate to mafic due to volcanoclastic nature of the rock, no specific visual boundaries. Upper and lower contacts are gradational.		40 20																					
																						295.3	61.6	
																							295.7	61.6
																							295.2	61.3
																							295.4	60.5

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data												Survey								
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.	Dip	
95	[47.6 - 196.1 m] Intermediate lapilli tuff (INLT). Light to medium green, same as 1.2-34.5m, but with zones of intermediate crystal tuff. Crystals are typically white, lath shaped, 1-3mm feldspar crystals approx. 10-20% concentration. Also, some visible local few mm thin elongate pumice lapilli. Compositional variation from intermediate to mafic due to volcanoclastic nature of the rock, no specific visual boundaries. Upper and lower contacts are gradational.		20																					

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data												Survey						
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.
			40 20																			
	[47.6 - 196.1 m] Intermediate lapilli tuff (INLT). Light to medium green, same as 1.2-34.5m, but with zones of intermediate crystal tuff. Crystals are typically white, lath shaped, 1-3mm feldspar crystals approx. 10-20% concentration. Also, some visible local few mm thin elongate pumice lapilli. Compositional variation from intermediate to mafic due to volcanoclastic nature of the rock, no specific visual boundaries. Upper and lower contacts are gradational.																					

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data												Survey									
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.	Dip		
190	[47.6 - 196.1 m] Intermediate lapilli tuff (INLT). Light to medium green, same as 1.2-34.5m, but with zones of intermediate crystal tuff. Crystals are typically white, lath shaped, 1-3mm feldspar crystals approx. 10-20% concentration. Also, some visible local few mm thin elongate pumice lapilli. Compositional variation from intermediate to mafic due to volcanoclastic nature of the rock, no specific visual boundaries. Upper and lower contacts are gradational.		40 20																		294.5	57.2			
195																							296.9	57	
200																								295.2	57.2
205																								296.9	56.9
210																								296.2	56.8
215	[196.1 - 202.9 m] Mafic tuff/ash tuff (MFTF). Green, fine grained mafic tuff to possible mafic lapilli tuff with possible qtz amygdules or shear related qtz blobs. Local few mm thin elongate chl altered possible lapilli. Upper contact is gradational and lower contact hidden within large qtz vein.																								
	[202.9 - 324.6 m] Intermediate tuff/ash tuff (INTF). Light to medium green with local beige sections; similar to 1.2-34.5m but much less lapilli to no lapilli and more sections of chl dominal alteration giving rock a pseudoclastic/shardy appearance. Local compositional variation from intermediate to mafic due to the volcanoclastic nature of the rock and no visible textural changes.																								

Project: Eskay
Drill Hole ID: EK10-140

Drill Hole ID: EK10-140

[illegible]

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data														Survey						
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.	Dip	
315	[202.9 - 324.6 m] Intermediate tuff/ash tuff (INTF). Light to medium green with local beige sections; similar to 1.2-34.5m but much less lapilli to no lapilli and more sections of chl domainal alteration giving rock a pseudoclastic/shardy appearance. Local compositional variation from intermediate to mafic due to the volcanoclastic nature of the rock and no visible textural changes.		20																					
320																						297.3	55.1	
325				[324.6 - 358.6 m] Mafic intrusive (MFIV). Grey green to grey brown, fine grained massive mafic with minor shear fabric. 345.3-353.6m, stronger shear fabric/foliation with local crenulations, possibly enhanced by alteration and brecciation. Towards the bottom of the unit there are zones with mior breccia and a lower gradational contact.																			298.3	54.8
330																								297.2
335																						297.9	54.9	
340																						297.9	54.9	
345																								

Project: Eskay

Drill Hole ID: EK10-140

[illegible]

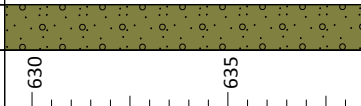
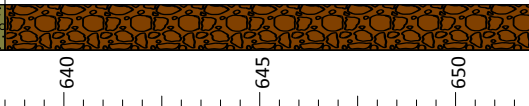
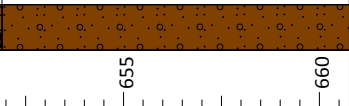
Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data													Survey					
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.
	[358.6 - 442.1 m] Mafic breccia (MFBX). Dark grey green to brown grey mafic breccia with local more massive mafic intrusive sections 10-50cm wide. Fragments are subangular, elongate, typically aligned, 1-10cm in size, and have mm thin chl amygdulites(?). Rock is matrix supported with dark green hyaloclastic chl altered matrix. Within the last 2m of the unit argillite mixed the mafic breccia, likely due to slumping.		40 ----- 20																			
																					297.9	54
380																						
385																					298.6	53.9
390																						
395																					297.9	53.5
400																					298.9	53.2
405																					298.1	53.1

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data														Survey			
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)
			40																		
			20																		

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data													Survey						
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.	Dip
	[530.5 - 539.1 m] Sediment - interbedded (SDIB). Dark to light grey interbedded sandstone and siltstone with intermittent beds of felsic tuff 10cm to 1m long towards the bottom contact.		40 20 -----																				
	[539.1 - 574.5 m] Felsic tuff/ash tuff (FSTF). Light green to grey, fine grained felsic tuff with local zones containing 1cm sized lapilli. Below 554.0m there is a variation in grain size and multiple sublithologies of siltstone and volcanic sandstone beds towards bottom contact. Eg. below 553.2m the felsic tuff gradually grades from sub mm tuff grains to 1mm sand sized grains, which grades into siltstone at 556.6. Also, possible slumping associated with the silty portions.																						
540																					300.8	50.5	
545																							
550																					301.5	50.3	
555																							
		[556.6 - 558.8 m] Sediment - siltstone (SDSI)																				302.2	50.1
560		[559.7 - 561.4 m] volcanic - sandstone (VOSA) [561.9 - 562.2 m] Sediment - siltstone (SDSI)																				302.2	50
565		[566 - 567.3 m] Sediment - siltstone (SDSI)																				302.2	49.7

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data													Survey						
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.	Dip
			40 20																				
	[539.1 - 574.5 m] Felsic tuff/ash tuff (FSTF). Light green to grey, fine grained felsic tuff with local zones containing 1cm sized lapilli. Below 554.0m there is a variation in grain size and multiple sublithologies of siltstone and volcanic sandstone beds towards bottom contact. Eg. below 553.2m the felsic tuff gradually grades from sub mm tuff grains to 1mm sand sized grains, which grades into siltstone at 556.6. Also, possible slumping associated with the silty portions.	[566 - 567.3 m] Sediment - siltstone (SDSI)																					
		[568.3 - 569 m] Sediment - siltstone (SDSI)																					
	[574.5 - 606.7 m] volcanic - sandstone (VOSA). Light green to grey pebbley volcanic sandstone, typically moderately sorted and matrix supported. Visible bedding with grading and local loading structures. Local zones of strong shearing made visible by strong sericite and chl alteration. These zones may contain more pumice clasts and less coherent clasts. Clasts are 3mm to 3.5cm long, subangular to subrounded with variable composition from intermediate to possible felsic volcanic tuff(?) clasts. Pumice fragments tend to be chl and/or sericite altered.																						

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data												Survey						
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.
			40 20																			
	[574.5 - 606.7 m] volcanic - sandstone (VOSA). Light green to grey pebbley volcanic sandstone, typically moderately sorted and matrix supported. Visible bedding with grading and local loading structures. Local zones of strong shearing made visible by strong sericite and chl alteration. These zones may contain more pumice clasts and less coherent clasts. Clasts are 3mm to 3.5cm long, subangular to subrounded with variable composition from intermediate to possible felsic volcanic tuff(?) clasts. Pumice fragments tend to be chl and/or sericite altered.																					
	[606.7 - 613.6 m] Mafic intrusive (MFIV). Green grey, fine grained, amygduloidal mafic intrusive with finer grained chill margins and local devitrification. Contains 10-30cm of VOSA beds within the intrusive due to interfingering by MFIV. Amygdules are chl and minor qtz. Upper contact is sharp and micro-faulted at low angle TCA resulting in a 10cm offset. Lower contact is sharp.																					
	[613.6 - 622 m] volcanic - sandstone (VOSA). Grey, fine grained, volcanic sandstone, moderate to well sorted, and with less pebbley sandstone sections as 574.5-606.7m. Contains subrounded, intermediate to felsic tuff and pumice clasts. Also, siltstone beds towards the bottom of the unit.																					
		[620.5 - 621.2 m] Sediment - siltstone (SDSI)																				
	[622 - 627.1 m] Mafic intrusive (MFIV). Same as 606.7-613.6. Fairly sharp contacts, and lower contact is brecciated.																					
	[627.1 - 638.5 m] volcanic - sandstone (VOSA). Similar to 574.5-606.7m. Light green, intermediate composition, and sheared with local crenulation cleavage. The crenulation cleavage is at low to moderate angles TCA. The shearing has caused the altered pumice fragments to resemble fiamme. Near the top and bottom contacts unit consists of finer sands.																					

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data												Survey							
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.	Dip
630	 [627.1 - 638.5 m] volcanic - sandstone (VOSA). Similar to 574.5-606.7m. Light green, intermediate composition, and sheared with local crenulation cleavage. The crenulation cleavage is at low to moderate angles TCA. The shearing has caused the altered pumice fragments to resemble fiamme. Near the top and bottom contacts unit consists of finer sands.		20	40																			
635																							
640				 [638.5 - 651.9 m] Sediment conglomerate (SDCG). Light green and green grey, polymict conglomerate. Generally matrix supporte towards the top section and more clast supported towards the bottom. Variable clast composition from felsic tuffs to porphyritic clast, to volcanic sandstone, and possible mafic clasts. Clasts are subrounded and have been stretched due to shearing and are several mm to ~30cm in size. Most clasts are chl to sericite altered. Local crenulation is visible.																		306.1	49
645																					306.5	48.9	
650																						307.3	48.8
655	 [651.9 - 663.6 m] Sediment - sandstone (SDSA). Grey, fine grained sandstone with zones of pebbley sandstone. Visible bedding and graded bedding showing between 654.7-655.4m the younging direction is up hole. Finer grained sands near top and towards bottom contact a 1m gradation into a more silt sized grains.																						
660																						307.6	48.7

Depth (m)	Major Lithology	Minor Lithology	% Sulphide	Assay & Geochemical (ICP) Data																Survey			
				Sample	From (m)	To (m)	Int. (m)	Cu (%)	Cu (ppm)	Pb (%)	Pb (ppm)	Zn (%)	Zn (ppm)	Au (g/t)	Au (ppb)	Ag (g/t)	Ag (ppm)	As (%)	As (ppm)	Sb (%)	Sb (ppm)	Azim.	Dip
725	[716.3 – 769.6 m] Mafic pillow breccia (MFPB). Light grey green to brown grey, mafic pillow breccia. For the most part this unit consists of insitu breccia with strong chl altered, hyaloclastite and locally fine Py matrix. Fragments are subangular to angular, several mm to 5cm wide of fine grained massive mafic pillow fragments. Also, contains zones of fine grained, massive pillows with strong chl altered pillow selvages. Pillows have devitrification textures toward outer margins and chl amygdules. Below 754.0m to end of unit, unit is more mafic breccia than pillowed breccia.		20	611727	723.3	724.3	1		43.8		3.4		110		0.5		0.1						
				611728	724.3	725.3	1		51.3		7.4		96		0.8		0.1				0.8		
730																				307.3	46.9		
735																				309.1	47		
740																							
745																				309	46.8		
750																				309.4	46.9		
																				309.8	46.9		

Project: Eskay

Drill Hole ID: EK10-140

[illegible]

APPENDIX B

Eskay Property 2010 Geochemical and Assay Data Tables

Drillcore Samples - Assays (Acme Labs - 7AR+G6 fire assay)

SAMPLE	HOLEID	From		Int.	Sample		LAB JOB	Ag (g/t)	Au (g/t)	Cu (%)	Mo (%)	Hg (%)	Bi (%)	Cr (%)	W (%)	Sr (%)	Na (%)	Cd (%)	S (%)
		(m)	To (m)		Type	(m)													
611501	EK10-136	-	-		Blank			<2	<0.005	<0.001	<0.001	<0.001	<0.01	0.001	<0.001	0.008	0.15	<0.001	<0.05
611502	EK10-136	451.4	452.3	0.9	Core			<2	0.05	0.009	0.002	<0.001	<0.01	<0.001	<0.001	0.004	0.05	<0.001	0.81
611503	EK10-136	459	459.5	0.5	Core			<2	0.049	0.04	<0.001	<0.001	<0.01	<0.001	<0.001	0.003	0.04	<0.001	1.09
611504	EK10-137	86.2	86.7	0.5	Core			<2	0.009	0.009	<0.001	<0.001	<0.01	<0.001	<0.001	0.008	0.02	<0.001	0.85
611505	EK10-137	30.4	30.8	0.4	Core			3	1.619	0.045	<0.001	<0.001	<0.01	<0.001	<0.001	0.004	0.03	0.003	1.55
611506	EK10-137	462	463	1	Core			2	0.196	0.011	<0.001	<0.001	<0.01	<0.001	<0.001	0.001	<0.01	<0.001	5.08
611507	EK10-137	463	464	1	Core			<2	0.542	0.008	<0.001	<0.001	<0.01	<0.001	<0.001	0.001	<0.01	<0.001	3.74
611508	EK10-137	464	464.5	0.5	Core			<2	0.71	0.002	<0.001	<0.001	<0.01	0.002	<0.001	0.002	<0.01	<0.001	1.2
611509	EK10-137	464.5	465	0.5	Core			<2	0.848	0.005	<0.001	<0.001	<0.01	<0.001	<0.001	0.001	<0.01	<0.001	2.1
611510	EK10-137	465	465.5	0.5	Core			<2	0.482	0.006	<0.001	<0.001	<0.01	<0.001	<0.001	0.001	<0.01	<0.001	1.23
611511	EK10-137	465.5	466	0.5	Core			11	0.814	0.047	<0.001	<0.001	<0.01	<0.001	<0.001	0.003	<0.01	<0.001	6.7
611512	EK10-137	466	466.5	0.5	Core			13	0.785	0.074	<0.001	<0.001	<0.01	<0.001	<0.001	0.004	<0.01	0.006	5.59
611513	EK10-137	466.5	467	0.5	Core			10	0.387	0.082	<0.001	<0.001	<0.01	0.001	<0.001	0.002	<0.01	0.001	4.11
611514	EK10-137	470	470.5	0.5	Core			11	1.267	0.086	<0.001	<0.001	<0.01	0.001	<0.001	0.003	<0.01	0.003	5.61
611515	EK10-137	470.5	471	0.5	Core			<2	0.155	0.038	<0.001	<0.001	<0.01	0.001	<0.001	0.002	<0.01	<0.001	0.89
611516	EK10-137	471	471.5	0.5	Core			11	0.37	0.044	<0.001	<0.001	<0.01	0.001	<0.001	0.009	<0.01	0.001	2.65
611517	EK10-137	471.5	472	0.5	Core			8	0.496	0.105	<0.001	<0.001	<0.01	0.001	<0.001	0.003	<0.01	<0.001	2.83
611518	EK10-137	472	472.5	0.5	Core			<2	0.079	0.027	<0.001	<0.001	<0.01	0.002	<0.001	0.003	<0.01	<0.001	0.78
611519	EK10-137	472.5	473	0.5	Core			25	2.355	0.339	<0.001	<0.001	<0.01	0.001	<0.001	0.004	<0.01	0.008	8.48
611520	EK10-137	473	474	1	Core			<2	0.085	0.006	<0.001	<0.001	<0.01	0.002	<0.001	0.002	<0.01	<0.001	0.99
611521	EK10-137	474	475	1	Core			2	0.294	0.011	<0.001	<0.001	<0.01	<0.001	<0.001	0.001	<0.01	<0.001	5.46
611522	EK10-137	467	467.5	0.5	Core			16	1.439	0.123	<0.001	<0.001	<0.01	<0.001	<0.001	0.002	<0.01	0.003	7.76
611523	EK10-137	467.5	468	0.5	Core			5	0.387	0.03	<0.001	<0.001	<0.01	<0.001	<0.001	0.001	<0.01	0.002	4.4
611524	EK10-137	468	468.5	0.5	Core			6	0.591	0.034	<0.001	<0.001	<0.01	0.001	<0.001	0.002	<0.01	0.001	7.94
611525	EK10-137	468	468.5	0.5	Dupl.			6	0.427	0.048	<0.001	<0.001	<0.01	0.002	<0.001	0.001	<0.01	0.001	5.12
611526	EK10-137	468.5	469	0.5	Core			2	0.437	0.015	<0.001	<0.001	<0.01	<0.001	<0.001	0.001	<0.01	<0.001	5.35
611527	EK10-137	469	469.5	0.5	Core			3	0.258	0.011	<0.001	<0.001	<0.01	0.001	<0.001	0.001	<0.01	<0.001	4.4
611528	EK10-137	469.5	470	0.5	Core			<2	0.129	0.009	<0.001	<0.001	<0.01	0.001	<0.001	0.003	<0.01	<0.001	2
611529	EK10-137	763.6	764.1	0.5	Core			<2	0.009	<0.001	<0.001	<0.001	<0.01	0.001	<0.001	<0.001	<0.01	<0.001	0.35
611530	EK10-138	72.7	73.7	1	Core			<2	0.437	0.007	<0.001	<0.001	<0.01	<0.001	<0.001	0.03	0.01	<0.001	1.41
611531	EK10-138	73.7	74.7	1	Core			2	1.744	0.026	<0.001	<0.001	<0.01	<0.001	<0.001	0.012	0.01	<0.001	3.42
611532	EK10-138	74.7	75.4	0.7	Core			<2	0.202	0.01	<0.001	<0.001	<0.01	<0.001	<0.001	0.022	0.02	<0.001	1.91
611533	EK10-138	556.4	556.9	0.5	Core			3	0.07	0.004	<0.001	<0.001	<0.01	<0.001	<0.001	0.004	0.01	<0.001	1.74
611534	EK10-138	556.9	557.8	0.9	Core			5	0.057	0.024	<0.001	<0.001	<0.01	<0.001	<0.001	0.004	0.02	<0.001	3.28
611535	EK10-138	557.8	558.4	0.6	Core			<2	0.04	0.005	<0.001	<0.001	<0.01	<0.001	<0.001	0.003	0.02	<0.001	2.44
611536	EK10-138	558.4	559.4	1	Core			3	0.077	0.004	<0.001	<0.001	<0.01	<0.001	<0.001	0.004	0.01	<0.001	2.68
611537	EK10-138	559.4	559.9	0.5	Core			4	0.105	0.003	<0.001	<0.001	<0.01	<0.001	<0.001	0.004	0.01	<0.001	1.78
611538	EK10-138	559.9	560.4	0.5	Core			3	0.391	0.009	<0.001	<0.001	<0.01	<0.001	<0.001	0.005	0.02	<0.001	3.19
611539	EK10-138	560.4	560.9	0.5	Core			<2	0.286	0.007	<0.001	<0.001	<0.01	<0.001	<0.001	0.007	0.01	<0.001	2.89
611540	EK10-138	560.9	561.5	0.6	Core			<2	0.041	0.005	<0.001	<0.001	<0.01	<0.001	<0.001	0.008	0.01	<0.001	1.45

Drillcore Samples - Assays (Acme Labs - 7AR+G6 fire assay)

SAMPLE	HOLEID	From (m)	To (m)	Int. (m)	Sample Type	LAB JOB	P (%)	K (%)	As (%)	Pb (%)	Zn (%)	Mn (%)	Fe (%)	Ca (%)	Sb (%)	Mg (%)	Al (%)	Ni (%)
611501	EK10-136	-	-		Blank	SMI10000496	0.071	0.56	<0.01	<0.01	<0.01	0.06	1.99	0.6	<0.001	0.51	1.11	<0.001
611502	EK10-136	451.4	452.3	0.9	Core	SMI10000497	0.091	0.16	<0.01	0.03	0.03	0.1	3.22	1.13	0.001	0.78	1.27	<0.001
611503	EK10-136	459	459.5	0.5	Core	SMI10000497	0.108	0.19	<0.01	0.09	0.17	0.08	3.55	1.18	<0.001	0.78	1.35	<0.001
611504	EK10-137	86.2	86.7	0.5	Core	SMI10000497	0.106	0.37	<0.01	<0.01	0.04	0.1	3.44	3.22	0.003	0.86	0.61	<0.001
611505	EK10-137	30.4	30.8	0.4	Core	SMI10000497	0.089	0.22	<0.01	0.06	0.42	0.14	4.1	4.81	<0.001	0.85	1.48	<0.001
611506	EK10-137	462	463	1	Core	SMI10000497	0.044	0.43	0.03	0.04	0.08	<0.01	4.42	0.17	<0.001	0.05	0.45	0.002
611507	EK10-137	463	464	1	Core	SMI10000497	0.034	0.48	0.1	0.05	0.1	<0.01	3.37	0.15	0.001	0.05	0.45	<0.001
611508	EK10-137	464	464.5	0.5	Core	SMI10000497	0.002	0.24	0.04	<0.01	<0.01	<0.01	1.29	0.09	<0.001	0.03	0.19	<0.001
611509	EK10-137	464.5	465	0.5	Core	SMI10000497	<0.001	0.3	0.02	0.01	0.08	<0.01	1.88	0.05	0.001	0.03	0.27	<0.001
611510	EK10-137	465	465.5	0.5	Core	SMI10000497	0.001	0.24	0.01	0.02	0.03	0.01	1.39	0.06	<0.001	0.05	0.23	<0.001
611511	EK10-137	465.5	466	0.5	Core	SMI10000497	0.001	0.35	0.05	0.37	0.59	0.02	5.78	0.11	0.003	0.11	0.37	<0.001
611512	EK10-137	466	466.5	0.5	Core	SMI10000497	0.004	0.34	0.03	1.09	1.28	0.02	4.4	0.21	0.003	0.08	0.33	<0.001
611513	EK10-137	466.5	467	0.5	Core	SMI10000497	0.007	0.29	0.02	0.58	0.34	<0.01	3.55	0.11	0.002	0.05	0.26	<0.001
611514	EK10-137	470	470.5	0.5	Core	SMI10000497	0.002	0.21	0.04	0.38	0.73	0.02	4.8	0.19	0.002	0.11	0.18	<0.001
611515	EK10-137	470.5	471	0.5	Core	SMI10000497	0.002	0.26	<0.01	0.05	0.07	<0.01	1	0.13	<0.001	0.07	0.22	<0.001
611516	EK10-137	471	471.5	0.5	Core	SMI10000497	0.003	0.28	0.02	0.12	0.27	0.04	2.55	0.57	0.005	0.2	0.26	<0.001
611517	EK10-137	471.5	472	0.5	Core	SMI10000497	0.002	0.32	0.02	0.25	0.23	0.01	2.54	0.17	0.002	0.08	0.3	<0.001
611518	EK10-137	472	472.5	0.5	Core	SMI10000497	0.002	0.21	<0.01	0.02	0.04	0.01	0.95	0.17	<0.001	0.07	0.17	<0.001
611519	EK10-137	472.5	473	0.5	Core	SMI10000497	0.002	0.29	0.05	0.97	1.89	0.03	6.76	0.33	0.004	0.14	0.28	<0.001
611520	EK10-137	473	474	1	Core	SMI10000497	0.001	0.24	<0.01	<0.01	0.01	0.01	1.09	0.11	<0.001	0.05	0.19	<0.001
611521	EK10-137	474	475	1	Core	SMI10000497	0.06	0.46	0.02	0.09	0.08	<0.01	4.76	0.23	0.002	0.06	0.44	<0.001
611522	EK10-137	467	467.5	0.5	Core	SMI10000497	0.026	0.34	0.05	0.78	0.77	0.01	6.58	0.15	0.003	0.08	0.33	<0.001
611523	EK10-137	467.5	468	0.5	Core	SMI10000497	0.03	0.35	0.02	0.22	0.4	<0.01	3.77	0.11	0.002	0.05	0.34	<0.001
611524	EK10-137	468	468.5	0.5	Core	SMI10000497	0.014	0.24	0.04	0.14	0.26	<0.01	6.88	0.1	0.002	0.05	0.21	<0.001
611525	EK10-137	468	468.5	0.5	Dupl.	SMI10000497	0.009	0.25	0.04	0.26	0.32	<0.01	4.49	0.08	0.002	0.06	0.23	<0.001
611526	EK10-137	468.5	469	0.5	Core	SMI10000497	0.043	0.38	0.03	0.03	0.19	0.01	4.93	0.13	0.001	0.12	0.39	<0.001
611527	EK10-137	469	469.5	0.5	Core	SMI10000497	0.013	0.23	0.02	0.04	0.1	0.01	4.08	0.09	0.001	0.09	0.22	<0.001
611528	EK10-137	469.5	470	0.5	Core	SMI10000497	0.011	0.29	0.01	0.04	0.06	0.01	2	0.2	<0.001	0.08	0.26	<0.001
611529	EK10-137	763.6	764.1	0.5	Core	SMI10000497	<0.001	0.44	<0.01	<0.01	0.04	0.02	1.21	0.15	<0.001	0.33	0.74	<0.001
611530	EK10-138	72.7	73.7	1	Core	SMI10000497	0.099	0.26	0.01	0.02	<0.01	0.2	3.52	5.8	<0.001	0.7	1.13	<0.001
611531	EK10-138	73.7	74.7	1	Core	SMI10000497	0.102	0.27	0.03	0.06	0.05	0.11	4.59	3.41	<0.001	0.44	0.4	<0.001
611532	EK10-138	74.7	75.4	0.7	Core	SMI10000497	0.209	0.27	<0.01	<0.01	<0.01	0.19	4.44	5.17	<0.001	0.64	0.46	<0.001
611533	EK10-138	556.4	556.9	0.5	Core	SMI10000497	0.063	0.36	<0.01	<0.01	<0.01	0.1	4.46	1.57	0.001	0.69	0.48	0.001
611534	EK10-138	556.9	557.8	0.9	Core	SMI10000497	0.074	0.37	<0.01	0.19	0.47	0.06	4.86	1.14	0.002	0.55	0.57	0.001
611535	EK10-138	557.8	558.4	0.6	Core	SMI10000497	0.071	0.33	<0.01	<0.01	0.01	0.08	4.56	1.15	<0.001	0.6	0.79	<0.001
611536	EK10-138	558.4	559.4	1	Core	SMI10000497	0.073	0.37	<0.01	<0.01	<0.01	0.08	4.6	1.37	0.002	0.57	0.5	<0.001
611537	EK10-138	559.4	559.9	0.5	Core	SMI10000497	0.072	0.39	0.01	<0.01	<0.01	0.06	3.7	1.07	0.001	0.53	0.51	0.001
611538	EK10-138	559.9	560.4	0.5	Core	SMI10000497	0.069	0.29	0.02	<0.01	<0.01	0.08	4.5	1.76	0.002	0.46	0.76	<0.001
611539	EK10-138	560.4	560.9	0.5	Core	SMI10000497	0.073	0.34	0.02	<0.01	0.02	0.11	5.37	2.55	<0.001	0.75	0.93	<0.001
611540	EK10-138	560.9	561.5	0.6	Core	SMI10000497	0.071	0.39	<0.01	<0.01	0.01	0.08	4.83	2.14	<0.001	0.71	0.71	0.001

Drillcore Samples - ICP Geochemical analyses (Acme Labs - 1DX)

Sample	Drillhole	From (m)	To (m)	Sample Type	LAB JOB	Au (ppb)	Ag (ppm)	Cu (ppm)	Mo (ppm)	Ti (%)	Ga (ppm)	B (ppm)	W (ppm)	Th (ppm)	V (ppm)	P (%)	Te (ppm)	Al (%)	Mg (%)
611601	EK10-136	0	0	Blank	SM110000558	<0.5	<0.1	3.4	0.2	0.162	5	<1	<0.1	5.2	52	0.125	<0.2	0.74	0.53
611602	EK10-136	407.4	407.9	Core	SM110000558	3169.1	1.8	248.7	21.2	0.004	7	<1	<0.1	0.1	31	0.025	0.5	1.66	1.99
611603	EK10-136	407.9	408.4	Core	SM110000558	493.4	0.8	229.5	8.1	0.005	14	<1	<0.1	0.2	54	0.046	<0.2	3.36	3.93
611604	EK10-136	593.4	594.2	Core	SM110000558	<0.5	0.8	40	14.7	0.001	3	1	<0.1	1	35	0.041	<0.2	1.07	0.2
611605	EK10-137	313.4	314.4	Core	SM110000558	46.4	0.5	20.5	2.9	0.002	8	<1	<0.1	0.5	39	0.043	0.3	2.09	2.22
611606	EK10-137	314.4	315.4	Core	SM110000558	13.1	0.8	19.9	5.2	0.002	5	<1	<0.1	0.7	32	0.047	0.5	1.27	1.02
611607	EK10-137	461	462	Core	SM110000558	9.5	1	29.1	1.8	0.001	<1	<1	<0.1	0.9	7	0.061	<0.2	0.32	0.06
611609	EK10-137	475	476	Core	SM110000558	227.7	9.6	429	4.4	<0.001	<1	2	0.2	3.5	3	0.016	1.1	0.26	0.35
611610	EK10-137	476	477	Core	SM110000584	879.2	6.9	238	3.6	<0.001	2	3	0.1	1.3	8	0.047	1.1	0.5	0.1
611611	EK10-137	460	461	Core	SM110000584	2.6	0.6	23.8	1.8	0.001	1	1	<0.1	0.9	8	0.068	<0.2	0.52	0.08
611612	EK10-138	374.1	375.1	Core	SM110000584	1.8	0.8	81.6	2.2	0.001	1	5	0.4	0.6	25	0.156	<0.2	0.52	1.28
611613	EK10-138	375.1	376.1	Core	SM110000584	0.7	1.2	91.4	4.4	0.001	1	4	0.3	0.9	21	0.149	<0.2	0.51	0.59
611614	EK10-138	376.1	377.1	Core	SM110000584	<0.5	1.9	124.6	1.8	0.001	1	3	0.2	0.7	11	0.096	<0.2	0.46	0.07
611615	EK10-138	377.1	378.1	Core	SM110000584	<0.5	1.6	72.6	11	0.001	<1	2	0.1	0.4	12	0.043	<0.2	0.4	0.12
611616	EK10-138	669.9	670.9	Core	SM110000584	9.1	0.2	15.5	0.6	0.004	7	2	<0.1	2.8	40	0.041	<0.2	3.09	0.49
611617	EK10-138	670.9	671.9	Core	SM110000584	1.7	0.1	36.3	1.1	0.003	7	2	<0.1	2	41	0.115	<0.2	3.03	0.43
611618	EK10-138	671.9	672.9	Core	SM110000581	1.3	0.3	45.2	15.8	0.006	7	2	<0.1	1.9	37	0.35	<0.2	2.81	0.38
611619	EK10-138	672.9	673.9	Core	SM110000581	4	0.1	33.3	0.3	0.005	7	2	<0.1	2.4	35	0.082	0.4	2.95	0.38
611620	EK10-138	673.9	674.9	Core	SM110000581	2.7	<0.1	35.2	0.3	0.009	9	3	<0.1	1.8	53	0.316	<0.2	3.38	0.48
611621	EK10-138	674.9	675.9	Core	SM110000581	4.7	0.2	35.1	7.5	0.006	8	2	<0.1	1.4	45	0.201	<0.2	3.31	0.55
611622	EK10-138	675.9	676.9	Core	SM110000581	2.3	0.1	33.3	0.6	0.005	10	2	<0.1	1.6	71	0.116	<0.2	3.75	0.75
611623	EK10-138	676.9	677.9	Core	SM110000581	5.1	0.1	37.2	0.4	0.006	10	2	<0.1	1.9	71	0.177	<0.2	3.68	0.74
611624	EK10-138	807.1	808.1	Core	SM110000581	<0.5	0.7	65.6	31.2	0.001	3	1	0.3	1.8	51	0.049	<0.2	0.69	0.57
611625	EK10-138	807.1	808.1	Dupl.	SM110000581	<0.5	0.8	69	29.2	0.001	3	1	0.3	2	46	0.048	<0.2	0.63	0.49
611626	EK10-138	808.1	808.6	Core	SM110000582	0.7	1.4	59.4	30.9	0.002	1	<1	0.2	0.8	43	0.039	<0.2	0.26	0.22
611627	EK10-138	816.1	817.1	Core	SM110000582	<0.5	0.5	102.3	35.4	0.046	2	1	0.3	3	62	0.043	<0.2	0.68	0.5
611628	EK10-138	817.1	818	Core	SM110000582	<0.5	1	62.8	30.6	0.054	2	1	0.4	4.4	31	0.029	<0.2	0.52	0.33
611629	EK10-138	799.4	800.4	Core	SM110000582	<0.5	<0.1	33.7	31.3	0.059	9	1	0.3	2.6	108	0.02	<0.2	2.38	3.11
611630	EK10-138	800.4	801.4	Core	SM110000582	<0.5	<0.1	29.6	28.9	0.071	10	<1	0.4	2.8	108	0.02	<0.2	2.67	3.45
611631	EK10-138	801.4	802.4	Core	SM110000582	<0.5	<0.1	30.7	53.7	0.065	9	<1	0.4	2.8	111	0.018	<0.2	2.57	3.13
611632	EK10-138	802.4	803.2	Core	SM110000582	<0.5	<0.1	30	86.1	0.056	7	2	0.1	2.7	104	0.018	<0.2	1.83	2.53
611633	EK10-138	803.4	804.4	Core	SM110000582	<0.5	<0.1	7.6	23.4	<0.001	4	2	0.1	19	8	0.008	<0.2	1.52	1.73
611634	EK10-138	798.9	799.4	Core	SM110000582	<0.5	<0.1	36.3	9.3	0.088	12	1	0.2	3	145	0.023	<0.2	3.54	4.82
611701	EK10-139	0	0	Blank	SM110000662	<0.5	0.7	5.7	0.2	0.15	5	<1	<0.1	5.3	50	0.121	<0.2	0.7	0.51
611702	EK10-139	387.8	388.3	Core	SM110000662	0.5	0.7	49.7	1.9	0.004	5	2	0.2	0.6	51	0.071	<0.2	1.27	0.52
611703	EK10-139	388.3	388.8	Core	SM110000662	0.7	0.9	39.1	6.5	0.006	5	3	0.8	1.7	66	0.544	<0.2	1.42	0.51
611704	EK10-139	388.8	389.8	Core	SM110000662	0.8	1.1	43.8	3.1	0.004	4	2	0.4	0.9	65	0.271	<0.2	1.28	0.5
611705	EK10-139	389.8	390.8	Core	SM110000662	<0.5	1.4	58.7	3.8	0.004	3	2	0.3	0.9	54	0.232	<0.2	0.97	0.33
611706	EK10-139	390.8	391.8	Core	SM110000662	0.7	0.9	54.1	3.5	0.003	3	2	0.2	0.7	35	0.17	<0.2	0.97	0.34
611707	EK10-139	391.8	392.8	Core	SM110000662	<0.5	2.3	55.9	10.9	0.002	2	3	0.3	0.7	63	0.246	0.3	0.55	0.32
611708	EK10-139	392.8	393.8	Core	SM110000662	<0.5	2.2	61.3	9	0.003	3	2	0.2	0.6	76	0.146	0.2	0.9	0.33
611709	EK10-139	393.8	394.8	Core	SM110000666	<0.5	1.2	44.1	10.7	0.004	9	2	0.2	0.5	95	0.061	<0.2	1.92	0.94

Drillcore Samples - ICP Geochemical analyses (Acme Labs - 1DX)

Sample	Drillhole	From (m)	To (m)	Sample Type	LAB JOB	Fe (%)	Hg (ppm)	Pb (ppm)	Cr (ppm)	Na (%)	La (ppm)	Ba (ppm)	U (ppm)	Cd (ppm)	Sr (ppm)	Tl (ppm)	Mn (ppm)	S (%)	Ca (%)
611601	EK10-136	0	0	Blank	SM110000558	2.28	<0.01	0.8	9	0.102	15	212	1.2	<0.1	48	0.1	290	<0.05	0.52
611602	EK10-136	407.4	407.9	Core	SM110000558	4.2	0.88	580.8	6	<0.001	4	20	<0.1	14.4	46	<0.1	1760	2.51	4.99
611603	EK10-136	407.9	408.4	Core	SM110000558	6.41	0.39	84.8	4	<0.001	5	21	<0.1	4.6	72	<0.1	3243	2.85	8.97
611604	EK10-136	593.4	594.2	Core	SM110000558	5.09	0.5	109.7	3	0.024	4	61	0.3	2.1	79	7	555	3.63	3.39
611605	EK10-137	313.4	314.4	Core	SM110000558	4.66	0.04	18.7	10	<0.001	3	71	0.2	<0.1	18	0.1	769	2.54	0.52
611606	EK10-137	314.4	315.4	Core	SM110000558	4.16	0.11	40.2	5	0.008	4	72	0.3	<0.1	18	0.1	517	3.24	0.41
611607	EK10-137	461	462	Core	SM110000558	4.14	0.06	43.6	2	0.003	4	37	0.3	0.9	13	0.3	144	4.52	0.23
611609	EK10-137	475	476	Core	SM110000558	4.35	2.91	3775.5	<1	0.003	4	21	1.3	29.8	105	0.3	352	4.68	0.81
611610	EK10-137	476	477	Core	SM110000584	4.07	1.57	1849.2	4	0.004	5	22	0.3	17.9	23	0.5	96	4.77	0.27
611611	EK10-137	460	461	Core	SM110000584	3	0.02	23.7	3	0.004	5	31	0.4	0.6	15	0.2	138	3.39	0.26
611612	EK10-138	374.1	375.1	Core	SM110000584	4.3	0.69	66.4	6	0.006	4	32	0.3	0.5	84	0.9	1272	2.28	4.33
611613	EK10-138	375.1	376.1	Core	SM110000584	4.82	0.71	39.4	7	0.005	3	23	0.4	1.3	46	0.8	569	4.44	2.08
611614	EK10-138	376.1	377.1	Core	SM110000584	1.72	0.39	21.6	9	0.006	5	94	0.3	<0.1	20	0.2	151	1.84	0.42
611615	EK10-138	377.1	378.1	Core	SM110000584	2.98	0.6	38	7	0.005	3	41	0.2	0.1	14	1.3	163	3.37	0.38
611616	EK10-138	669.9	670.9	Core	SM110000584	6.18	0.03	8.7	2	0.029	14	259	0.6	<0.1	20	<0.1	445	<0.05	0.13
611617	EK10-138	670.9	671.9	Core	SM110000584	5.82	0.02	15.6	5	0.024	12	272	0.4	<0.1	23	<0.1	412	<0.05	0.26
611618	EK10-138	671.9	672.9	Core	SM110000581	5.33	0.04	45.8	4	0.031	15	267	0.8	0.1	46	<0.1	427	<0.05	0.91
611619	EK10-138	672.9	673.9	Core	SM110000581	5.9	0.03	11.9	3	0.034	10	279	0.4	<0.1	41	<0.1	382	0.05	0.24
611620	EK10-138	673.9	674.9	Core	SM110000581	7.13	0.04	8.8	7	0.028	17	266	0.6	<0.1	44	<0.1	509	<0.05	0.88
611621	EK10-138	674.9	675.9	Core	SM110000581	7.18	0.04	42.2	4	0.027	12	284	0.4	<0.1	29	<0.1	524	0.21	0.58
611622	EK10-138	675.9	676.9	Core	SM110000581	8.24	0.06	16.5	16	0.028	14	205	0.3	0.1	30	<0.1	809	0.1	0.74
611623	EK10-138	676.9	677.9	Core	SM110000581	8.6	0.05	12.8	3	0.023	20	154	0.4	<0.1	24	0.1	831	0.09	0.64
611624	EK10-138	807.1	808.1	Core	SM110000581	1.9	0.2	32.3	9	0.011	12	20	1.7	4.2	94	1.9	2556	1.72	10.38
611625	EK10-138	807.1	808.1	Dupl.	SM110000581	2.19	0.22	34.7	9	0.012	13	21	2.3	4.4	70	2.1	1564	2.05	6.65
611626	EK10-138	808.1	808.6	Core	SM110000582	2.1	0.27	33	15	0.018	13	15	0.8	2.9	17	2.1	153	2.22	1.04
611627	EK10-138	816.1	817.1	Core	SM110000582	2.94	0.09	43.2	11	0.013	12	22	8.7	5.2	12	0.3	182	1.96	0.9
611628	EK10-138	817.1	818	Core	SM110000582	2.54	0.14	31.5	8	0.012	14	21	3.3	2.6	10	1.3	171	1.99	0.72
611629	EK10-138	799.4	800.4	Core	SM110000582	10.66	0.08	11.8	97	0.026	2	9	2.5	0.2	10	0.6	854	8.34	1.1
611630	EK10-138	800.4	801.4	Core	SM110000582	9.44	0.09	9	103	0.03	3	11	1.3	0.1	17	0.2	924	7.03	1.96
611631	EK10-138	801.4	802.4	Core	SM110000582	11.57	0.13	11.7	92	0.033	2	10	1.5	0.2	22	0.4	905	9.03	2.27
611632	EK10-138	802.4	803.2	Core	SM110000582	5.57	0.16	15.1	88	0.028	4	19	4.1	0.3	26	0.8	737	4.3	2.99
611633	EK10-138	803.4	804.4	Core	SM110000582	1.46	0.04	26.1	4	0.014	16	15	12.7	0.5	32	0.5	364	0.41	1.82
611634	EK10-138	798.9	799.4	Core	SM110000582	5.98	0.02	7.1	128	0.035	3	11	1.1	<0.1	14	<0.1	973	2.18	0.84
611701	EK10-139	0	0	Blank	SM110000662	2.06	<0.01	7.4	9	0.086	14	243	1.5	<0.1	39	0.1	279	<0.05	0.49
611702	EK10-139	387.8	388.3	Core	SM110000662	3.91	0.2	13.1	14	0.019	8	79	<0.1	0.3	20	0.1	232	1.88	0.29
611703	EK10-139	388.3	388.8	Core	SM110000662	4.08	0.18	10.6	29	0.025	23	87	1.1	1.4	114	0.3	439	2.13	2.26
611704	EK10-139	388.8	389.8	Core	SM110000662	4.35	0.31	11.9	23	0.015	14	66	0.3	1.1	219	0.3	3084	2.53	8.53
611705	EK10-139	389.8	390.8	Core	SM110000662	4.39	0.51	17.1	17	0.022	10	44	0.2	1.1	47	0.4	247	3.36	0.78
611706	EK10-139	390.8	391.8	Core	SM110000662	4.09	0.37	12.4	9	0.016	8	50	0.1	1.4	42	0.3	315	2.91	0.77
611707	EK10-139	391.8	392.8	Core	SM110000662	4.55	0.57	12.5	17	0.016	10	63	1.1	13.1	310	0.9	1861	3.76	8.76
611708	EK10-139	392.8	393.8	Core	SM110000662	4.18	0.63	16.4	23	0.022	9	43	0.3	10.4	29	0.7	308	3.47	0.44
611709	EK10-139	393.8	394.8	Core	SM110000666	5.66	0.28	10.5	14	0.013	5	71	0.1	5.1	40	0.4	511	2.68	0.66

Drillcore Samples - ICP Geochemical analyses (Acme Labs - 1DX)

Sample	Drillhole	From (m)	To (m)	Sample Type	LAB JOB	Au (ppb)	Ag (ppm)	Cu (ppm)	Mo (ppm)	Ti (%)	Ga (ppm)	B (ppm)	W (ppm)	Th (ppm)	V (ppm)	P (%)	Te (ppm)	Al (%)	Mg (%)
611710	EK10-139	394.8	395.8	Core	SM110000666	<0.5	1.2	45.7	11.4	0.003	2	<1	0.1	0.5	54	0.148	<0.2	0.45	0.16
611711	EK10-139	395.8	396.8	Core	SM110000666	0.9	0.8	23.8	8.5	0.002	3	2	0.2	2.2	15	0.118	<0.2	0.8	0.28
611712	EK10-139	396.8	397.8	Core	SM110000666	<0.5	0.8	14.9	17.9	0.001	2	2	0.2	1.9	4	0.005	<0.2	0.6	0.21
611713	EK10-139	397.8	398.8	Core	SM110000666	<0.5	0.8	19.6	10.3	0.002	2	2	0.2	1.9	11	0.057	<0.2	0.52	0.17
611714	EK10-139	398.8	399.8	Core	SM110000666	<0.5	0.3	7.2	21.2	0.002	5	3	0.2	3.9	3	0.004	<0.2	1.27	0.59
611715	EK10-139	399.8	400.3	Core	SM110000666	0.5	0.2	3.8	24.8	0.003	7	3	<0.1	5.1	<2	0.001	<0.2	2.01	1.3
611723	EK10-139	604	604.5	Core	SM110000664	1.1	0.2	6.4	0.4	0.001	7	1	0.4	5.2	<2	0.004	<0.2	1.11	0.75
611724	EK10-140	721.3	722.3	Core	SM110000665	0.6	<0.1	38.7	16.3	0.175	10	<1	0.4	1.7	132	0.033	<0.2	2.93	3.71
611725	EK10-140	721.3	722.3	Dupl.	SM110000665	0.7	<0.1	38.5	5.1	0.191	11	1	0.4	1.5	147	0.036	<0.2	3.09	4.11
611726	EK10-140	722.3	723.3	Core	SM110000665	<0.5	<0.1	53.2	6.7	0.188	10	1	0.6	0.8	164	0.039	<0.2	3.11	4.35
611727	EK10-140	723.3	724.3	Core	SM110000665	<0.5	<0.1	43.8	4.1	0.197	10	2	0.3	0.6	164	0.041	<0.2	3.24	4.34
611728	EK10-140	724.3	725.3	Core	SM110000665	0.8	<0.1	51.3	3.1	0.222	10	3	0.4	0.7	154	0.038	<0.2	3.32	3.64
611729	EK10-140	758.5	759.5	Core	SM110000665	<0.5	<0.1	36.1	7	0.169	12	3	0.9	3.7	130	0.022	<0.2	4.24	5.28
611730	EK10-140	759.5	760.6	Core	SM110000665	<0.5	<0.1	42.3	6.4	0.184	13	2	1.5	3.8	148	0.025	<0.2	4.01	5.65
611731	EK10-140	760.6	761.3	Core	SM110000665	<0.5	<0.1	33.1	4.2	0.138	8	<1	1.9	2.8	123	0.022	<0.2	2.42	3.74

Drillcore Samples - ICP Geochemical analyses (Acme Labs - 1DX)

Sample	Drillhole	From (m)	To (m)	Sample Type	LAB JOB	Fe (%)	Hg (ppm)	Pb (ppm)	Cr (ppm)	Na (%)	La (ppm)	Ba (ppm)	U (ppm)	Cd (ppm)	Sr (ppm)	Tl (ppm)	Mn (ppm)	S (%)	Ca (%)
611710	EK10-139	394.8	395.8	Core	SM110000666	2.12	0.46	7.7	15	0.026	7	78	0.4	24.3	93	1.1	305	1.75	1.63
611711	EK10-139	395.8	396.8	Core	SM110000666	2.35	0.21	39.2	9	0.014	13	100	1.3	3.3	203	0.8	1406	1.44	5.97
611712	EK10-139	396.8	397.8	Core	SM110000666	1.92	0.12	93.9	5	0.008	14	68	1.2	1.5	337	0.7	2813	1.16	9.33
611713	EK10-139	397.8	398.8	Core	SM110000666	1.71	0.19	47.3	11	0.01	10	72	0.8	2.3	181	0.9	1677	1.27	5.1
611714	EK10-139	398.8	399.8	Core	SM110000666	2.42	0.03	17.3	4	0.009	33	86	1.6	0.4	107	1.2	908	0.91	3.42
611715	EK10-139	399.8	400.3	Core	SM110000666	2.96	0.02	14.4	2	0.009	33	98	2.7	0.3	43	0.5	223	0.45	0.25
611723	EK10-139	604	604.5	Core	SM110000664	1.36	0.06	18.3	3	0.024	58	28	1.5	0.4	9	0.1	150	0.12	0.11
611724	EK10-140	721.3	722.3	Core	SM110000665	5.79	0.17	11.2	64	0.055	2	40	1.5	0.5	19	0.2	367	3.55	1.43
611725	EK10-140	721.3	722.3	Dupl.	SM110000665	5.49	0.09	7.2	69	0.071	2	54	1.3	0.3	20	0.1	403	2.87	1.29
611726	EK10-140	722.3	723.3	Core	SM110000665	5.4	0.1	6.1	62	0.073	2	61	0.8	0.2	26	0.1	366	2.85	0.72
611727	EK10-140	723.3	724.3	Core	SM110000665	4.39	0.06	3.4	61	0.108	2	63	0.6	0.3	29	<0.1	355	1.5	0.77
611728	EK10-140	724.3	725.3	Core	SM110000665	5.54	0.08	7.4	54	0.1	2	46	0.9	0.3	25	<0.1	306	3	1.29
611729	EK10-140	758.5	759.5	Core	SM110000665	4.77	0.01	8	117	0.058	4	22	1.9	0.1	34	<0.1	271	0.92	1.61
611730	EK10-140	759.5	760.6	Core	SM110000665	5.31	0.02	10.6	131	0.073	4	23	2.4	<0.1	33	<0.1	252	1.17	0.9
611731	EK10-140	760.6	761.3	Core	SM110000665	3.62	0.02	7.8	113	0.086	4	15	1.8	<0.1	27	<0.1	238	0.91	0.93

Drillcore Samples - ICP Geochemical analyses (Acme Labs - 1DX)

Drillhole	From (m)	To (m)	Sample Type	LAB JOB	K (%)	Ni (ppm)	Se (ppm)	Sb (ppm)	Co (ppm)	Bi (ppm)	Zn (ppm)	Sc (ppm)	As (ppm)
EK10-136	0	0	Blank	SM110000558	0.38	2.9	<0.5	<0.1	5.6	<0.1	39	1.6	<0.5
EK10-136	407.4	407.9	Core	SM110000558	<0.01	1.4	10.1	3.9	5.7	0.4	2534	1.6	123.3
EK10-136	407.9	408.4	Core	SM110000558	<0.01	1.7	4	1.4	9.7	0.2	904	2.7	88.7
EK10-136	593.4	594.2	Core	SM110000558	0.17	21.5	8.4	12.6	90.3	0.1	561	1.9	183.7
EK10-137	313.4	314.4	Core	SM110000558	0.13	5.5	2.9	2.1	12.5	0.5	45	1.7	161.2
EK10-137	314.4	315.4	Core	SM110000558	0.24	5	1.1	6	10	0.4	26	1.5	112.3
EK10-137	461	462	Core	SM110000558	0.26	10.5	1.1	4.5	9.9	0.1	225	1	180.2
EK10-137	475	476	Core	SM110000558	0.24	4	7.5	225.3	4.2	0.3	5558	0.9	340.2
EK10-137	476	477	Core	SM110000584	0.35	9.8	6.2	81.1	7.5	0.4	3397	1.1	292.1
EK10-137	460	461	Core	SM110000584	0.35	12.1	0.8	3	9.5	0.1	121	1.6	114.1
EK10-138	374.1	375.1	Core	SM110000584	0.33	10.4	1.3	51.9	20.3	<0.1	101	3.3	285.9
EK10-138	375.1	376.1	Core	SM110000584	0.35	26.2	4.8	44.6	20.6	<0.1	161	2.9	596.1
EK10-138	376.1	377.1	Core	SM110000584	0.32	63.3	3.4	35.5	14.6	<0.1	12	1.2	115
EK10-138	377.1	378.1	Core	SM110000584	0.28	53.7	2.3	24.1	13.4	<0.1	14	1.3	622.3
EK10-138	669.9	670.9	Core	SM110000584	0.27	3	<0.5	0.6	11.1	0.2	107	2	0.8
EK10-138	670.9	671.9	Core	SM110000584	0.31	4	<0.5	0.9	13.7	0.2	111	3.9	0.9
EK10-138	671.9	672.9	Core	SM110000581	0.29	4.1	<0.5	0.8	13.9	0.5	116	3.5	0.9
EK10-138	672.9	673.9	Core	SM110000581	0.27	3.5	<0.5	0.4	10.4	0.2	96	3.3	<0.5
EK10-138	673.9	674.9	Core	SM110000581	0.3	4.1	<0.5	0.7	14.3	<0.1	131	5.3	0.5
EK10-138	674.9	675.9	Core	SM110000581	0.26	4.1	<0.5	0.9	21.9	0.3	146	4.3	19.3
EK10-138	675.9	676.9	Core	SM110000581	0.24	8	<0.5	0.7	31	0.2	149	7	11.1
EK10-138	676.9	677.9	Core	SM110000581	0.23	3.9	<0.5	0.7	24.5	0.1	152	5	2.8
EK10-138	807.1	808.1	Core	SM110000581	0.18	105.4	5.3	63.6	3.9	0.1	446	2.3	63.7
EK10-138	807.1	808.1	Dupl.	SM110000581	0.17	105.2	5.9	63.1	4.5	0.1	450	2.1	74.8
EK10-138	808.1	808.6	Core	SM110000582	0.11	86.7	3.3	49.1	5.1	0.1	327	1.1	74.6
EK10-138	816.1	817.1	Core	SM110000582	0.21	193.7	5	8	9.8	0.2	671	3.3	70.4
EK10-138	817.1	818	Core	SM110000582	0.2	114.7	3.6	20	3.9	0.2	345	2.6	75.7
EK10-138	799.4	800.4	Core	SM110000582	0.05	45.7	<0.5	2.1	23.7	0.1	44	8.5	36.1
EK10-138	800.4	801.4	Core	SM110000582	0.06	46.4	0.6	2.2	22.1	0.1	43	9.8	29.8
EK10-138	801.4	802.4	Core	SM110000582	0.04	49	0.9	2.3	21.3	0.1	45	9	59.3
EK10-138	802.4	803.2	Core	SM110000582	0.06	95.1	1.5	5.2	20.4	0.1	56	7.7	76
EK10-138	803.4	804.4	Core	SM110000582	0.29	26.2	1	6	1.8	0.5	86	1.3	11.3
EK10-138	798.9	799.4	Core	SM110000582	0.05	60.1	<0.5	0.9	28.1	0.1	52	13.4	8.1
EK10-139	0	0	Blank	SM110000662	0.45	3.4	<0.5	0.2	5.6	<0.1	40	1.3	0.9
EK10-139	387.8	388.3	Core	SM110000662	0.19	19.9	4.9	2.1	8.5	0.1	143	3.3	26.2
EK10-139	388.3	388.8	Core	SM110000662	0.25	18.3	6.3	2.2	9.8	0.2	179	4.6	37.8
EK10-139	388.8	389.8	Core	SM110000662	0.19	19	8.5	2.7	7.5	0.1	137	3.3	47.8
EK10-139	389.8	390.8	Core	SM110000662	0.22	23	9.4	4.5	8.8	0.1	179	2.7	57
EK10-139	390.8	391.8	Core	SM110000662	0.3	19.1	8.9	6.9	8.1	0.1	187	2.4	54.6
EK10-139	391.8	392.8	Core	SM110000662	0.19	30	19.9	16	5.1	<0.1	604	3.3	99.9
EK10-139	392.8	393.8	Core	SM110000662	0.25	32.2	19.4	10.8	8	0.1	639	2.6	96.8
EK10-139	393.8	394.8	Core	SM110000666	0.21	17.3	13.3	3.5	6.8	0.1	342	4.1	42.3

Drillcore Samples - ICP Geochemical analyses (Acme Labs - 1DX)

Drillhole	From (m)	To (m)	Sample Type	LAB JOB	K (%)	Ni (ppm)	Se (ppm)	Sb (ppm)	Co (ppm)	Bi (ppm)	Zn (ppm)	Sc (ppm)	As (ppm)
EK10-139	394.8	395.8	Core	SM110000666	0.15	26.6	17.1	5.3	5.2	<0.1	1294	2.4	48.4
EK10-139	395.8	396.8	Core	SM110000666	0.26	15.3	9.7	3.4	2.7	0.2	364	1.1	37.3
EK10-139	396.8	397.8	Core	SM110000666	0.17	22.8	11.1	3.7	1.2	0.1	391	0.9	30
EK10-139	397.8	398.8	Core	SM110000666	0.18	13.4	7.4	3.2	2.4	0.1	290	0.8	30.9
EK10-139	398.8	399.8	Core	SM110000666	0.32	3.7	1.8	1.4	2.8	0.2	117	0.6	13.1
EK10-139	399.8	400.3	Core	SM110000666	0.37	1.5	0.8	0.8	0.5	0.3	119	0.4	3.3
EK10-139	604	604.5	Core	SM110000664	0.33	0.1	<0.5	0.7	0.2	0.3	147	0.3	1.1
EK10-140	721.3	722.3	Core	SM110000665	0.04	117.9	1.6	2.6	40	<0.1	119	10	20.4
EK10-140	721.3	722.3	Dupl.	SM110000665	0.07	77	1.1	1.7	37	<0.1	89	11.3	13
EK10-140	722.3	723.3	Core	SM110000665	0.06	67.4	0.7	1.6	38.9	<0.1	80	11.4	12.9
EK10-140	723.3	724.3	Core	SM110000665	0.07	59.8	<0.5	0.5	36.5	<0.1	110	10.3	6
EK10-140	724.3	725.3	Core	SM110000665	0.06	58.8	0.7	0.8	41.6	<0.1	96	11.4	11.5
EK10-140	758.5	759.5	Core	SM110000665	0.04	55.3	<0.5	0.4	26.5	0.1	50	13.6	3.3
EK10-140	759.5	760.6	Core	SM110000665	0.06	62.3	<0.5	0.4	30.9	0.1	54	15.4	3.5
EK10-140	760.6	761.3	Core	SM110000665	0.05	52.4	<0.5	0.6	29.1	0.1	40	10.6	6.6

Drillhole Samples - Lithogeochemical analyses (Acme Labs - 4A+4B)

Sample	Drillhole	From (m)	To (m)	SiO2 (wt.%)	Al2O3 (wt.%)	Fe2O3 (wt.%)	MgO (wt.%)	CaO (wt.%)	Na2O (wt.%)	K2O (wt.%)	TiO2 (wt.%)	P2O5 (wt.%)	MnO (wt.%)	Cr2O3 (wt.%)	Ni (ppm)	Sc (ppm)	LOI (wt.%)	Tot_C (wt.%)	Tot_S (wt.%)	SUM (wt.%)
611552	EK10-136	86.2	86.4	47.74	15.11	8.32	6.02	6.97	0.49	4.39	0.66	0.15	0.23	0.006	42	29	9.5	1.44	0.2	99.87
611553	EK10-136	115.8	116	46.99	13.62	10.53	6.82	6.1	1.25	4.27	0.52	0.47	0.17	0.013	24	34	8.8	1.26	0.03	99.85
611554	EK10-136	495.5	495.7	60.78	17.95	4.9	2.1	1.11	3.26	4.46	0.7	0.62	0.09	<0.002	<20	18	3.8	0.05	1.42	99.96
611555	EK10-136	592.9	593	54.79	23.84	7.21	1.03	0.61	1.06	4.83	1.01	0.11	0.04	<0.002	<20	17	5.1	0.12	0.03	99.95
611556	EK10-136	608.7	608.8	58.06	20.08	7.61	1.78	0.82	2.89	3.56	0.6	0.17	0.04	<0.002	<20	16	4.1	0.04	<0.02	99.96
611557	EK10-136	647.7	647.8	47.02	15.06	12.73	7.58	5.28	3.36	0.61	1.46	0.23	0.3	0.003	<20	44	6.1	0.6	0.83	99.82
611558	EK10-136	720.7	720.9	47.9	14.98	9.14	9.17	10.65	2.12	0.59	0.95	0.07	0.22	0.028	51	44	4	0.23	0.04	99.81
611560	EK10-137	480.2	480.4	57.72	17.95	7.36	1.95	0.31	0.05	6.07	0.89	0.11	0.08	0.012	69	44	7.3	0.45	3.97	99.92
611561	EK10-137	573.4	573.6	60.65	16.04	6.14	1.96	1.41	0.1	4.88	0.9	0.03	0.06	0.008	30	20	7.6	1.47	0.3	99.96
611562	EK10-137	583.4	583.6	58.93	18.26	4.7	0.83	4.28	2.21	3.84	0.6	0.21	0.1	<0.002	<20	6	5.8	0.93	<0.02	99.99
611563	EK10-137	598.5	598.7	47.02	14.99	10.73	9.87	6.44	2.92	0.32	0.86	0.08	0.2	0.018	52	43	6.3	0.39	0.03	99.81
611564	EK10-137	631	631.2	77.66	10.86	1.54	0.27	0.42	1.61	6.82	0.08	0.01	0.01	0.007	<20	1	0.6	0.07	0.14	100.01
611565	EK10-137	673	673.2	64.35	15.58	4.34	2.31	3.07	3.35	2.58	0.1	0.01	0.05	<0.002	<20	2	4.1	0.6	0.08	99.95
611566	EK10-137	692.4	692.6	79.59	11.63	1.08	0.42	0.23	6.17	0.16	0.07	0.01	<0.01	0.006	<20	1	0.6	0.03	0.05	100.02
611567	EK10-137	745.3	745.5	77.61	11.04	1.42	0.19	0.22	0.15	8.4	0.08	0.01	0.02	0.005	<20	1	0.8	0.04	0.23	100.01
611568	EK10-137	770.8	771	71.32	14.64	1.9	1.67	0.6	0.19	6.59	0.1	<0.01	0.03	<0.002	<20	2	2.8	0.13	0.15	99.96
611569	EK10-138	589.7	589.8	48.51	18.18	7.87	8.93	2.7	5.6	0.2	0.73	0.13	0.09	0.011	93	35	6.8	0.54	0.11	99.81
611570	EK10-138	565.4	565.6	61.56	18.22	7.09	1.63	0.62	2.26	3.73	0.68	0.07	0.05	0.002	23	13	3.8	0.12	<0.02	99.96
611571	EK10-138	668	668.2	63.81	17.83	4.54	1.02	2.14	2.62	3.19	0.45	0.12	0.11	<0.002	<20	6	3.9	0.4	0.03	99.98
611572	EK10-138	689.2	689.4	50.15	17.1	9.56	7.97	3.24	4.35	0.52	0.7	0.14	0.18	0.009	55	34	5.9	0.27	0.58	99.85
611573	EK10-138	717.8	718	47.33	14.26	10.28	10.3	9.54	1.42	0.08	0.84	0.06	0.18	0.029	46	42	5.5	0.09	0.1	99.8
611574	EK10-138	748.9	749.1	49.87	14.88	11.67	5.42	9.08	2.3	1.62	1.5	0.22	0.22	0.006	<20	46	3	0.25	0.63	99.86
611576	EK10-138	804.4	804.6	67.39	15.78	1.73	3.88	0.45	0.25	5.57	0.15	0.03	0.03	<0.002	<20	6	4.6	0.11	0.19	99.95
611577	EK10-138	841.4	841.6	75.66	12.14	2.11	2.36	0.2	2.07	2.84	0.11	<0.01	0.01	<0.002	<20	2	2.4	<0.02	<0.02	99.98
611578	EK10-138	842.4	842.6	84.66	8.24	0.57	0.21	0.43	2.63	2.54	0.06	0.01	<0.01	0.004	<20	<1	0.6	0.04	0.04	100.03
611752	EK10-139	404.7	404.9	78.99	9.08	1.13	0.21	0.57	2.47	3.47	0.08	0.02	0.01	0.002	<20	<1	3.8	0.13	0.18	100
611753	EK10-139	407.5	407.6	43.3	13.67	9.98	8.37	8.98	1.46	2.53	0.84	0.07	0.1	0.025	46	39	10.4	1.71	1.66	99.82
611754	EK10-139	477.2	477.3	73.64	13.8	1.89	0.83	0.75	0.53	5.7	0.09	<0.01	0.03	<0.002	<20	2	2.6	0.15	0.08	99.98
611780	EK10-140	609.5	609.7	40.99	21.14	12.9	10.52	0.63	1.55	4.13	0.73	0.15	0.07	0.019	53	30	6.7	0.09	0.02	99.76
611781	EK10-140	699.1	699.2	55.58	16.92	7.58	6.07	1.46	5.47	0.81	1.06	0.29	0.09	0.002	44	24	4.4	0.19	0.53	99.86
611782	EK10-140	708	708.1	52.7	18.65	6.35	4.74	2.31	3.67	4.84	1.17	0.3	0.12	0.004	40	25	4.9	0.39	0.42	99.88
611783	EK10-140	721.1	721.2	48.71	17.04	9.57	7.08	3.33	3.26	3.69	0.73	0.11	0.06	0.011	49	36	6.1	0.06	3.33	99.85
611784	EK10-140	781.9	782	48.37	14.37	12.53	7.98	7.45	3.32	0.13	1.37	0.16	0.13	0.004	<20	45	3.9	0.03	0.09	99.79
611785	EK10-140	803.7	803.8	69.94	15.77	1.69	3.22	0.37	0.32	4.78	0.12	<0.01	<0.01	<0.002	<20	2	3.6	0.06	1.06	99.93

Drillhole Samples - Lithogeochemical analyses (Acme Labs - 4A+4B)

Sample	Drillhole	From (m)	To (m)	Ba (ppm)	Be (ppm)	Co (ppm)	Cs (ppm)	Ga (ppm)	Hf (ppm)	Nb (ppm)	Rb (ppm)	Sn (ppm)	Sr (ppm)	Ta (ppm)	Th (ppm)	U (ppm)	V (ppm)	W (ppm)	Zr (ppm)	Y (ppm)	La (ppm)	Ce (ppm)
611552	EK10-136	86.2	86.4	2563	<1	32.4	4.4	14.1	1.8	4.1	103.1	<1	130.9	0.3	1.5	1	207	1.3	71.2	18.9	6.9	13.3
611553	EK10-136	115.8	116	1792	<1	34.4	2.6	12.7	1	2.4	72.9	<1	394	0.1	1.7	1	285	<0.5	30.7	13.3	8.5	16.9
611554	EK10-136	495.5	495.7	1602	1	22.1	1.8	23.3	2.6	5.4	130	1	73.2	0.3	2.5	1.4	260	3	96.5	15.7	5.1	12.8
611555	EK10-136	592.9	593	2467	1	10.8	4.1	26.7	5.2	12.7	119.6	2	180.4	0.9	8.6	4.4	204	1.6	206.5	32.8	28.3	55.9
611556	EK10-136	608.7	608.8	1823	2	15.3	7.5	21.3	4.7	9.3	109.2	2	250.2	0.5	7.9	2.6	141	1.2	165.3	31.7	20	41.9
611557	EK10-136	647.7	647.8	301	<1	34	1.8	18.5	2.6	3.7	10	<1	194.3	0.3	1.3	0.7	382	1.3	86.9	29.7	6.3	14.9
611558	EK10-136	720.7	720.9	73	<1	44.9	1.1	14.2	1.5	0.8	15.5	<1	168.1	<0.1	<0.2	<0.1	292	0.5	48.9	18	1.6	4.7
611560	EK10-137	480.2	480.4	850	1	49.1	6.9	26.5	1.8	2.5	164.7	7	20.3	0.2	1	0.8	281	7.1	64.7	20.6	2.6	6.4
611561	EK10-137	573.4	573.6	1590	2	13.2	5.2	19.1	4.8	7.2	134.3	2	44.8	0.5	4.4	1.3	77	3.2	154.7	38.7	17.4	34.9
611562	EK10-137	583.4	583.6	1835	2	4.9	4.7	19.6	3.6	8.7	97.8	<1	196.4	0.5	6.2	3.2	93	0.8	146	25	22.5	45
611563	EK10-137	598.5	598.7	331	<1	41.8	1.8	15.1	1.3	1	8.8	<1	271.2	<0.1	<0.2	0.1	280	<0.5	41.4	19.8	1.7	4.9
611564	EK10-137	631	631.2	674	5	0.4	1.1	9.4	6.5	29.8	110.2	2	85.9	1.8	12.8	7.8	<8	2.2	169.1	60.5	39.6	82.7
611565	EK10-137	673	673.2	465	6	<0.2	3.1	21.6	9.2	39.5	95.5	4	237.5	2.7	17.5	9.9	<8	2.6	228	54.4	36.7	91.6
611566	EK10-137	692.4	692.6	94	3	0.2	0.2	12.3	6.1	36.4	5.3	2	132.1	2	12.7	7.5	<8	1.5	159.8	53.1	32.8	68.5
611567	EK10-137	745.3	745.5	217	5	0.2	2	19.7	6.6	29.1	157.9	4	73.1	1.8	12.7	7.5	<8	2.6	181.3	63.5	45	91
611568	EK10-137	770.8	771	285	6	<0.2	5.8	31	9.1	39.4	211.1	6	34.6	2.6	17	10.2	<8	0.8	223.9	84	61.5	125
611569	EK10-138	589.7	589.8	197	1	46.4	1.4	16.8	1.3	2.5	3.7	<1	196.9	0.2	0.8	1.3	240	0.9	52.3	18.9	3.8	8.6
611570	EK10-138	565.4	565.6	1828	2	11.4	6.6	20	4.7	10	110	2	104.5	0.7	8.4	2.4	133	1.7	167	31.6	27.9	57.5
611571	EK10-138	668	668.2	1573	<1	7.6	1.9	19	4.5	8.5	80.7	3	152.7	0.6	7.8	3.6	68	0.8	186.4	28.2	25.1	49.3
611572	EK10-138	689.2	689.4	360	<1	35.2	0.9	15.1	1.5	2.2	9.5	<1	293.2	0.2	0.8	1.7	228	0.9	48.1	20.9	3.2	8
611573	EK10-138	717.8	718	15	<1	42.9	1	15.6	1.2	1.1	1.6	<1	37	<0.1	0.2	<0.1	280	<0.5	37.9	21.1	1.6	5
611574	EK10-138	748.9	749.1	540	1	33.6	2	18.7	2.4	3.7	32.3	<1	220.9	0.2	1.1	0.6	415	0.5	86.5	34.8	6.9	16.5
611576	EK10-138	804.4	804.6	172	7	1	8.1	22.7	6.9	50	255.6	4	25.3	4.1	36.5	20.6	14	0.6	130.5	29.5	29.5	56.2
611577	EK10-138	841.4	841.6	164	3	0.4	2.6	21.5	6.8	28.1	108.5	4	62.3	1.9	12.4	6.5	<8	1.1	182.4	62.5	46.3	89.3
611578	EK10-138	842.4	842.6	275	3	0.4	0.6	10.9	3.8	16.2	46	2	111.6	1.3	8.7	6.1	<8	5	110.3	39	45.7	88.4
611752	EK10-139	404.7	404.9	1467	2	0.3	0.9	12.3	5.7	22	67.8	2	114.8	1.4	9.5	6.3	<8	3.5	169.2	51.4	40.5	78.7
611753	EK10-139	407.5	407.6	507	<1	39.1	4.8	13.7	1.1	0.9	62.9	<1	370.8	0.2	<0.2	0.2	249	29.1	38	21	1.2	4.6
611754	EK10-139	477.2	477.3	223	9	0.3	6.9	27.8	8.9	38.5	338.7	5	39	2.6	16.8	10	<8	2.6	219	95	48.6	103.4
611780	EK10-140	609.5	609.7	1914	2	26.9	3.8	21.3	2.5	7.8	89.3	<1	96.9	0.3	3	1.9	197	1.4	103.9	27.6	5.6	13.6
611781	EK10-140	699.1	699.2	440	2	34.1	1.9	18.8	3	9.1	17.8	<1	245.1	0.2	2.6	2.2	240	1.3	134.2	25.2	12.6	28.8
611782	EK10-140	708	708.1	1242	2	31.7	2.4	19.4	3.3	10.1	108.3	<1	132	0.4	3.9	2	245	0.9	145.2	29.2	12.8	31.2
611783	EK10-140	721.1	721.2	959	<1	39.1	1.2	15.8	1.7	3.1	68.1	<1	324.1	<0.1	0.8	1.5	260	0.8	57.1	20.1	4.3	10.4
611784	EK10-140	781.9	782	113	<1	35	2.1	16.8	2.3	3.6	3	<1	314.7	<0.1	1.3	0.6	405	1.1	95.4	32.4	6.3	16
611785	EK10-140	803.7	803.8	415	5	0.6	8	27	8.6	37.4	160.8	4	23.4	2.3	18	4.7	<8	1	247.6	50	38.6	87.9

Drillhole Samples - Lithogeochemical analyses (Acme Labs - 4A+4B)

Sample	Drillhole	From (m)	To (m)	Pr (ppm)	Nd (ppm)	Sm (ppm)	Eu (ppm)	Gd (ppm)	Tb (ppm)	Dy (ppm)	Ho (ppm)	Er (ppm)	Tm (ppm)	Yb (ppm)	Lu (ppm)	Ag (ppm)	As (ppm)	Au (ppm)	Bi (ppm)	Cd (ppm)	Cu (ppm)	Hg (ppm)
611552	EK10-136	86.2	86.4	1.71	7.8	1.94	0.75	2.61	0.49	3.17	0.69	2.12	0.27	2.09	0.33	0.1	11.1	1.6	<0.1	<0.1	47.3	0.01
611553	EK10-136	115.8	116	2.1	9.8	2.16	0.63	2.37	0.39	2.33	0.47	1.29	0.18	1.37	0.2	<0.1	<0.5	<0.5	<0.1	<0.1	102	0.02
611554	EK10-136	495.5	495.7	1.74	8.6	2.35	0.64	2.63	0.45	2.55	0.57	1.77	0.26	1.93	0.33	0.3	50.7	26.1	<0.1	<0.1	11.9	0.01
611555	EK10-136	592.9	593	6.53	28.1	5.38	1.3	5.19	0.87	5.28	1.17	3.5	0.51	3.82	0.58	<0.1	1.5	<0.5	0.3	0.1	28.5	0.02
611556	EK10-136	608.7	608.8	4.77	21.6	4.56	1.33	4.9	0.83	4.96	1.07	3.24	0.47	3.58	0.52	0.2	<0.5	27.5	<0.1	<0.1	16.8	0.01
611557	EK10-136	647.7	647.8	2.14	11.8	3.38	1.18	4.3	0.82	5.25	1.17	3.27	0.47	3.79	0.54	<0.1	<0.5	<0.5	<0.1	0.2	6.4	<0.01
611558	EK10-136	720.7	720.9	0.79	4.9	1.71	0.69	2.54	0.49	3.22	0.69	2.1	0.28	2.19	0.3	<0.1	10.2	0.7	<0.1	0.3	49.9	<0.01
611560	EK10-137	480.2	480.4	0.9	4.9	1.56	0.43	2.28	0.49	3.34	0.77	2.36	0.34	2.67	0.38	1.2	216.2	177.9	0.5	0.2	92.4	0.08
611561	EK10-137	573.4	573.6	4.54	20.6	4.46	0.96	5.12	0.97	6.25	1.41	4.34	0.65	4.31	0.68	<0.1	18.6	<0.5	0.1	0.1	12.3	0.05
611562	EK10-137	583.4	583.6	5.28	23.4	4.61	1.35	4.61	0.72	4.29	0.86	2.59	0.4	2.69	0.42	0.2	1.1	6.6	0.1	0.2	6.7	0.03
611563	EK10-137	598.5	598.7	0.82	4.5	1.66	0.7	2.68	0.53	3.56	0.82	2.44	0.37	2.38	0.36	<0.1	4.3	<0.5	<0.1	0.1	38.8	0.01
611564	EK10-137	631	631.2	9.74	41.7	8.62	0.11	9.41	1.68	10.67	2.25	7.02	1.14	7.53	1.16	0.2	11.1	<0.5	0.2	0.3	4.4	0.02
611565	EK10-137	673	673.2	11.57	49.8	10.89	0.1	11.44	1.94	11.39	2.11	6.12	0.97	7.26	1.11	0.2	1.6	<0.5	0.3	0.2	4.7	0.02
611566	EK10-137	692.4	692.6	8.04	34.1	7.56	0.07	8.02	1.43	9.17	2	6.42	1.05	7.48	1.12	0.2	1.5	<0.5	0.2	0.3	4.4	0.03
611567	EK10-137	745.3	745.5	10.37	42.7	8.24	0.1	8.86	1.58	9.93	2.25	7.03	1.13	7.62	1.16	<0.1	14.1	<0.5	0.2	<0.1	4	<0.01
611568	EK10-137	770.8	771	14.33	59.4	12.18	0.12	12.95	2.33	14.57	3.07	9.4	1.5	9.92	1.52	0.3	2.7	<0.5	0.3	0.3	3.9	0.03
611569	EK10-138	589.7	589.8	1.15	5.7	1.71	0.34	2.43	0.47	3.07	0.71	2.25	0.33	2.11	0.31	<0.1	38.6	0.9	<0.1	0.1	55	0.03
611570	EK10-138	565.4	565.6	6.71	28.5	4.93	1.21	5.11	0.92	5.65	1.21	3.51	0.53	3.55	0.52	0.3	2.1	3	<0.1	<0.1	6.8	0.03
611571	EK10-138	668	668.2	5.8	24.6	4.53	1.05	4.33	0.77	4.7	0.95	2.99	0.44	3.16	0.49	0.2	6.3	8.3	0.3	0.3	17.6	0.03
611572	EK10-138	689.2	689.4	1.17	6.8	1.95	0.61	2.64	0.52	3.46	0.75	2.29	0.35	2.31	0.35	<0.1	0.6	1.6	<0.1	0.2	32.7	0.01
611573	EK10-138	717.8	718	0.88	5.4	1.77	0.75	2.66	0.53	3.63	0.78	2.32	0.36	2.29	0.34	<0.1	2.3	0.9	<0.1	0.1	53.7	<0.01
611574	EK10-138	748.9	749.1	2.4	12.1	3.54	1.29	4.69	0.91	5.72	1.24	3.7	0.56	3.77	0.58	<0.1	1.5	1.1	<0.1	0.2	13.2	<0.01
611576	EK10-138	804.4	804.6	6.13	22	4.03	0.21	3.92	0.72	4.59	0.93	2.86	0.5	4.25	0.71	<0.1	4.1	1.1	0.6	0.2	4.2	0.03
611577	EK10-138	841.4	841.6	10.62	42.9	8.58	0.12	8.6	1.59	9.9	2.14	6.63	1.02	6.97	1.06	0.1	0.7	1.7	0.2	0.2	3.1	0.01
611578	EK10-138	842.4	842.6	10.2	41.9	7.48	0.11	7.2	1.25	7.46	1.46	4.26	0.63	4.09	0.6	<0.1	3.3	0.9	<0.1	<0.1	5.3	<0.01
611752	EK10-139	404.7	404.9	8.9	38	7.2	0.15	7.69	1.3	7.85	1.72	5.6	0.87	6.63	0.94	0.2	0.7	2.9	0.2	0.2	3	0.01
611753	EK10-139	407.5	407.6	0.75	5	1.83	0.62	2.83	0.51	3.28	0.71	2.17	0.33	2.25	0.33	<0.1	8.2	2.4	<0.1	0.2	46.7	0.02
611754	EK10-139	477.2	477.3	12.52	56	12.23	0.12	13.63	2.37	14.67	3.07	9.71	1.55	11.08	1.55	0.2	<0.5	1	0.3	0.2	3.1	0.02
611780	EK10-140	609.5	609.7	1.61	7.2	2.45	0.61	3.68	0.74	4.61	1.02	2.92	0.46	2.99	0.45	<0.1	0.8	<0.5	<0.1	0.1	45.6	<0.01
611781	EK10-140	699.1	699.2	3.3	14	3.39	0.55	4.15	0.71	4.38	0.88	2.67	0.41	2.52	0.4	<0.1	4.5	<0.5	<0.1	0.2	29.1	0.02
611782	EK10-140	708	708.1	3.6	14.4	3.56	1.72	4.23	0.79	4.5	1.04	3.01	0.46	3.16	0.46	<0.1	3.5	<0.5	<0.1	0.1	35.9	0.04
611783	EK10-140	721.1	721.2	1.28	7.1	1.6	0.72	2.6	0.5	3.26	0.71	2.01	0.31	1.95	0.28	<0.1	8.8	<0.5	<0.1	<0.1	42.6	0.08
611784	EK10-140	781.9	782	2.05	8.8	2.87	1.1	4.31	0.84	5.42	1.19	3.34	0.57	3.75	0.55	<0.1	1.9	<0.5	<0.1	0.1	18.2	<0.01
611785	EK10-140	803.7	803.8	10.31	40.3	7.88	0.11	8.42	1.64	9.78	1.87	5.12	0.82	5.66	0.96	0.2	0.8	<0.5	0.3	0.3	4.7	0.03

Drillhole Samples - Lithogeochemical analyses (Acme Labs - 4A+4B)

Sample	Drillhole	From (m)	To (m)	Mo (ppm)	Ni (ppm)	Pb (ppm)	Sb (ppm)	Se (ppm)	Tl (ppm)	Zn (ppm)
611552	EK10-136	86.2	86.4	0.9	34.1	3.6	0.6	<0.5	<0.1	74
611553	EK10-136	115.8	116	0.5	19.2	2.2	0.2	<0.5	<0.1	73
611554	EK10-136	495.5	495.7	0.8	1.7	12.7	1.9	<0.5	0.1	22
611555	EK10-136	592.9	593	0.3	2.6	6.3	<0.1	<0.5	<0.1	93
611556	EK10-136	608.7	608.8	0.8	7.3	21.9	0.2	<0.5	<0.1	145
611557	EK10-136	647.7	647.8	9.4	5.3	1.5	<0.1	0.7	<0.1	94
611558	EK10-136	720.7	720.9	0.1	34.8	0.4	<0.1	<0.5	<0.1	58
611560	EK10-137	480.2	480.4	1.1	64.6	79	9	1.4	0.4	47
611561	EK10-137	573.4	573.6	0.2	7.5	9.8	2.6	<0.5	<0.1	59
611562	EK10-137	583.4	583.6	0.4	1.1	4.6	0.1	<0.5	<0.1	87
611563	EK10-137	598.5	598.7	0.4	24.6	1.1	0.3	<0.5	<0.1	68
611564	EK10-137	631	631.2	2.5	0.6	28.2	0.8	0.7	<0.1	106
611565	EK10-137	673	673.2	3	<0.1	19.5	0.4	<0.5	0.1	121
611566	EK10-137	692.4	692.6	0.2	0.5	18.6	0.2	<0.5	<0.1	106
611567	EK10-137	745.3	745.5	2.6	0.4	11	0.5	<0.5	<0.1	19
611568	EK10-137	770.8	771	4.5	0.2	32.9	0.6	0.6	0.2	150
611569	EK10-138	589.7	589.8	0.2	51.3	3.7	1.1	<0.5	<0.1	72
611570	EK10-138	565.4	565.6	<0.1	4.3	2.2	0.3	<0.5	0.1	117
611571	EK10-138	668	668.2	0.4	1.6	42.6	0.2	<0.5	<0.1	74
611572	EK10-138	689.2	689.4	0.4	58.2	1.6	<0.1	<0.5	<0.1	105
611573	EK10-138	717.8	718	0.5	41.8	0.6	<0.1	<0.5	<0.1	64
611574	EK10-138	748.9	749.1	1.3	5	2	0.2	<0.5	<0.1	93
611576	EK10-138	804.4	804.6	6.8	6.5	33.8	1.4	0.6	0.2	59
611577	EK10-138	841.4	841.6	0.8	0.3	14.3	0.3	<0.5	<0.1	97
611578	EK10-138	842.4	842.6	0.7	0.8	8.4	0.3	<0.5	<0.1	27
611752	EK10-139	404.7	404.9	1.3	0.3	9.8	0.4	0.6	<0.1	37
611753	EK10-139	407.5	407.6	0.2	41.7	1.3	0.5	<0.5	0.1	66
611754	EK10-139	477.2	477.3	<0.1	0.3	18.7	0.3	0.5	0.2	114
611780	EK10-140	609.5	609.7	0.2	41.5	0.6	<0.1	<0.5	<0.1	118
611781	EK10-140	699.1	699.2	4.3	38.1	3.6	5.9	<0.5	<0.1	99
611782	EK10-140	708	708.1	0.6	33.4	4.5	6.2	<0.5	<0.1	78
611783	EK10-140	721.1	721.2	2.2	54	5.6	0.8	0.8	<0.1	44
611784	EK10-140	781.9	782	0.8	9.5	1.6	0.5	<0.5	<0.1	91
611785	EK10-140	803.7	803.8	0.3	0.5	14.8	0.2	<0.5	<0.1	140

APPENDIX C

List of Mineral Tenures for the Eskay Property

Tenure Number	Claim Name	Owner	Tenure Type	Tenure Sub Type	Map Number	Issue Date	Good To Date	Status	Area (ha)
251344	COUL 1	202689 (100%)	Mineral	Claim	104B058	1986/feb/28	2014/jan/31	GOOD	500.0
251345	COUL 2	202689 (100%)	Mineral	Claim	104B058	1986/feb/28	2014/jan/31	GOOD	500.0
251346	COUL 3	202689 (100%)	Mineral	Claim	104B058	1986/feb/28	2014/jan/31	GOOD	500.0
251347	COUL 4	202689 (100%)	Mineral	Claim	104B058	1986/feb/28	2014/jan/31	GOOD	500.0
251358	UNUK 1	202689 (100%)	Mineral	Claim	104B059	1986/feb/28	2014/jan/31	GOOD	500.0
251360	UNUK 11	202689 (100%)	Mineral	Claim	104B059	1986/feb/28	2014/jan/31	GOOD	500.0
251361	UNUK 12	202689 (100%)	Mineral	Claim	104B059	1986/feb/28	2014/jan/31	GOOD	500.0
251374	UNUK 13	202689 (100%)	Mineral	Claim	104B068	1986/feb/28	2014/jan/31	GOOD	400.0
251375	UNUK 14	202689 (100%)	Mineral	Claim	104B059	1986/feb/28	2014/jan/31	GOOD	400.0
251379	UNUK 22	202689 (100%)	Mineral	Claim	104B059	1986/feb/28	2014/jan/31	GOOD	500.0
251844	LANCE 3	202689 (100%)	Mineral	Claim	104B069	1987/apr/28	2014/jan/31	GOOD	450.0
251845	LANCE 4	202689 (100%)	Mineral	Claim	104B069	1987/apr/28	2014/jan/31	GOOD	450.0
252352	SKOOKUM	202689 (100%)	Mineral	Claim	104B068	1989/jan/13	2014/jan/31	GOOD	400.0
252872	SIB 27	202689 (100%)	Mineral	Claim	104B068	1989/jun/29	2016/jan/31	GOOD	25.0
252876	SIB 31	202689 (100%)	Mineral	Claim	104B068	1989/jun/29	2016/jan/31	GOOD	25.0
253015	POLO 7	202689 (100%)	Mineral	Claim	104B058	1989/sep/04	2014/jan/31	GOOD	500.0
253016	POLO 8	202689 (100%)	Mineral	Claim	104B058	1989/sep/04	2014/jan/31	GOOD	500.0
253146	AFTOM #7	202689 (100%)	Mineral	Claim	104B069	1989/sep/16	2014/jan/31	GOOD	400.0
253147	AFTOM #9	202689 (100%)	Mineral	Claim	104B068	1989/sep/15	2014/jan/31	GOOD	500.0
253152	AFTOM #14	202689 (100%)	Mineral	Claim	104B069	1989/sep/13	2014/jan/31	GOOD	500.0
253153	AFTOM #15	202689 (100%)	Mineral	Claim	104B069	1989/sep/13	2014/jan/31	GOOD	500.0
253154	AFTOM #16	202689 (100%)	Mineral	Claim	104B069	1989/sep/18	2014/jan/31	GOOD	400.0
253155	AFTOM #18	202689 (100%)	Mineral	Claim	104B068	1989/sep/17	2014/jan/31	GOOD	400.0
253156	AFTOM #19	202689 (100%)	Mineral	Claim	104B068	1989/sep/16	2014/jan/31	GOOD	500.0
253157	AFTOM #20	202689 (100%)	Mineral	Claim	104B068	1989/sep/17	2014/jan/31	GOOD	500.0
253176	P-MAC #1	202689 (100%)	Mineral	Claim	104B058	1989/sep/14	2014/jan/31	GOOD	25.0
253177	P-MAC #2	202689 (100%)	Mineral	Claim	104B058	1989/sep/14	2014/jan/31	GOOD	25.0
253178	P-MAC #3	202689 (100%)	Mineral	Claim	104B058	1989/sep/14	2014/jan/31	GOOD	25.0
253179	P-MAC #4	202689 (100%)	Mineral	Claim	104B058	1989/sep/14	2014/jan/31	GOOD	25.0
253180	P-MAC #5	202689 (100%)	Mineral	Claim	104B058	1989/sep/14	2014/jan/31	GOOD	25.0
253181	P-MAC #6	202689 (100%)	Mineral	Claim	104B058	1989/sep/14	2014/jan/31	GOOD	25.0
253182	P-MAC #7	202689 (100%)	Mineral	Claim	104B058	1989/sep/14	2014/jan/31	GOOD	25.0
253183	P-MAC #8	202689 (100%)	Mineral	Claim	104B058	1989/sep/14	2014/jan/31	GOOD	25.0
253184	P-MAC #9	202689 (100%)	Mineral	Claim	104B058	1989/sep/14	2014/jan/31	GOOD	25.0
253185	P-MAC #10	202689 (100%)	Mineral	Claim	104B058	1989/sep/14	2014/jan/31	GOOD	25.0
253240	POLO 13	202689 (100%)	Mineral	Claim	104B058	1989/sep/15	2014/jan/31	GOOD	125.0
253295	FRED 15	202689 (100%)	Mineral	Claim	104B068	1989/oct/11	2014/jan/31	GOOD	375.0
255254	S.I.B. #1	202689 (100%)	Mineral	Claim	104B068	1972/may/31	2014/dec/15	GOOD	25.0
255255	S.I.B. #2	202689 (100%)	Mineral	Claim	104B068	1972/may/31	2014/dec/15	GOOD	25.0
255256	S.I.B. #3	202689 (100%)	Mineral	Claim	104B068	1982/may/31	2014/dec/15	GOOD	25.0
255257	S.I.B. #4	202689 (100%)	Mineral	Claim	104B068	1972/may/31	2014/dec/15	GOOD	25.0
304070	RAMBO 1	202689 (100%)	Mineral	Claim	104B058	1991/sep/09	2014/jan/31	GOOD	25.0
304072	RAMBO 3	202689 (100%)	Mineral	Claim	104B058	1991/sep/09	2014/jan/31	GOOD	25.0
304074	RAMBO 5	202689 (100%)	Mineral	Claim	104B058	1991/sep/09	2014/jan/31	GOOD	25.0
305317	FOG 1	202689 (100%)	Mineral	Claim	104B058	1991/oct/05	2014/jan/31	GOOD	25.0
305318	FOG 2	202689 (100%)	Mineral	Claim	104B058	1991/oct/05	2014/jan/31	GOOD	25.0
305319	FOG 3	202689 (100%)	Mineral	Claim	104B058	1991/oct/05	2014/jan/31	GOOD	25.0
305320	FOG 4	202689 (100%)	Mineral	Claim	104B058	1991/oct/05	2014/jan/31	GOOD	25.0
305321	FOG 5	202689 (100%)	Mineral	Claim	104B058	1991/oct/05	2014/jan/31	GOOD	25.0
305322	FOG 6	202689 (100%)	Mineral	Claim	104B058	1991/oct/05	2014/jan/31	GOOD	25.0

Tenure Number	Claim Name	Owner	Tenure Type	Tenure Sub Type	Map Number	Issue Date	Good To Date	Status	Area (ha)
306723	NOOT 1	202689 (100%)	Mineral	Claim	104B058	1991/nov/29	2014/jan/31	GOOD	500.0
306724	NOOT 2	202689 (100%)	Mineral	Claim	104B058	1991/nov/29	2014/jan/31	GOOD	500.0
306725	NOOT 3	202689 (100%)	Mineral	Claim	104B058	1991/nov/29	2014/jan/31	GOOD	500.0
311923	LINK FR	202689 (100%)	Mineral	Claim	104B058	1992/jul/24	2014/jan/31	GOOD	25.0
313285	CALVIN	202689 (100%)	Mineral	Claim	104B069	1992/sep/17	2014/jan/31	GOOD	500.0
329001		202689 (100%)	Mineral	Lease	104B068	1996/sep/06	2011/sep/06	GOOD	823.0
367934	PUD 1	202689 (100%)	Mineral	Claim	104B058	1999/feb/25	2014/jan/31	GOOD	500.0
367935	PUD 2	202689 (100%)	Mineral	Claim	104B058	1999/feb/25	2014/jan/31	GOOD	100.0
367943	MEGAN 1	202689 (100%)	Mineral	Claim	104B058	1999/feb/25	2015/jan/31	GOOD	25.0
367944	MEGAN 2	202689 (100%)	Mineral	Claim	104B058	1999/feb/25	2015/jan/31	GOOD	25.0
373867	STO 2	202689 (100%)	Mineral	Claim	104B058	1999/dec/15	2014/jan/31	GOOD	125.0
384019	JOHN 1	202689 (100%)	Mineral	Claim	104B058	2001/feb/12	2014/jan/31	GOOD	400.0
384020	JOHN 2	202689 (100%)	Mineral	Claim	104B058	2001/feb/12	2014/jan/31	GOOD	400.0
387231	IRVING 1	202689 (100%)	Mineral	Claim	104B069	2001/jun/04	2014/jan/31	GOOD	500.0
387233	IRVING 3	202689 (100%)	Mineral	Claim	104B069	2001/jun/04	2014/jan/31	GOOD	500.0
387237	BELL 1	202689 (100%)	Mineral	Claim	104B069	2001/jun/04	2014/jan/31	GOOD	500.0
387238	BELL 2	202689 (100%)	Mineral	Claim	104B069	2001/jun/04	2014/jan/31	GOOD	500.0
387239	BELL 3	202689 (100%)	Mineral	Claim	104B069	2001/jun/04	2014/jan/31	GOOD	375.0
387240	BELL 4	202689 (100%)	Mineral	Claim	104B069	2001/jun/04	2014/jan/31	GOOD	500.0
387241	BELL 5	202689 (100%)	Mineral	Claim	104B069	2001/jun/04	2014/jan/31	GOOD	200.0
387245	BELL 6	202689 (100%)	Mineral	Claim	104B069	2001/jun/04	2014/jan/31	GOOD	250.0
387248	BELL 7	202689 (100%)	Mineral	Claim	104B069	2001/jun/04	2014/jan/31	GOOD	175.0
387249	BELL 8	202689 (100%)	Mineral	Claim	104B069	2001/jun/04	2014/jan/31	GOOD	125.0
389463	TOON 1	202689 (100%)	Mineral	Claim	104B058	2001/sep/10	2014/jan/31	GOOD	50.0
389464	TOON 2	202689 (100%)	Mineral	Claim	104B058	2001/sep/10	2014/jan/31	GOOD	300.0
390911	HARRY 1	202689 (100%)	Mineral	Claim	104B058	2001/nov/16	2014/jan/31	GOOD	500.0
390912	HARRY 2	202689 (100%)	Mineral	Claim	104B058	2001/nov/16	2014/jan/31	GOOD	375.0
390913	HARRY 3	202689 (100%)	Mineral	Claim	104B058	2001/nov/16	2014/jan/31	GOOD	500.0
390914	SC 1	202689 (100%)	Mineral	Claim	104B049	2001/nov/16	2014/jan/31	GOOD	500.0
390915	SC 2	202689 (100%)	Mineral	Claim	104B049	2001/nov/16	2014/jan/31	GOOD	500.0
390916	SC 3	202689 (100%)	Mineral	Claim	104B049	2001/nov/16	2014/jan/31	GOOD	500.0
390917	SC 4	202689 (100%)	Mineral	Claim	104B049	2001/nov/16	2014/jan/31	GOOD	500.0
390918	SC 5	202689 (100%)	Mineral	Claim	104B049	2001/nov/16	2014/jan/31	GOOD	500.0
390919	SC 6	202689 (100%)	Mineral	Claim	104B049	2001/nov/16	2014/jan/31	GOOD	500.0
390920	SC 7	202689 (100%)	Mineral	Claim	104B049	2001/nov/16	2014/jan/31	GOOD	500.0
390921	SC 8	202689 (100%)	Mineral	Claim	104B049	2001/nov/16	2014/jan/31	GOOD	500.0
392425	HARRY 4	202689 (100%)	Mineral	Claim	104B058	2002/mar/22	2014/jan/31	GOOD	500.0
392426	HARRY 5	202689 (100%)	Mineral	Claim	104B058	2002/mar/22	2014/jan/31	GOOD	100.0
392427	KING 1	202689 (100%)	Mineral	Claim	104B068	2002/mar/22	2012/jan/31	GOOD	75.0
392428	KING 2	202689 (100%)	Mineral	Claim	104B058	2002/mar/22	2014/jan/31	GOOD	400.0
392429	KING 3	202689 (100%)	Mineral	Claim	104B058	2002/mar/22	2012/jan/31	GOOD	450.0
392430	KING 4	202689 (100%)	Mineral	Claim	104B058	2002/mar/22	2014/jan/31	GOOD	450.0
392431	KING 5	202689 (100%)	Mineral	Claim	104B058	2002/mar/22	2014/jan/31	GOOD	450.0
392432	KING 6	202689 (100%)	Mineral	Claim	104B057	2002/mar/22	2014/jan/31	GOOD	300.0
392433	KING 7	202689 (100%)	Mineral	Claim	104B057	2002/mar/22	2014/jan/31	GOOD	450.0
392438	TC 13	202689 (100%)	Mineral	Claim	104B060	2002/mar/21	2014/jan/31	GOOD	500.0
392439	TC 14	202689 (100%)	Mineral	Claim	104B060	2002/mar/21	2014/jan/31	GOOD	500.0
392440	VALCANO 1	202689 (100%)	Mineral	Claim	104B068	2002/mar/22	2014/jan/31	GOOD	450.0
392441	VALCANO 2	202689 (100%)	Mineral	Claim	104B078	2002/mar/22	2014/jan/31	GOOD	450.0
392442	VALCANO 3	202689 (100%)	Mineral	Claim	104B078	2002/mar/22	2014/jan/31	GOOD	400.0

Tenure Number	Claim Name	Owner	Tenure Type	Tenure Sub Type	Map Number	Issue Date	Good To Date	Status	Area (ha)
392443	VALCANO 4	202689 (100%)	Mineral	Claim	104B078	2002/mar/22	2014/jan/31	GOOD	400.0
392444	VALCANO 5	202689 (100%)	Mineral	Claim	104B078	2002/mar/23	2014/jan/31	GOOD	225.0
392445	VALCANO 6	202689 (100%)	Mineral	Claim	104B077	2002/mar/23	2014/jan/31	GOOD	450.0
392446	VALCANO 7	202689 (100%)	Mineral	Claim	104B077	2002/mar/23	2014/jan/31	GOOD	450.0
392447	VALCANO 8	202689 (100%)	Mineral	Claim	104B068	2002/mar/22	2014/jan/31	GOOD	400.0
392448	VALCANO 9	202689 (100%)	Mineral	Claim	104B068	2002/mar/22	2014/jan/31	GOOD	400.0
392449	CALVIN 2	202689 (100%)	Mineral	Claim	104B069	2002/mar/23	2014/jan/31	GOOD	350.0
392450	CALVIN 3	202689 (100%)	Mineral	Claim	104B069	2002/mar/23	2014/jan/31	GOOD	350.0
392451	CALVIN 4	202689 (100%)	Mineral	Claim	104B069	2002/mar/23	2014/jan/31	GOOD	250.0
392452	CALVIN 5	202689 (100%)	Mineral	Claim	104B069	2002/mar/23	2014/jan/31	GOOD	500.0
392453	GINGRASS 1	202689 (100%)	Mineral	Claim	104B059	2002/mar/21	2014/jan/31	GOOD	150.0
392454	GINGRASS 2	202689 (100%)	Mineral	Claim	104B059	2002/mar/21	2014/jan/31	GOOD	500.0
392455	GINGRASS 3	202689 (100%)	Mineral	Claim	104B059	2002/mar/21	2014/jan/31	GOOD	300.0
392456	GINGRASS 4	202689 (100%)	Mineral	Claim	104B059	2002/mar/21	2014/jan/31	GOOD	225.0
392457	GINGRASS 5	202689 (100%)	Mineral	Claim	104B059	2002/mar/21	2014/jan/31	GOOD	300.0
392458	IRVING 5	202689 (100%)	Mineral	Claim	104B069	2002/mar/23	2014/jan/31	GOOD	225.0
392459	IRVING 6	202689 (100%)	Mineral	Claim	104B069	2002/mar/23	2014/jan/31	GOOD	450.0
394157	LANCE 5	202689 (100%)	Mineral	Claim	104B069	2002/jun/09	2014/jan/31	GOOD	150.0
394158	MEGAN 3	202689 (100%)	Mineral	Claim	104B058	2002/jun/09	2014/jan/31	GOOD	100.0
394159	MEGAN 4	202689 (100%)	Mineral	Claim	104B058	2002/jun/08	2014/jan/31	GOOD	75.0
394160	SKI	202689 (100%)	Mineral	Claim	104B068	2002/jun/09	2014/jan/31	GOOD	125.0
394161	DWAYNE 2	202689 (100%)	Mineral	Claim	104B058	2002/jun/08	2014/jan/31	GOOD	175.0
394162	AFT	202689 (100%)	Mineral	Claim	104B068	2002/jun/09	2014/jan/31	GOOD	50.0
394163	SHIRLEY	202689 (100%)	Mineral	Claim	104B068	2002/jun/09	2014/jan/31	GOOD	75.0
394164	FREDDY 1	202689 (100%)	Mineral	Claim	104B068	2002/jun/09	2014/jan/31	GOOD	75.0
394165	FREDDY 2	202689 (100%)	Mineral	Claim	104B058	2002/jun/09	2014/jan/31	GOOD	75.0
404668	SUL 1	202689 (100%)	Mineral	Claim	104B048	2003/aug/07	2014/jan/31	GOOD	500.0
404669	SUL 2	202689 (100%)	Mineral	Claim	104B048	2003/aug/07	2014/jan/31	GOOD	500.0
527171	KUT M	202689 (100%)	Mineral	Claim	104B	2006/feb/06	2016/jan/31	GOOD	231.619
527172		202689 (100%)	Mineral	Claim	104B	2006/feb/06	2016/jan/31	GOOD	17.813
527177		202689 (100%)	Mineral	Claim	104B	2006/feb/06	2014/jan/31	GOOD	320.871
527180		202689 (100%)	Mineral	Claim	104B	2006/feb/06	2016/jan/31	GOOD	35.623
527241		202689 (100%)	Mineral	Claim	104B	2006/feb/07	2014/jan/31	GOOD	178.253
528422	SIB FIXUP 1	202689 (100%)	Mineral	Claim	104B	2006/feb/16	2014/jan/31	GOOD	284.541
528661		202689 (100%)	Mineral	Claim	104B	2006/feb/20	2014/jan/31	GOOD	142.474
528664	SIB FIXUP 2	202689 (100%)	Mineral	Claim	104B	2006/feb/20	2014/jan/31	GOOD	35.617
528665	SIB FIXUP 3	202689 (100%)	Mineral	Claim	104B	2006/feb/20	2014/jan/31	GOOD	17.806
528666	ST ANDREW 2	202689 (100%)	Mineral	Claim	104B	2006/feb/20	2014/jan/31	GOOD	17.815
530906	ST ANDREW 1	202689 (100%)	Mineral	Claim	104B	2006/mar/31	2011/oct/15	GOOD	178.103
530907	ST ANDREW 3	202689 (100%)	Mineral	Claim	104B	2006/mar/31	2011/oct/15	GOOD	249.275
530910	ST ANDREW 1	202689 (100%)	Mineral	Claim	104B	2006/mar/31	2011/jan/29	GOOD	17.807
541059	ST ANDREW 2	202689 (100%)	Mineral	Claim	104B	2006/sep/11	2014/jan/31	GOOD	17.8115
566735	ST ANDREW 3	202689 (100%)	Mineral	Claim	104B	2007/sep/26	2014/jan/31	GOOD	17.8115
566739	ST ANDREW 4	202689 (100%)	Mineral	Claim	104B	2007/sep/26	2014/jan/31	GOOD	160.6031
566751	ST ANDREW 1	202689 (100%)	Mineral	Claim	104B	2007/sep/26	2014/jan/31	GOOD	249.777
566752	ST ANDREW 2	202689 (100%)	Mineral	Claim	104B	2007/sep/26	2014/jan/31	GOOD	17.8468
566752	ST ANDREW 3	202689 (100%)	Mineral	Claim	104B	2007/sep/26	2014/jan/31	GOOD	17.8373
566844	ST ANDREW 4	202689 (100%)	Mineral	Claim	104B	2007/sep/26	2014/jan/31	GOOD	17.8373
566845	KHBER PASS 1	202689 (100%)	Mineral	Claim	104B	2007/sep/27	2011/jan/29	GOOD	106.892
566845	KHBER PASS 2	202689 (100%)	Mineral	Claim	104B	2007/sep/27	2011/oct/15	GOOD	320.668
570110	KUT ABC	202689 (100%)	Mineral	Claim	104B	2007/nov/15	2011/jan/29	GOOD	658.237
570253	ST ANDREW 1	202689 (100%)	Mineral	Claim	104B	2007/nov/19	2011/jan/29	GOOD	177.679

Tenure Number	Claim Name	Owner	Tenure Type	Tenure Sub Type	Map Number	Issue Date	Good To Date	Status	Area (ha)
570254	ST ANDREW 2	202689 (100%)	Mineral	Claim	104B	2007/nov/19	2011/jan/29	GOOD	266.567
570258	ST ANDREW 3	202689 (100%)	Mineral	Claim	104B	2007/nov/19	2011/apr/29	GOOD	568.871
570937	INEL WEST 1	202689 (100%)	Mineral	Claim	104B	2007/nov/28	2011/jan/29	GOOD	284.689
598285		202689 (100%)	Mineral	Claim	104B	2009/feb/01	2011/oct/15	GOOD	444.9753
598286		202689 (100%)	Mineral	Claim	104B	2009/feb/01	2011/jan/29	GOOD	427.1677
598292		202689 (100%)	Mineral	Claim	104B	2009/feb/01	2011/jan/29	GOOD	444.9872
598293		202689 (100%)	Mineral	Claim	104B	2009/feb/01	2011/jan/29	GOOD	409.5644
598294		202689 (100%)	Mineral	Claim	104B	2009/feb/01	2011/jan/29	GOOD	267.082
598300		202689 (100%)	Mineral	Claim	104B	2009/feb/01	2011/jan/29	GOOD	35.5915
598666	RESURRECTION (	202689 (100%)	Mineral	Claim	104B	2009/feb/03	2011/jan/29	GOOD	17.7958
600290	KHYBER PASS 3	202689 (100%)	Mineral	Claim	104B	2009/mar/03	2011/oct/15	GOOD	356.3848
663823	KHYBER PASS 4	202689 (100%)	Mineral	Claim	104B	2009/nov/02	2011/jan/29	GOOD	427.7448
663824	KHYBER PASS 5	202689 (100%)	Mineral	Claim	104B	2009/nov/02	2011/jan/29	GOOD	445.5524
831390	ST ANDREW 5	202689 (100%)	Mineral	Claim	104B	2010/aug/12	2011/aug/12	GOOD	284.4663
831393	ST ANDREW 5	202689 (100%)	Mineral	Claim	104B	2010/aug/12	2011/aug/12	GOOD	284.5923
831397	ST ANDREW 7	202689 (100%)	Mineral	Claim	104B	2010/aug/12	2011/aug/12	GOOD	71.1714
833742	NEW SNIP 1	202689 (100%)	Mineral	Claim	104B	2010/sep/16	2011/sep/16	GOOD	445.2922
833743	NEW SNIP 2	202689 (100%)	Mineral	Claim	104B	2010/sep/16	2011/sep/16	GOOD	356.129
833744	NEW SNIP 3	202689 (100%)	Mineral	Claim	104B	2010/sep/16	2011/sep/16	GOOD	374.1421
834157	NEW SNIP 4	202689 (100%)	Mineral	Claim	104B	2010/sep/23	2011/sep/23	GOOD	71.2607
834369	NEW SNIP 5	202689 (100%)	Mineral	Claim	104B	2010/sep/27	2011/sep/27	GOOD	142.5995
834370	NEW SNIP 5	202689 (100%)	Mineral	Claim	104B	2010/sep/27	2011/sep/27	GOOD	356.5346

APPENDIX D

Lithology Codes (used in cross sections, Figs. 4-7)

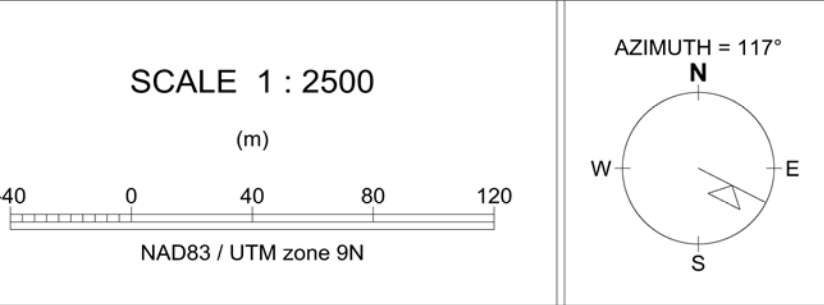
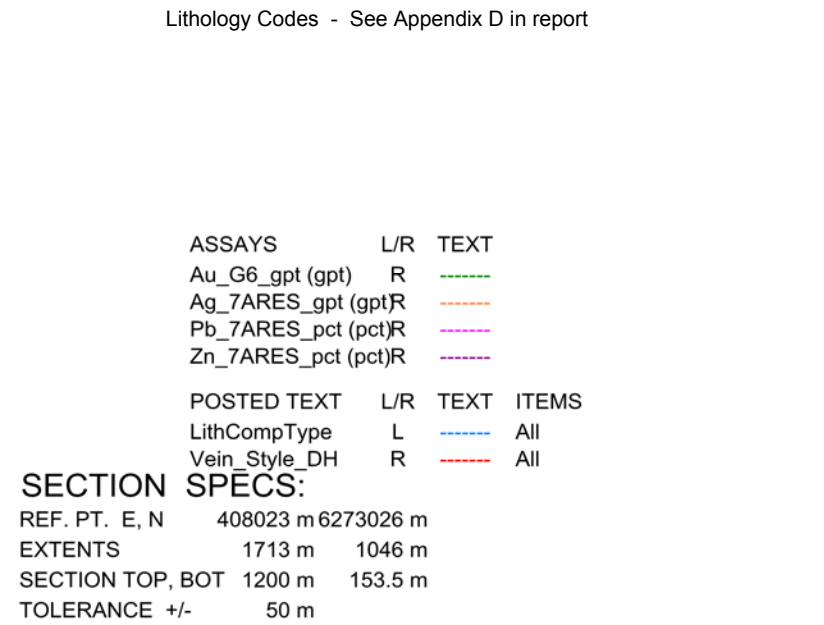
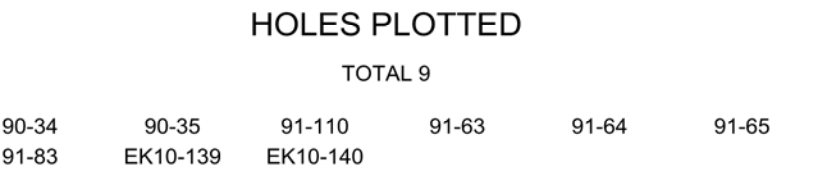
Lithology Rock Codes

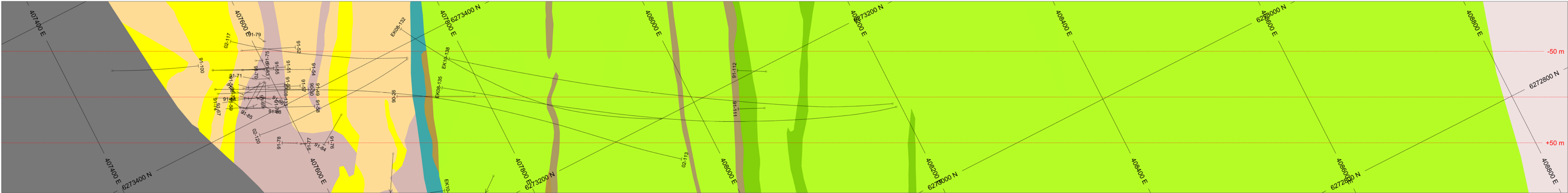
Lithological Composition = 1st two letters of rock code (XX--)

BS	Basalt	VO	Volcanic	DH	Drillhole
MF	Mafic				
AN	Andesite	DI	Diorite	ST	Structure
IN	Intermediate				
DC	Dacite	SD	Sedimentary rock	AX	Altered
RD	Rhyodacite	LI	Lithic		
RH	Rhyolite	CH	Chert		
FS	Felsic				

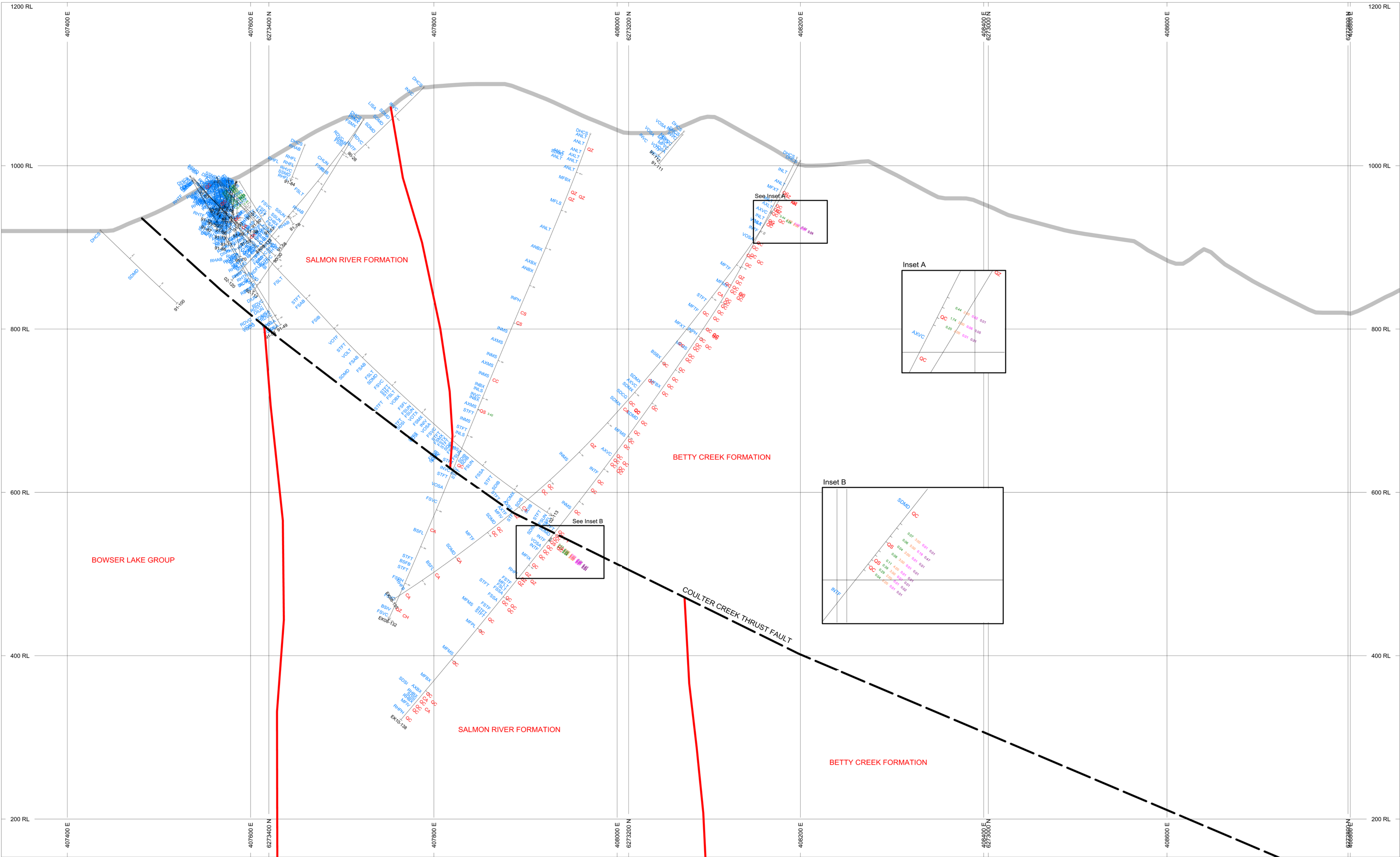
Lithological Type = 2nd two letters of rock code (--XX)

AB	autobreccia	CG	conglomerate	OB	overburden
BX	breccia	MD	mudstone	CS	casing
FB	flow breccia	SI	siltstone		
FL	flow	SA	sandstone		
IV	intrusion				
LS	lapillistone	IB	interbedded	FT	fault
LT	lapilli tuff	MS	massive		
PH	porphyry	MX	mixed		
PL	pillowed				
PB	pillow breccia	UN	undifferentiated		
TF	tuff				
VC	volcaniclastic				





HOLES PLOTTED					
TOTAL 37					
02-113	02-117	02-120	90-26	90-30	91-100
91-111	91-112	91-49	91-50	91-51	91-53
91-54	91-55	91-56	91-57	91-58	91-59
91-67	91-68	91-69	91-70	91-71	91-72
91-73	91-74	91-75	91-76	91-84	91-85
91-86	91-87	91-88	EK08-132	EK08-133	EK08-135
EK10-138					

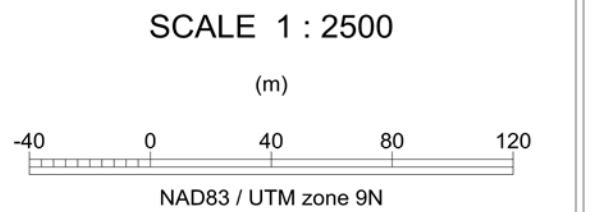


Lithology Codes - See Appendix D in report

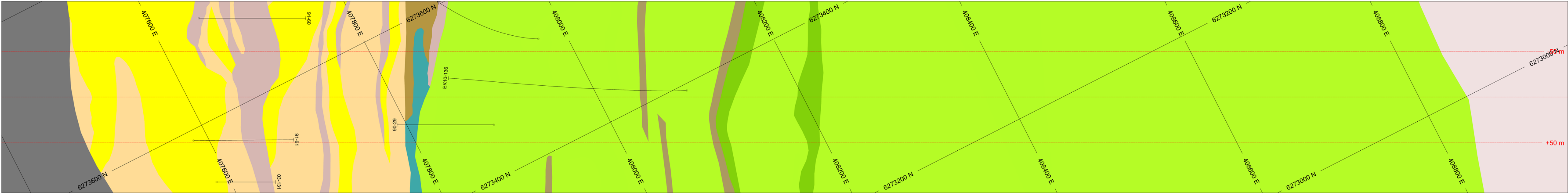
ASSAYS		L/R	TEXT	
Au_G6_gpt	(gpt)	R	-----	
Ag_7ARES_gpt	(gpt)	R	-----	
Pb_7ARES_pct	(pct)	R	-----	
Zn_7ARES_pct	(pct)	R	-----	
POSTED TEXT		L/R	TEXT	ITEMS
LithCompType	L	-----	-----	All
Vein_Style_DH	R	-----	-----	All

SECTION SPECS:

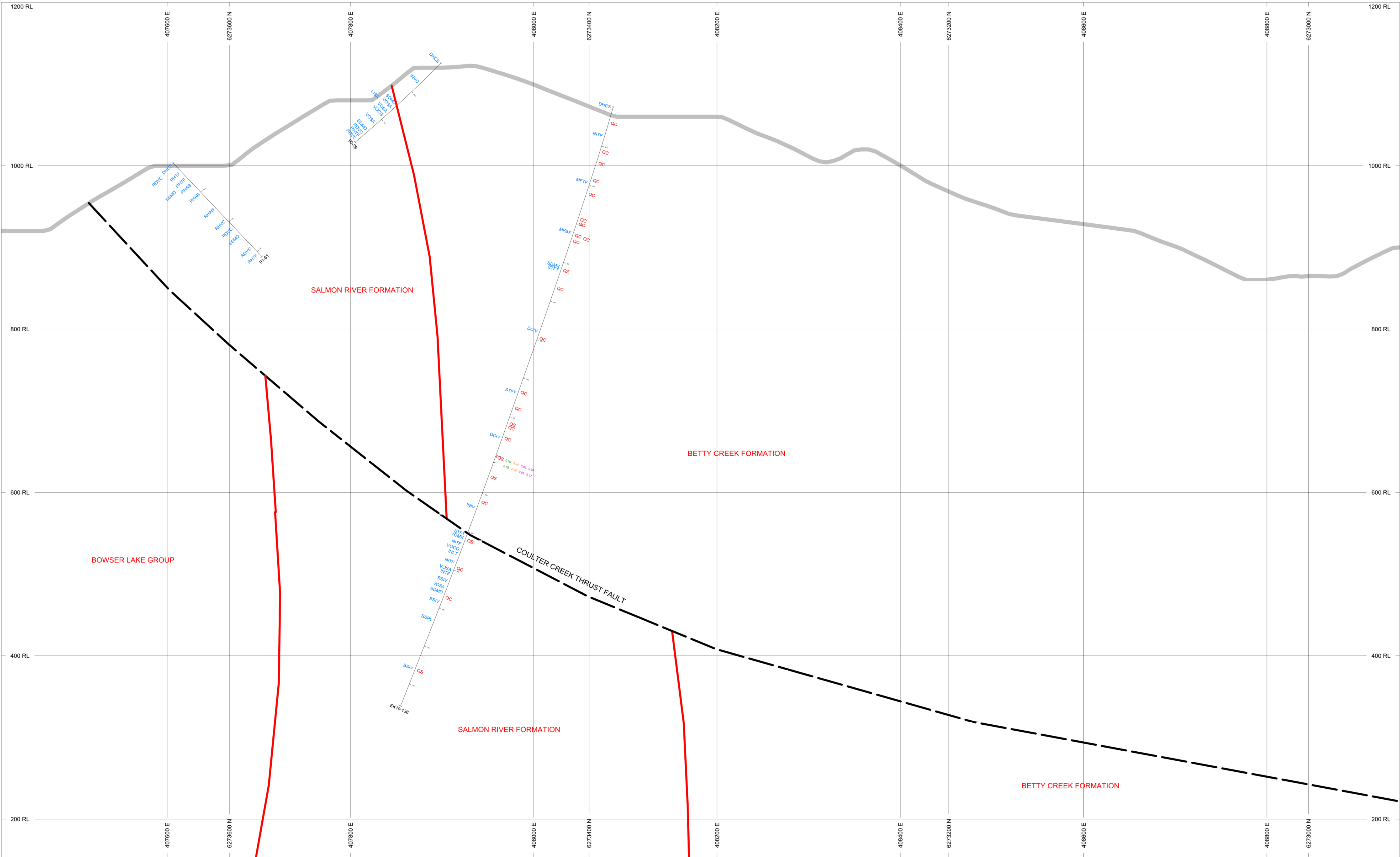
REF. PT. E, N	408091 m	6273160 m
EXTENTS	1713 m	1046 m
SECTION TOP, BOT	1200 m	153.5 m
TOLERANCE +/-	50 m	



Eskay Mining Corporation
Eskay-SIB Property
Section 2 (Fig.5)
Cambria Geosciences Inc. 2010/11/30



HOLES PLOTTED
TOTAL 3
90-29 91-61 EK10-136

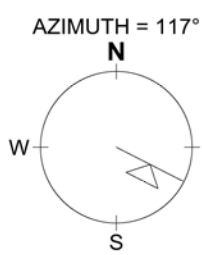
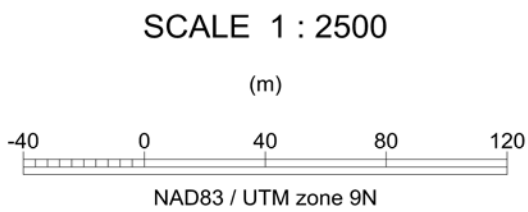


Lithology Codes - See Appendix D in report

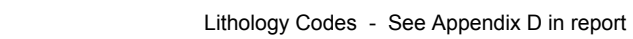
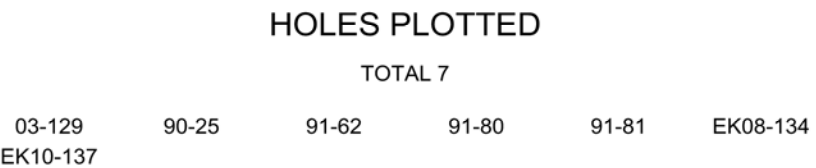
ASSAYS		L/R	TEXT	
Au_G6_gpt (gpt)	R	-----		
Ag_7ARES_gpt (gpt)	R	-----		
Pb_7ARES_pct (pct)	R	-----		
Zn_7ARES_pct (pct)	R	-----		
POSTED TEXT		L/R	TEXT	ITEMS
LithCompType	L	-----	All	
Vein_Style_DH	R	-----	All	

SECTION SPECS:

REF. PT. E, N	408162 m	6273338 m
EXTENTS	1713 m	1046 m
SECTION TOP, BOT	1200 m	153.5 m
TOLERANCE +/-	50 m	



Eskay Mining Corporation
Eskay-SIB Property
Section 3 (Fig.6)
Cambria Geosciences Inc. 2010/11/30

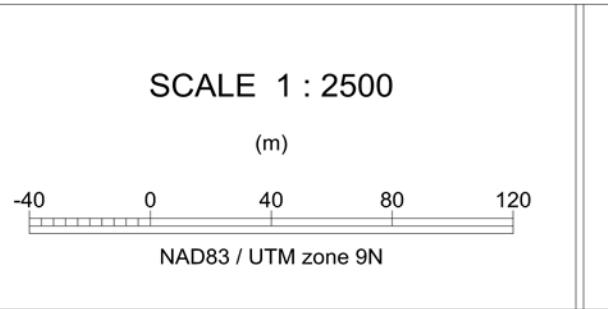


ASSAYS	L/R	TEXT
Au_G6_gpt (gpt)	R	-----
Ag_7ARES_gpt (gpt)	R	-----
Pb_7ARES_pct (pct)	R	-----
Zn_7ARES_pct (pct)	R	-----

	POSTED TEXT	L/R	TEXT	ITEMS
	LithCompType	L	-----	All
	Vein_Style_DH	R	-----	All

SECTION SPECS:

REF. PT. E, N	408273 m	6273517 m
EXTENTS	1713 m	1046 m
SECTION TOP, BOT	1200 m	153.5 m
TOLERANCE +/-	50 m	



Eskay Mining Corporation
Eskay-SIB Property
Section 4 (Fig.7)
Cambria Geosciences Inc. 2010/11/30

Lithological Legend

Eskay (SIB) Property, Northwest, British Columbia

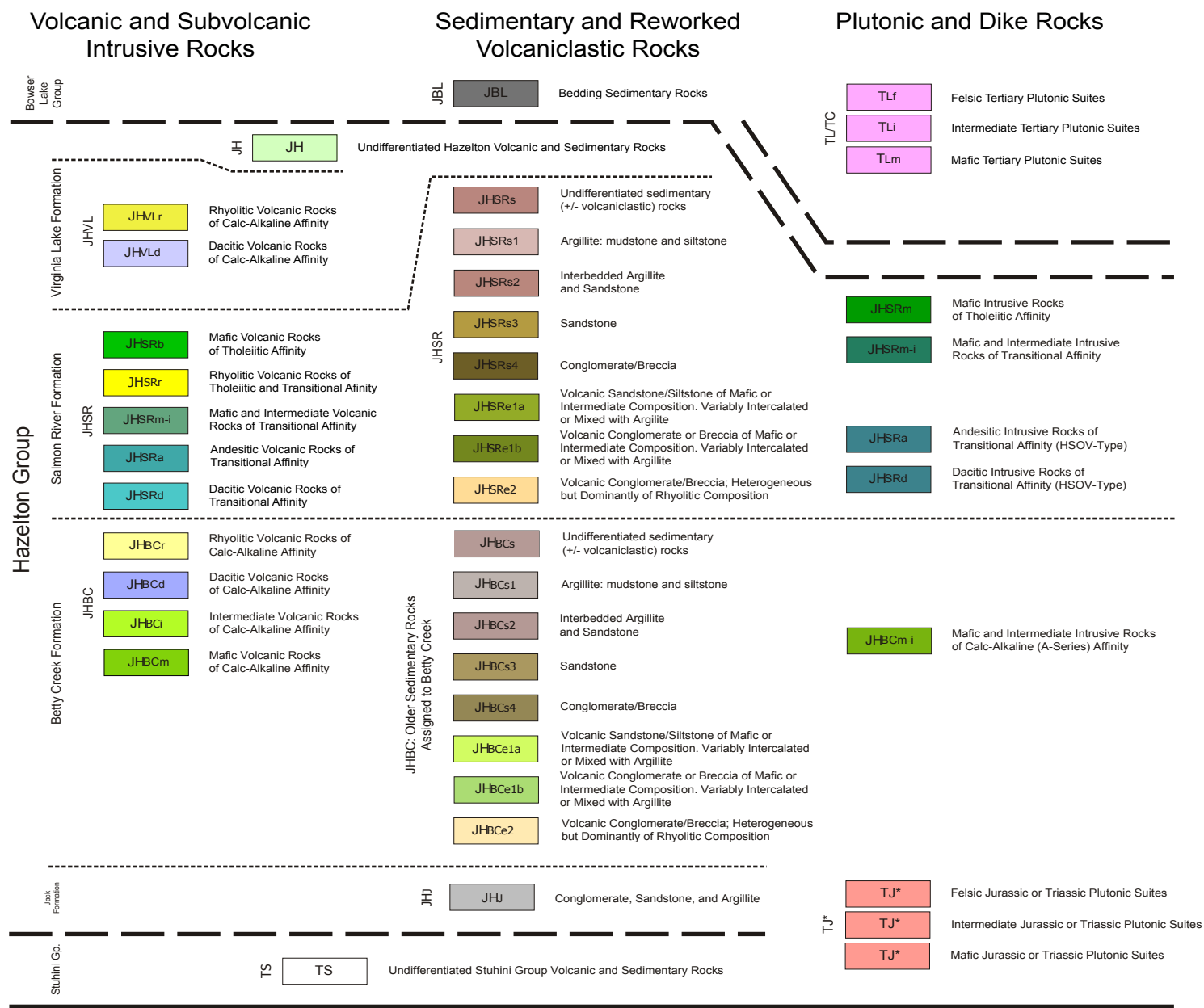
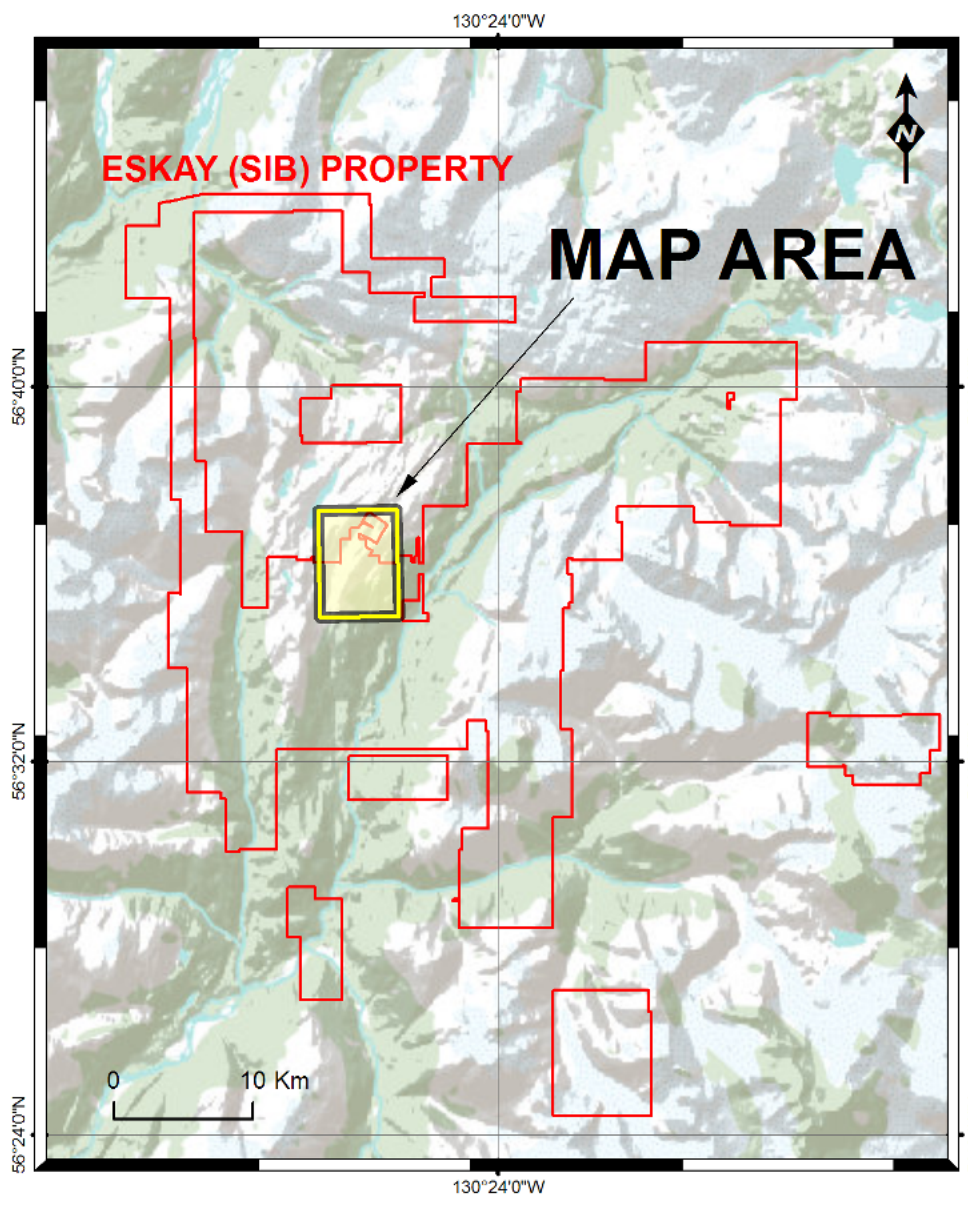
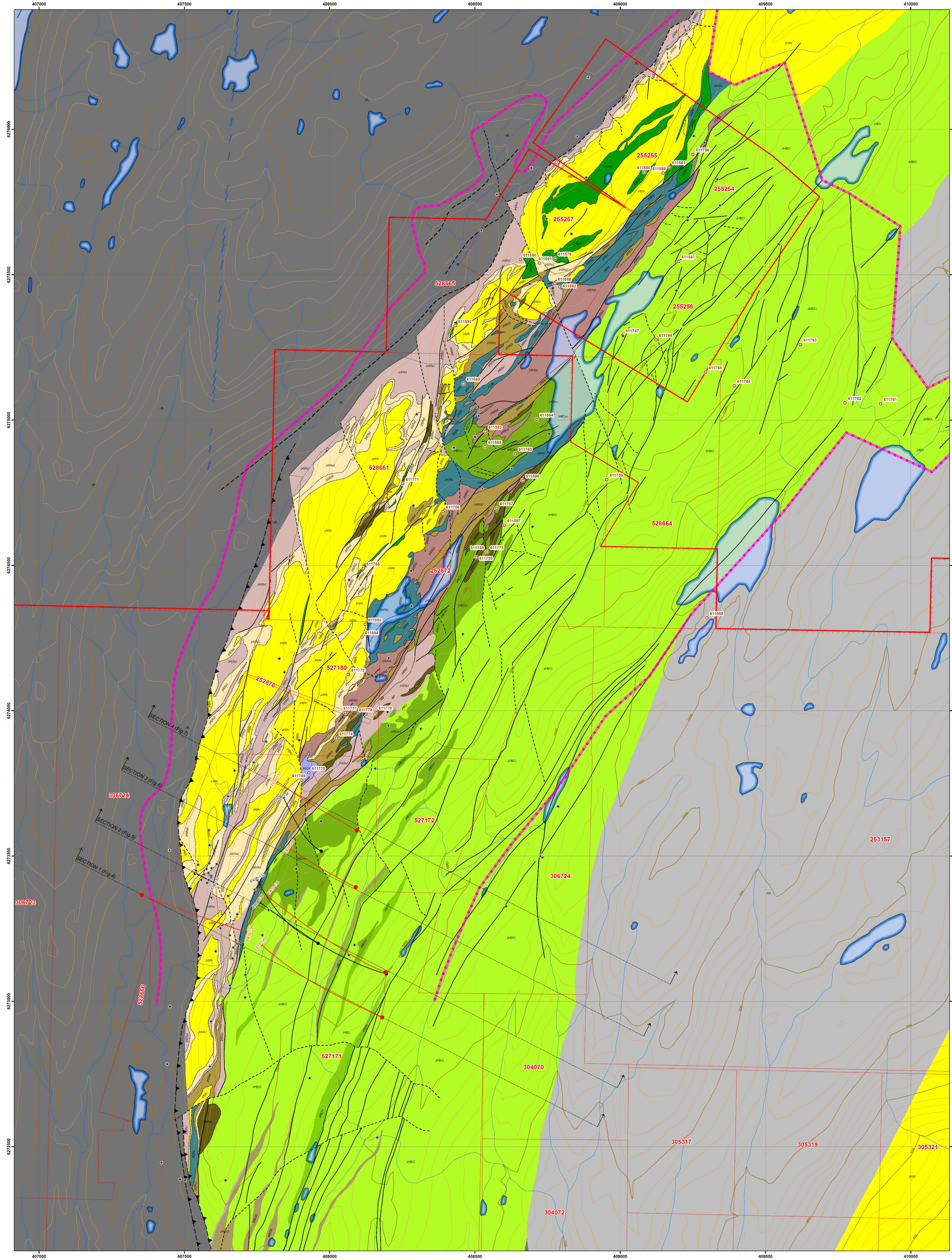
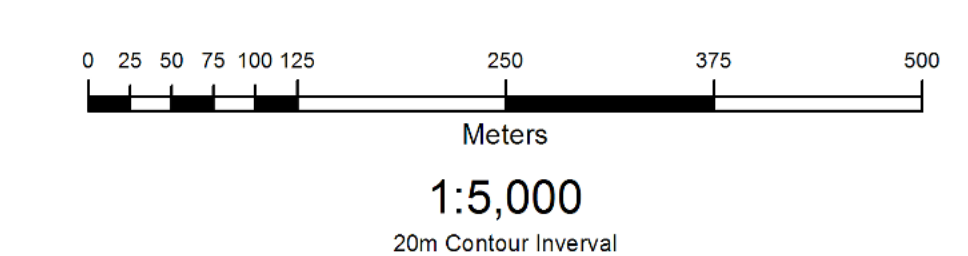


Fig. 8



- Legend**
- 2010 Drilling
 - 2008 Drilling
 - Pre 2008 Drill Collar
 - Lithochemical Sample Location
 - Section Line
 - EsKay (SIB) Claim Boundary
 - Tenure Boundary
 - Lithological Contact
 - Boundaries Of Mapping
 - Faults**
 - Thrust Fault
 - Argillite Creek Fault
 - Shear
 - Fault - General
 - Cross Fault
 - River
 - Lake
 - Index Contour
 - Intermediate Contour

Lithological Legend
See Fig. 8 in report



ESKAY MINING CORP.	
ESKAY (SIB) PROPERTY Northwest British Columbia, Canada	
GEOLOGY	
Map Author: D. Melvett	Cambria Geosciences Inc.
Datum: NAD83	Date: 2010-12-08
Projection: UTM - 09	Map: 1

